

1 ***Candidatus* Ethanoperedens, a thermophilic genus of archaea mediating the**
2 **anaerobic oxidation of ethane**

3 **Running title: Ethane oxidation by “*Ca. Ethanoperedens*”**

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14 **ABSTRACT:** Cold seeps and hydrothermal vents deliver large amounts of methane and other
15 gaseous alkanes into marine surface sediments. Consortia of archaea and partner bacteria
16 thrive on the oxidation of these alkanes and its coupling to sulfate reduction. The inherently
17 slow growth of the involved organisms and the lack of pure cultures have impeded the
18 understanding of the molecular mechanisms of archaeal alkane degradation. Here, using
19 hydrothermal sediments of the Guaymas Basin (Gulf of California) and ethane as substrate we
20 cultured microbial consortia of a novel anaerobic ethane oxidizer *Candidatus* Ethanoperedens
21 thermophilum (GoM-Arc1 clade) and its partner bacterium *Candidatus* Desulfosphaerulum
22 auxilii previously known from methane-oxidizing consortia. The sulfate reduction activity of
23 the culture doubled within one week, indicating a much faster growth than in any other
24 alkane-oxidizing archaea described before. The dominance of a single archaeal phylotype in

25 this culture allowed retrieving a closed genome of *Ca. Ethanoperedens*, a sister genus of the
26 recently reported ethane oxidizer *Candidatus Argoarchaeum*. The metagenome-assembled
27 genome of *Ca. Ethanoperedens* encoded for a complete methanogenesis pathway including a
28 methyl-coenzyme M reductase (MCR) that is highly divergent from those of methanogens
29 and methanotrophs. Combined substrate and metabolite analysis showed ethane as sole
30 growth substrate and production of ethyl-coenzyme M as activation product. Stable isotope
31 probing showed that the enzymatic mechanisms of ethane oxidation in *Ca. Ethanoperedens* is
32 fully reversible, thus its enzymatic machinery has potential for the biotechnological
33 development of microbial ethane production from carbon dioxide.

34

35 **IMPORTANCE:** In the seabed gaseous alkanes are oxidized by syntrophic microbial
36 consortia that thereby reduce fluxes of these compounds into the water column. Because of
37 the immense quantities of seabed alkane fluxes, these consortia are key catalysts of the global
38 carbon cycle. Due to their obligate syntrophic lifestyle, the physiology of alkane-degrading
39 archaea remains poorly understood. We have now cultivated a thermophilic, relatively fast-
40 growing ethane oxidizer in partnership with a sulfate-reducing bacterium known to aid in
41 methane oxidation, and have retrieved the first complete genome of a short-chain alkane-
42 degrading archaeon. This will greatly enhance the understanding of non-methane alkane
43 activation by non-canonical methyl-coenzyme M reductase enzymes, and provide insights
44 into additional metabolic steps and the mechanisms underlying syntrophic partnerships.
45 Ultimately, this knowledge could lead to the biotechnological development of alkanogenic
46 microorganisms to support the carbon neutrality of industrial processes.

47

48 **KEYWORDS:** Alkane degradation, archaea, syntrophy, methyl-coenzyme M reductase,
49 model organism, hydrothermal vents

50 **Etymology.** *Ethanoperedens. ethano*, (new Latin): pertaining to ethane; *peredens* (Latin):
51 consuming, devouring; *thermophilum*. (Greek): heat-loving. The name implies an organism
52 capable of ethane oxidation at elevated temperatures.

53 **Locality.** Enriched from hydrothermally heated, hydrocarbon-rich marine sediment of the
54 Guaymas Basin at 2000 m water depth, Gulf of California, Mexico.

55 **Diagnosis.** Anaerobic, ethane-oxidizing archaeon, mostly coccoid, about 0.7 μm in diameter,
56 forms large irregular cluster in large dual-species consortia with the sulfate- reducing partner
57 bacterium '*Candidatus Desulfosphaerium auxilii*'.

58

59 INTRODUCTION

60 In deep marine sediments, organic matter undergoes thermocatalytic decay, resulting
61 in the formation of natural gas (methane to butane) and crude oil. If not capped, the gas
62 fraction will rise towards the sediment surface due to buoyancy, porewater discharge and
63 diffusion. Most of the gas is oxidized within the sediments coupled to the reduction of the
64 abundant electron acceptor sulfate [1, 2]. Responsible for the anaerobic oxidation of alkanes
65 are either free-living bacteria or microbial consortia of archaea and bacteria. Most free-living
66 bacteria use alkyl succinate synthases to activate the alkane, forming succinate-bound alkyl
67 units as primary intermediates [3]. Usually, these alkanes are completely oxidized, and this
68 process is coupled to sulfate reduction in the same cells, as has been shown, for example in
69 the deltaproteobacterial butane-degrading strain BuS5 [4]. However, alkane oxidation in
70 seafloor sediments is to a large extent performed by dual species consortia of archaea and
71 bacteria [5, 6]. As close relatives of methanogens, the archaea in those consortia activate

72 alkanes as thioethers and completely oxidize the substrates to CO₂. The electrons released
73 during alkane oxidation are consumed by the sulfate-reducing partner bacteria.

74 The anaerobic methane-oxidizing archaea (ANME) activate methane using methyl-
75 coenzyme M reductases (MCR) that are highly similar to those of methanogens, forming
76 methyl-coenzyme M as primary intermediate [7]. The methyl group is oxidized via a reversal
77 of the methanogenesis pathway [8]. Thermophilic archaea of the genus *Candidatus*
78 *Syntrophoarchaeum* thrive on the oxidation of butane and propane. In contrast to ANME, they
79 contain four highly divergent MCR variants, which generate butyl- and propyl-coenzyme M
80 (CoM) as primary intermediates [9]. Based on genomic and transcriptomic evidence the CoM-
81 bound alkyl units are transformed to fatty acids and oxidized further via beta-oxidation. The
82 reactions transforming the CoM-bound alkyl units to CoA-bound fatty acids and the enzymes
83 performing such reactions are so far unknown. The CoA-bound acetyl units are completely
84 oxidized in the Wood-Ljungdahl pathway including the upstream part of the methanogenesis
85 pathway. In hydrogenotrophic methanogens, the enzymes of this pathway are used to reduce
86 CO₂-forming methyl-tetrahydromethanopterin for methanogenesis and for biomass
87 production. In *Ca. Syntrophoarchaeum* this pathway is used in reverse direction for the
88 complete oxidation of acetyl-CoA. Both, the thermophilic ANME-1 and *Ca.*
89 *Syntrophoarchaeum* form dense consortia with their sulfate-reducing partner bacterium
90 *Candidatus Desulfobrevibacter* (HotSeep-1 clade) [10, 11]. The transfer of reducing equivalents
91 between the alkane oxidizing archaea and their partners is likely mediated by pili-based
92 nanowires and cytochromes produced by both consortial partners [12]. For a critical view on
93 electron transfer in AOM consortia see [13].

94 Sulfate-dependent ethane oxidation has been described multiple times in slurries of
95 marine sediments [4, 14, 15]. The first functional description of this process based on a cold-
96 adapted culture derived from Gulf of Mexico sediments [5]. In this culture, *Candidatus*

97 Argoarchaeum (formerly known as GoM-Arc1 clade) activates ethane with the help of
98 divergent MCRs that are phylogenetically placed on a distinct branch next to those of *Ca.*
99 Syntrophoarchaeum. Based on the presence of all enzymes of the Wood-Ljungdahl pathway
100 that can be used for acetyl-CoA oxidation, it has been suggested that the CoM-bound ethyl
101 groups are transferred to CoA-bound acetyl units. The required intermediates for this reaction
102 mechanism are so far unknown [5]. *Ca.* Argoarchaeum forms unstructured consortia with yet
103 unidentified bacterial partners and grows slowly with substrate turnover rates comparable to
104 AOM [5]. Additional metagenome assembled genomes (MAGs) of the GoM-Arc1 clade
105 derived from the Guaymas Basin and the Gulf of Mexico have similar gene contents,
106 suggesting that these GoM-Arc1 archaea are ethane oxidizers [16, 17].

107 To date, the understanding of short-chain alkane metabolizing archaea mainly relies
108 on comparison of their genomic information with those of methanogens that are well-
109 characterized with regard to their enzymes. Due to the slow growth of the alkane-oxidizing
110 archaea and the resulting lack of sufficient biomass, specific biochemical traits remain
111 unknown. For instance, the structural modifications of non-canonical MCRs, or the proposed
112 transformation of the CoM-bound alkyl to CoA-bound acetyl units in the short-chain alkane
113 degraders has not been proven. Here, we describe a faster growing, thermophilic ethane-
114 oxidizing culture from sediments of the Guaymas Basin. Metagenomic analyses of Guaymas
115 Basin sediments revealed a great diversity of potential alkane degraders with divergent MCR
116 enzymes [9, 18]. With ethane as sole energy source and sulfate as electron acceptor we
117 obtained well-growing meso- and thermophilic ethane-degrading enrichment cultures from
118 these sediments. Their low strain diversity makes them particularly suitable for assessing the
119 pathways of the anaerobic oxidation of ethane.

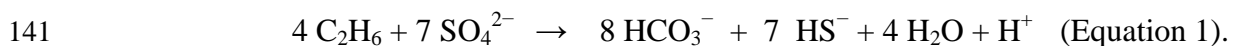
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121 **RESULTS AND DISCUSSION**

122 **Establishment of meso- and thermophilic ethane-oxidizing enrichment cultures.**

123 Sediments were sampled from the gas- and oil-rich sediments covered by sulfur-oxidizing
124 mats of the Guaymas Basin. From these sediments and artificial seawater medium, a slurry
125 was produced under anoxic conditions and distributed into replicate bottles. These bottles
126 were supplied with an ethane headspace (2 atm), and incubated at 37°C and 50°C. Additional
127 growth experiments were performed with methane and controls were set up with a nitrogen
128 atmosphere. As measure of metabolic activity, sulfide concentrations were tracked over time
129 (further details see methods). Both methane and ethane additions resulted in the formation of
130 15 mM sulfide within 4 months. Instead, nitrogen controls produced only little sulfide (<
131 2mM) that likely corresponds to the degradation of alkanes and organic matter from the
132 original sediment. Subsequent dilution (1:3) of the ethane and methane cultures and further
133 incubation with the corresponding substrates showed faster, exponentially increasing sulfide
134 production in the ethane culture, suggesting robust growth of the ethane-degrading
135 community (Fig. 1 A). After three consecutive dilution steps, virtually sediment-free cultures
136 were obtained. These cultures produced approximately 10 mM sulfide in 8 weeks. All further
137 experiments were conducted with the faster growing 50°C culture (Ethane50). Sequencing of
138 metagenomes, however, was done on both, the 50°C and 37°C (Ethane37) culture.

139 A stoichiometric growth experiment with the Ethane50 culture (Fig. 1B) showed that
140 ethane is completely oxidized while sulfate was reduced to sulfide according to the formula:



142 An experiment tracking the exponential development of sulfide over time suggested doubling
143 times of only 6 days at low sulfide concentrations of <5 mM (Fig. 1B), which is substantially
144 faster than estimated for thermophilic AOM consortia with about 60 days [10], and also faster
145 than the cold-adapted anaerobic ethane-oxidizing cultures [5]. Sulfide concentrations over
146 5 mM seemed to suppress activity and growth of the ethane-oxidizing microorganisms (Fig.

147 1C). Hence, flow-through bioreactors could be beneficial to increase biomass yields of
148 anaerobic ethane degraders.

149 **Microbial composition of the Ethane50 culture.** Amplified archaeal and bacterial 16S
150 rRNA genes of the original sediment and early, still sediment-containing cultures (150 days of
151 incubation) were sequenced to track the development of microbial compositions over time
152 (for primers, see Table S1). The original sediment contained large number of ANME-1 and
153 the putative partner bacteria *Ca. Desulfofervidus*. The AOM culture got further enriched in
154 ANME-1 archaea and *Ca. Desulfofervidus*, whereas in the Ethane50 culture the GoM-Arc1
155 clade increased from < 0.1% in the original sediment to roughly 35% of all archaea (Fig. 2A).
156 Notably, the relative abundance of *Ca. Desulfofervidus* increased also in the Ethane50
157 culture. This indicates that *Ca. Desulfofervidus* was also involved as partner bacterium in the
158 thermophilic ethane culture.

159 To visualize the cells involved in the anaerobic oxidation of ethane, oligonucleotide
160 probes specific for the GoM-Arc1 clade and *Ca. Desulfofervidus* were applied on the
161 Ethane50 culture using catalyzed reporter deposition fluorescence in situ hybridization
162 (CARD-FISH; for probes see Table S1). The Ethane50 culture contained large and tightly
163 packed consortia with sizes of up to 40 μm in diameter formed by GoM-Arc1 and *Ca.*
164 *Desulfofervidus* cells (Fig 2D,E). In the consortia, archaea and bacteria grew spatially
165 separated. These large consortia apparently develop from small but already dense consortia
166 found in the inoculate, similar to what was found for cold-adapted AOM consortia [19]. Such
167 a separation of the partner organisms is also characteristic for consortia in the butane-
168 degrading culture [9] and for most AOM consortia [20]. In contrast, in thermophilic AOM
169 consortia of ANME-1 and *Ca. Desulfofervidus*, the partner cells appear well-mixed [21]. The
170 Ethane50 culture differs from the cold-adapted ethane oxidizing culture in which *Ca.*
171 *Argoarchaeum* forms rather loose assemblages with yet uncharacterized bacteria [5].

172 To analyze the metabolic potential of the microorganisms involved in ethane
173 degradation, Ethane37 and Ethane50 cultures were subjected to transcriptomic and genomic
174 analysis. The 16S rRNA sequences extracted from the shot-gun RNA reads of the Ethane50
175 culture were strongly dominated by GoM-Arc1 (60%) and *Ca. Desulfofervidus* (20%; Fig.
176 2C), supporting a crucial role of these two organisms in thermophilic ethane degradation.
177 Long-read DNA sequencing for the Ethane50 culture resulted in a partial genome of GoM-
178 Arc1 with 76.2% completeness (GoM-Arc1_E50_DN), whereas by applying this approach to
179 the Ethane37 culture we obtained a closed genome of the GoM-Arc1 archaeon (GoM-
180 Arc1_E37). Both GoM-Arc1 genomes share an average nucleotide identity (ANI) of 98%,
181 hence a complete consensus genome for Ethane50 (GoM-Arc1_E50) was obtained by
182 mapping long reads of the Ethane50 culture on the closed GoM-Arc1_E37 genome (see
183 Material and Methods and Table S2). GoM-Arc1_E50 had a size of 1.92 MB and a GC
184 content of 46.5%. To assess the genomic diversity of archaea of the GoM-Arc1 clade,
185 additionally a MAG of GoM-Arc1 from the Loki's Castle hydrothermal vent field (GoM-
186 Arc1-LC) with a completeness of 68% and eight single-cell amplified genomes (SAGs) from
187 different cold seeps and different completenesses (10 to 59%) were retrieved (Table S2). The
188 MAG GoM-Arc1-LC and the eight single cells have an average nucleotide identity (ANI) of
189 over 90% suggesting that they belong to the same species. The 16S rRNA gene identity is in
190 the range of 99.5% supporting a definition as same species and shows that the same species of
191 GoM-Arc1 can be found in diverse seep sites (Table S2 and Figure S1). Together with several
192 MAGs of the GoM-Arc1 clade archaea from public databases [5, 17, 18] these MAGs now
193 provide an extensive database for the genomic description of the GoM-Arc1 clade. All GoM-
194 Arc1 clade genomes have an estimated size smaller than 2 Mb, which is in the range of the
195 other thermophilic alkane degraders, such as *Ca. Syntrophoarchaeum* (1.5-1.7 Mb) and
196 ANME-1 (1.4-1.8 Mb) [9, 22]. The genome is however much smaller than the 3.5 Mb

197 genome of the mesophilic sister lineage *Ca. Methanoperedens*. This organism is thriving on
198 methane and is capable to reduce nitrate or metals without partner bacteria [23, 24].

199 All GoM-Arc1 genomes contain the genes encoding the enzymes of the
200 methanogenesis pathway, including a highly similar divergent-type MCR and the Wood-
201 Ljungdahl pathway, but no pathway for beta-oxidation of longer fatty acids. Hence it is likely
202 that all members of this clade are ethane oxidizers. Based on 16S rRNA gene phylogeny and a
203 genome tree based on 32 marker genes, the GoM-Arc1 clade divides into two sub-clusters.
204 According to a 16S rRNA gene identity of ~95% (Fig. S1) and an average amino acid identity
205 (AAI) of ~63% (Fig. 3A; Table S2), these clusters should represent two different genera. One
206 cluster contains the recently described ethane oxidizer *Candidatus* Argoarchaeum
207 ethanivorans and genomes derived from cold environments including the Gulf of Mexico and
208 the moderately heated Loki's Castle seeps [25]. The second cluster includes the thermophilic
209 GoM-Arc1 strains found in the Ethane50 and the Ethane37 culture and sequences of other
210 MAGs from the Guaymas Basin [16, 18]. Based on the substrate specificity (see results
211 below) and its optimal growth at elevated temperatures, we propose to name the Ethane50
212 strain of GoM-Arc1 *Candidatus* Ethanoperedens thermophilum (Ethanoperedens, Latin for
213 nourishing on ethane; thermophilum, Latin for heat loving).

214 **Genomic and catabolic features of *Ca. Ethanoperedens*.** The main catabolic pathways of
215 *Ca. Ethanoperedens* are a complete methanogenesis and a Wood-Ljungdahl pathway (Fig. 4).
216 Its genome encodes for only one MCR. The three MCR subunits $\alpha\beta\gamma$ are on a single operon.
217 The amino acid sequence of the single alpha subunit (*mcrA*) of *Ca. Ethanoperedens* is
218 phylogenetically most closely related with the recently described divergent-type MCR of *Ca.*
219 *Argoarchaeum* with an amino acid identity of 69%, but also with all other *mcrA* sequences of
220 GoM-Arc1 archaea [5, 12, 16, 18]. These MCRs form a distinct cluster in comparison to other
221 divergent MCRs and to the canonical MCRs of methanogens and methanotrophs (Fig. 3B).

222 The similarity of GoM-Arc1 *mcrA* sequences to the described canonical and non-canonical
223 sequences is below 43% and changes in the amino acid sequences are also found in the highly
224 conserved active site of the enzyme (Fig. S2). The relative expression of the *mcr* subunits
225 compared to all reads mapping to *Ca. Ethanoperedens* (reads per kilo base per million mapped
226 reads; RPKM; i.e. *mcrA* = 9790) is at least two times higher than the expression of all other
227 genes of the main catabolic pathway (Fig. 4; Table S3). The relative *mcr* expression of *Ca.*
228 *Ethanoperedens* is higher than the expression of the multiple *mcr* genes in *Ca.*
229 *Syntrophoarchaeum*, but lower than the expression of *mcr* in thermophilic ANME-1 archaea
230 [9, 22]. The relatively low expression of *mcr* in short-chain alkane oxidizing archaea can be
231 explained by the properties of their substrates. Short-chain alkane oxidation releases larger
232 amounts of energy than methane oxidation. Furthermore, the cleavage of C-H bonds in multi-
233 carbon compounds requires less energy than the cleavage of C-H bonds of methane [26],
234 hence less MCR might be required to supply the organism with sufficient energy.

235 To test the substrates activated by the MCR of *Ca. Ethanoperedens*, we supplied
236 different alkanes to the active Ethane50 culture replicates and analysed the extracted
237 metabolites. Cultures supplied with ethane show the *m/z* 168.9988 of the authentic ethyl-CoM
238 standard (Fig. 5A,B), which was not observed in the control incubation without substrate.
239 Moreover, addition of 30% 1-¹³C-ethane resulted in the increase of masses expected for 1-
240 ¹³C-ethyl-CoM and 2-¹³C-ethyl-CoM (Fig. 5C). This confirms that *Ca. Ethanoperedens*
241 produces ethyl-CoM from ethane. To test substrate specificity of *Ca. Ethanoperedens*, we
242 provided culture replicates with four different gaseous alkanes (methane, ethane, propane and
243 *n*-butane, and a mix of all four substrates). Besides the ethane-amended culture, sulfide was
244 only produced in the Ethane50 culture supplied with the substrate mix (Fig. S3). In agreement
245 to this, no other alkyl-CoM variant apart from ethyl-CoM was detected (Fig. 5A). This shows
246 that the MCR of *Ca. Ethanoperedens* and most likely all MCR enzymes of GoM-Arc1 archaea

247 (Fig. 3B) activate ethane, but no or only trace amounts of methane and other alkanes. The
248 high substrate specificity of the MCR is crucial for GoM-Arc1 archaea, as they miss the
249 enzymatic machinery to use larger CoM-bound intermediates, since they lack the fatty acid
250 degradation pathway that is required to degrade butane and propane [9]. *Ca. Ethanoperedens*
251 contains and expresses a complete methyl transferase (*mtr*). This enzyme might cleave small
252 amounts of methyl-CoM that might be formed as side reaction of the MCR. The methyl unit
253 would be directly transferred to the methylene-H4MPT reductase (*mer*) and oxidized in the
254 upstream part of the methanogenesis pathway to CO₂ (Fig. 4).

255 Based on the observed net reaction and the genomic information, *Ca. Ethanoperedens*
256 completely oxidizes ethane to CO₂. In this pathway, coenzyme A bound acetyl units are
257 oxidized in the Wood-Ljungdahl pathway and the upstream part of the methanogenesis
258 pathway (Fig. 4). Our model, however, does not explain how CoM-bound ethyl groups are
259 oxidized to acetyl units and ligated to CoA. Similar transformations are required in the other
260 multi-carbon alkane-oxidizing archaea such as *Ca. Syntrophoarchaeum* and *Ca.*
261 *Argoarchaeum* [5, 9]. Those oxidation reactions lack biochemical analogues, hence genomic
262 information alone allows only indirect hints on their function. In *Ca. Ethanoperedens*, a
263 release of ethyl-units and transformation as free molecules (ethanol to acetate) is unlikely,
264 because a subsequently required formation of acetyl-CoA from acetate would require CoA
265 ligases, which are not present in the genome. Instead, the transformation of ethyl into acetyl
266 units could be performed by a tungstate-containing aldehyde ferredoxin oxidoreductase
267 (AOR) that could catalyse the oxidation with cofactors such as CoM or CoA. In the archaeon
268 *Pyrococcus furiosus* AORs transform aldehydes to the corresponding carboxylic acid [27].
269 Both, *Ca. Ethanoperedens* and *Ca. Argoarchaeum* genomes, contain three *aor* copies, and in
270 all cases these genes are either located in close proximity or on operons with genes of the
271 methanogenesis pathway. We detected a high expression of two of the three *aor* genes

272 (RPKM *aor* = 3805 and 7928), indicating a viable function of the enzymes. Likewise, very
273 high protein concentrations of these enzymes were shown for *Ca. Argoarchaeum* [5],
274 supporting the hypothesis of a critical function. An *aor* gene is also present in the butane
275 oxidizer *Ca. Syntrophoarchaeum*, yet its expression is rather moderate [9], which questions its
276 role in the catabolic pathway of this organism. In contrast ANME archaea do not contain or
277 overexpress *aor* genes, likely because the encoded enzymes have no central role in their
278 metabolism. We searched the cell extracts for potential intermediates in the pathway, but
279 based on retention time and mass we were not able to detect potential intermediates such as
280 ethyl-CoA. Similarly, acetyl-CoA, the substrate of the Wood-Ljungdahl pathway, was not
281 detected. A lack of detection however does not exclude those compounds as intermediate.
282 Instead, the compound turnover might be very fast, which could be required for an efficient
283 net reaction. Additionally, a mass spectrometric detection of unknown intermediates could be
284 hindered by compound instability or loss during the extraction. Further metabolite studies and
285 enzyme characterizations are required to understand the role of AOR in alkane oxidation

286 Acetyl-CoA, the product formed of the above proposed reactions, can be introduced
287 into the Wood-Ljungdahl pathway. The acetyl group is decarboxylated by the highly
288 expressed acetyl-CoA decarbonylase/ synthase (ACDS), and the remaining methyl group is
289 transferred to tetrahydromethanopterin (H₄-MPT). The formed methyl-H₄-MPT can then be
290 further oxidized to CO₂ following the reverse methanogenesis pathway (Fig. 4). *Ca.*
291 *Ethanoperedens* lacks genes for sulfate or nitrate reduction, similarly to other genomes of the
292 GoM-Arc1 clade. The electrons produced in the oxidation of ethane thus need to be
293 transferred to the sulfate-reducing partner bacterium *Ca. D. auxilii*, as previously shown for
294 the anaerobic oxidation of methane and butane. In co-cultures of *Ca. Argoarchaeum* and their
295 partner bacteria, Chen and co-workers (2019) suggest the transfer of reducing equivalents via
296 zero-valent sulfur between the loosely aggregated *Ca. Argoarchaeum* and its partner

297 bacterium, analogous to the hypothesis of Milucka, Ferdelman [28]. In the Ethane50 culture
298 such a mode of interaction is highly unlikely, as the partner *Ca. D. auxilii* is an obligate
299 sulfate reducer, incapable of sulfur disproportionation [11]. Based on genomic information,
300 direct electron transfer appears to be more likely. Alkane-oxidizing archaea and their partner
301 bacterium *Ca. D. auxilii* produce cytochromes and pili-based nanowires when supplied with
302 their substrate [9, 29, 30]. Also *Ca. Ethanoperedens* contains 11 different genes for
303 cytochromes with expression values of up to 14800 RPKM representing some of the highest
304 expressed genes in the culture (Table S3). Interestingly, *Ca. Ethanoperedens* also contains and
305 expresses a type IV pilin protein with a high RPKM value of 11246. The partner bacterium
306 *Ca. Desulfofervidus* also shows a high expression of pili and cytochromes under ethane
307 supply, showing their potential importance for the interaction of these two organisms in the
308 syntrophic coupling of ethane oxidation to sulfate reduction.

309 **Environmental distribution of GoM-Arc1 archaea.** 16S rRNA gene sequences clustering
310 with *Ca. Ethanoperedens* and *Ca. Argoarchaeum* have been found in hydrocarbon-rich marine
311 environments like cold seep and hot vent environments, including asphalt seeps in the Gulf of
312 Mexico and the Guaymas Basin hydrothermal vents in the Gulf of California [31-33]. In some
313 environments like oil seeps of the Gulf of Mexico and gas-rich barite chimneys of the Loki's
314 Castle, 16S rRNA gene surveys have shown that up to 30% of archaeal gene sequences
315 belonged to the GoM-Arc1 clade [12]. To estimate absolute abundances and potential
316 partnerships of GoM-Arc1 in the environment, we performed CARD-FISH on samples from
317 different seep and vent sites across the globe (Fig. 6). With up to 10^8 cells per ml, archaea of
318 the GoM-Arc1 clade were particularly abundant in cold seep sediments in the northern Gulf
319 of Mexico (St. 156). This cold seep transports thermogenic hydrocarbon gases that are
320 particularly enriched in short-chain alkanes [34, 35]. Other cold seep and hot vent sediments
321 from the Guaymas Basin, Hydrate Ridge and Amon Mud Volcano contain between 10^5 and

322 10^6 GoM-Arc1 cells per ml of sediment, which represents 1-5% of the archaeal community
323 (Fig. 6A). At all sites we found that GoM-Arc1 associates with partner bacteria. At the
324 hydrothermally-heated site in the Guaymas Basin, GoM-Arc1 aggregated with *Ca.*
325 *Desulfofervidus*, the partner bacterium of the Ethane37 and Ethane50 cultures. At Loki's
326 Castle, GoM-Arc1 and *Ca. Desulfofervidus* were co-occurring in barite chimneys based on
327 sequence information, yet they were not found to form the same tight consortia as at other
328 sites. At the temperate site Katakolo Bay in Greece, GoM-Arc1 archaea formed consortia with
329 very large, yet unidentified vibrioform bacteria (Fig. 6B-F). These cells hybridized with a
330 probe for *Deltaproteobacteria* but not with probes for known partner bacteria (for probes see
331 Table S1). At the cold seep sites, the associated cells could not be stained with probes for the
332 known partner bacteria of cold-adapted ANME including SEEP-SRB1, and SEEP-SRB2, and
333 also not with that for *Ca. Desulfofervidus*. It remains an important question as to how the
334 archaea can select only few specific types of bacteria as partner in the anaerobic alkane
335 oxidation, and for which specific traits they are selected. Based on their global presence in
336 hydrocarbon-rich environments, GoM-Arc1 archaea could be considered as key player in the
337 anaerobic oxidation of ethane in marine sediments. Its role would be similar to the role of
338 ANME archaea in AOM.

339 **Future possible applications of *Ca. Ethanoperedens*.** Archaea of the GoM-Arc1 cluster are
340 likely the dominant, if not the only organisms capable of the anaerobic oxidation of ethane in
341 the global seafloor. An important further task is to assess deep oil and gas reservoirs for their
342 diversity of ethane oxidizers. The rapid growth of *Ca. Ethanoperedens* and the streamlined
343 genome make it a model organism for the study of anaerobic ethanotrophy in archaea. The
344 biochemistry of short-chain alkane-oxidizing archaea will be of high interest for future
345 biotechnological applications. An organism using the metabolism of *Ca. Ethanoperedens* in
346 the reverse direction should be able to produce ethane, similar to methane production by

347 methanogens. Yet, there is scarce isotopic evidence for the existence of ethanogenic organism
348 in nature [36]. Furthermore, under common environmental conditions thermodynamics favor
349 the production of methane from inorganic carbon over the production of ethane. To test the
350 general reversibility of the ethane oxidation pathway, we incubated the active Ethane50
351 culture with ^{13}C -labelled inorganic carbon and traced the label transfer into ethane. Within 18
352 days $\delta^{13}\text{C}$ -ethane values increased from -3‰ to $+120\text{‰}$, whereas isotopic compositions in the
353 non-labeled culture remained stable (Fig. S4). Considering the forward rate and ethane stock,
354 the back reaction amounts to 1.5‰ to 3% of the forward reaction, which is in the range for
355 back fluxes of carbon measured in AOM [21, 37]. This experiment shows that the ethane
356 oxidation pathway is fully reversible. To test the net ethane formation in the Ethane50 culture,
357 we removed sulfate from culture aliquots and added hydrogen as electron donor. These
358 cultures formed between 1 and 17 $\mu\text{mol l}^{-1}$ ethane within 27 days (Table S6). The ethane
359 production was however a very small fraction of 0.08% of the ethane oxidation rate in
360 replicate incubations with ethane and sulfate. No ethane was formed in the presence of
361 hydrogen and sulfate. We interpret the ethane formation in the culture as enzymatic effect in
362 the ethane-oxidizing consortia. Hydrogenases will fuel reducing equivalents into the pathway,
363 which may ultimately lead to the reduction of carbon dioxide to ethane. A growing culture
364 could not be established under these conditions, however, the experiments suggest that related
365 or genetically modified methanogenic archaea could thrive as ethanogens. A complete
366 understanding of the pathway and enzymes of GoM-Arc1 archaea, however, is required to
367 develop the biotechnological potential of an ethanogenic organism. To allow energy-
368 conserving electron flows in this organism, a genetically modified methanogen should be used
369 as host organism. For a targeted modification of such archaea, the pathway of ethane
370 oxidation must be completely understood, and research should focus especially on the
371 transformation of coenzyme M bound ethyl units to coenzyme A bound acetyl units.

372

373 MATERIAL AND METHODS

374 **Inoculum and establishment of alkane-oxidizing cultures.** This study bases on samples
375 collected during R/V ATLANTIS cruise AT37-06 with submersible *Alvin* to the Guaymas
376 Basin vent area in December 2016 (for locations see Table S4). A sediment sample was
377 collected by push coring within a hydrothermal area marked by conspicuous orange-type
378 *Beggiatoa* mats (Dive 4869, Core 26, 27°0.4505'N 111°24.5389'W, 2001 m water depth,
379 December 20, 2016). The sampling site was located in the hydrothermal area where, during a
380 previous *Alvin* visit, sediment cores containing locally ¹³C-enriched ethane had indicated
381 ethane-oxidizing microbial activity [33]. In situ temperature measurements using the *Alvin*
382 heat flow probe revealed a steep temperature gradient reaching 80°C at 30-40 cm sediment
383 depth. The retrieved samples contained large amounts of natural gas as observed by bubble
384 formation. Soon after recovery, the overlying *Beggiatoa* mat was removed, and the top 10 cm
385 of the sediment were filled into 250 ml Duran bottles, which were gas-tight sealed with butyl
386 rubber stoppers. In the home laboratory, sediments were transferred into an anoxic chamber.
387 There a sediment slurry (20% sediment and 80% medium) was produced with synthetic
388 sulfate reducer (SR) medium (pH 7.0) [38, 39], and distributed into replicate bottles (sediment
389 dry weight per bottle 1.45 g). These bottles were amended with methane or ethane (0.2 MPa),
390 or kept with a N₂ atmosphere without alkane substrate. These samples were incubated at
391 37°C, 50°C and 70°C. To determine substrate-dependent sulfide production rates, sulfide
392 concentrations were measured every 2-4 weeks using a copper sulfate assay [40]. Ethane-
393 dependent sulfide production was observed at 37°C and 50°C, but not at 70°C. When the
394 sulfide concentration exceeded 15 mM the cultures were diluted (1:3) in SR medium and
395 resupplied with ethane. Repeated dilutions led to virtually sediment-free, highly active

396 cultures within 18 months. A slight decrease of the initial pH value to 6.5 led to increased
397 ethane oxidation activity and faster growth in the culture.

398 **Quantitative substrate turnover experiment.** The Ethane50 culture was equally distributed
399 in 6×150 ml serum flasks using 20 ml inoculum and 80 ml medium. Three replicate cultures
400 were amended with 0.05 MPa ethane in 0.1 MPa $N_2:CO_2$, while 3 negative controls were
401 amended with 0.15 MPa $N_2:CO_2$. Both treatments were incubated at 50°C. Weekly, 0.5 ml
402 headspace gas samples were analyzed for ethane content using an Agilent 6890 gas
403 chromatograph in split-less mode equipped with a packed column (Supelco Porapak Q,
404 $6ft \times 1/8' \times 2.1$ mm SS, oven temp 80°C). The carrier gas was helium (20 ml per minute) and
405 hydrocarbons were detected by flame ionization detection. Each sample was analyzed in
406 triplicates and quantified against ethane standards of 5, 10 and 100%. Derived concentrations
407 were converted into molar amounts by taking the headspace size, pressure and temperature
408 into account. Results were corrected for sampled volumes. Sulfide concentrations were
409 measured as described above. To determine sulfate concentrations 1 ml of sample was fixed
410 in 0.5 ml zinc acetate. Samples were diluted 1:50 with deionized water (MilliQ grade; >18.5
411 M Ω) and samples were measured using non-suppressed ion chromatography (Metrohm 930
412 Compact IC Metrosep A PCC HC/4.0 preconcentration and Metrosep A Suup 5 -150/ 4.0
413 chromatography column).

414 **DNA extraction, 16S rRNA gene amplification and tag sequencing.** DNA was extracted
415 from the different cultures and the original sediment with the MoBio power soil DNA
416 extraction kit (MO BIO Laboratories Inc., Carlsbad, USA) using a modified protocol. 20 ml
417 of the culture was pelleted via centrifugation ($5.000 \times g$; 10 min). The pellet was resuspended
418 in saline phosphate buffer (PBS) and transferred to the *PowerBeat Tube* (manufacturer
419 information needed). The cells were lysed by three cycles of freezing in liquid nitrogen (20
420 sec) and thawing (5 min at 60°C). After cooling down to room temperature, 10 μ l of

421 proteinase K (20 mg ml⁻¹) were added and incubated for 30 min at 55°C. Subsequently, 60 µl
422 of solution C1 (contains SDS) were added and the tubes were briefly centrifuged. The
423 samples were homogenized 2 times for 30 sec at 6.0 M/S using a FastPrep-24 instrument (MP
424 Biomedicals, Eschwege, Germany). In between the runs, the samples were kept on ice for 5
425 min. After these steps, the protocol was followed further according to the manufacturer's
426 recommendations. DNA concentrations were measured using a Qubit 2.0 instrument
427 (Invitrogen, Carlsbad, USA). 2 ng of DNA was used for amplicon PCR and the product used
428 for 16S rDNA amplicon library preparation following the 16S Metagenomic Sequencing
429 Library Preparation guide provided by Illumina. The Arch349F - Arch915R primer pair was
430 used to amplify the archaeal V3-V5 region and the Bact341F - Bact785R primer pair for the
431 bacterial V3-V4 region (Table S1). Amplicon libraries for both Archaea and Bacteria were
432 sequenced on an Illumina MiSeq instrument (2×300 bp paired end run, v3 chemistry) at
433 CeBiTec (Bielefeld, Germany). After analysis adapters and primer sequences were clipped
434 from the retrieved sequences using cutadapt [41](v1.16) with 0.16 (-e) as maximum allowed
435 error rate and no indels allowed. Resulting reads were analysed using the SILVAngs pipeline
436 using the default parameters (<https://ngs.arb-silva.de/silvangs/>) [42-44].

437 **Extraction of high quality DNA, library preparation and sequencing of genomic DNA.**

438 Biomass from 200 ml of the Ethane50 and Ethane37 cultures was pelleted by centrifugation
439 and resuspended in 450 µl of extraction buffer. Genomic DNA was retrieved based on a
440 modified version of the protocol described in [45], including three extraction steps.
441 Resuspended pellet was frozen in liquid N₂ and thawed in a water bath at 65°C Another 1350
442 µl of extraction buffer were added. Cells were digested enzymatically by proteinase K
443 (addition of 60 µl 20 mg/ml, incubation at 37°C for 1.5 h under constant shaking at 225 rpm),
444 and chemically lysed (addition of 300 µl 20% SDS for 2 h at 65°C). Samples were
445 centrifuged (20 min, 13.000 × g) and the clear supernatant transferred to a new tube. 2 ml of

446 chloroform: isoamylalcohol (16:1, v/v) were added to the extract, mixed by inverting and
447 centrifuged for 20 minutes at $13.000 \times g$. The aqueous phase was transferred to a new tube,
448 mixed with 0.6 volumes of isopropanol, and stored over night at $-20^{\circ} C$ for DNA
449 precipitation. The DNA was redissolved in water at $65^{\circ}C$ for 5 min and then centrifuged for
450 40 min at $13.000 \times g$. The supernatant was removed and the pellet washed with ice-cold
451 ethanol (80%) and centrifugation for 10 min at $13.000 \times g$. The ethanol was removed and the
452 dried pellet was resuspended in PCR-grade water. This procedure yielded 114 μg and 145 μg
453 high quality genomic DNA from the Ethane37 and the Ethane50 culture, respectively.
454 Samples were sequenced with Pacific Biosciences Sequel as long amplicon (4 – 10 kb) and
455 long read gDNA library at the Max Planck-Genome-Centre (Cologne, Germany). To evaluate
456 the microbial community we extracted 16S rRNA gene reads using Metaxa2 [46] and
457 taxonomically classified them using the SILVA ACT online service [47]. For assembly either
458 HGAP4 (Implemented in the SMRTlink software by PacBio) or Canu
459 (<https://github.com/marbl/canu>) were used. The closed GoM-Arc1 genome from the Ethane37
460 culture was prepared manually by the combination of assemblies from the two before
461 mentioned tools. The final genome was polished using the resequencing tool included in the
462 SMRT Link software by PacBio. For not circularized de-novo genomes, the resulting contigs
463 were mapped via minimap2 (<https://github.com/lh3/minimap2>; parameter: '-x asm10') to a
464 reference genome. The reference consensus genomes were prepared using the resequencing
465 tool implemented in the SMRTLink software of PacBio using either the circular GoM-Arc1
466 de-novo genome from this study or the publically available *Ca. Desulfofervidus* genome
467 (Accession: NZ_CP013015.1) as reference. Final genomes were automatically annotated
468 using Prokka [48] and the annotation refined manually using the NCBI Blast interface [49].
469 Average nucleotide and amino acid identities were calculated using Enveomics tools [50].

470 **Single cell genomics.** Anoxic sediment aliquots were shipped to the Bigelow Laboratory
471 Single Cell Genomics Center (SCGC; <https://scgc.bigelow.org>). Cells were separated, sorted
472 and lysed, and total DNA was amplified by multiple displacement amplification. Single cell
473 DNA was characterized by 16S rRNA gene tag sequences [12, 51]. The single cell amplified
474 DNA from Gulf of Mexico was analyzed sequenced as described before in [12]. Single cell
475 amplified DNA from Amon Mud Volcano AAA-792_C10 was sequenced with Hiseq3000
476 and MiSeq technology and reads were assembled using SPAdes [52] with the single cell
477 mode. Assembled reads were binned based on tetra-nucleotides, coverage and taxonomy
478 using MetaWatt [53]. The final SAG was evaluated for completeness and contamination using
479 CheckM [54]. Genome annotation was performed as described above.

480 **Extraction of RNA, reverse transcription, sequencing and read processing.** Extraction
481 and sequencing of total RNA was prepared in triplicates. RNA was extracted from 150 ml
482 active Ethane50 culture grown in separate bottles at 50°C. Total RNA was extracted and
483 purified as described in [9] using the Quick-RNA MiniPrep Kit (Zymoresearch, Irvine, CA,
484 USA) and RNeasy MinElute Cleanup Kit (QIAGEN, Hilden, Germany). Per sample at least
485 150 ng of high-quality RNA were obtained. RNA library was prepared with the TrueSeq
486 Stranded Total RNA Kit (Illumina). An rRNA depletion step was omitted. The samples were
487 sequenced on an Illumina Nextseq with v2 chemistry and 1×150bp read length. The
488 sequencing produced ~50 Gb reads per sample. Adaptors and contaminant sequences were
489 removed and reads were quality trimmed to Q10 using bbduk v36.49 from the BBMAP
490 package. For phylogenetic analysis of the active community 16S rRNA reads were recruited
491 and classified based on SSU SILVA Release 132 [47] using phyloFlash [55]. Trimmed reads
492 were mapped to the closed genomes of the *Candidatus* Ethanoperedens thermophilum and *Ca.*
493 *Desulfofervidus* using Geneious Prime 2019.2.1 (<https://www.geneious.com>) with a minimum
494 mapping quality of 30%. The expression level of each gene was quantified by counting the

495 number of unambiguously mapped reads per gene using Geneious. To consider gene length,
496 read counts were converted to reads per kilobase per million mapped reads (RPKM).

497 **Phylogenetic analysis of 16S rRNA genes, marker genes and *mcrA* amino acid**
498 **sequences.** A 16S rRNA gene based phylogenetic tree was calculated using publically
499 available 16S rRNA sequences from the SSU Ref NR 128 SILVA database [42]. The tree was
500 constructed using ARB [56] and the FastTree 2 package [57] using a 50% similarity filter.
501 Sequence length for all 16S rRNA genes was at least 1100 bp. After tree calculation, partial
502 sequences retrieved from single cells were included into the tree. ARB [56] was used for
503 visualization of the final tree. The marker gene tree was calculated using 126 publically
504 available genomes and genomes presented in this study. The tree was calculated based on
505 aligned amino acid sequences of 32 marker genes picked from known archaeal marker genes
506 (Table S5) [58]. For the preparation of the aligned marker gene amino acid sequences we used
507 the phylogenomic workflow of Anvio 5.5 [59]. The marker gene phylogeny was calculated
508 using RaxML version 8.2.10 [60] with the PROTGAMMAAUTO model and LG likelihood
509 amino acid substitution. 1000 fast bootstraps were calculated to find the optimal tree
510 according to RaxML convergence criteria. The software iTol v3 was used for tree
511 visualization [61]. The *mcrA* amino acid phylogenetic tree was calculated using 358
512 sequences that are publically available or presented in this study. The sequences were
513 manually aligned using the Geneious Prime 2019.2.1 (<https://www.geneious.com>) interface
514 and 1060 amino acid positions considered. The aligned sequences were masked using Zorro
515 (<https://sourceforge.net/projects/probmask/>) and a phylogenetic tree was calculated using
516 RaxML version 8.2.10 [60] using the PROTGAMMAAUTO model and LG likelihood amino
517 acid substitution. 1000 fast bootstraps were calculated. The tree was visualized with iTol v3
518 [61].

519 **Catalyzed reported deposition fluorescence in situ hybridization (CARD-FISH).** Aliquots
520 of the Ethane50 culture and environmental samples were fixed for 1 h in 2% formaldehyde,
521 washed three times in PBS (pH = 7.4): ethanol 1:1 and stored in this solution. Aliquots were
522 sonicated (30 sec; 20% power; 20% cycle; Sonoplus HD70; Bandelin) and filtered on GTTP
523 polycarbonate filters (0.2 μm pore size; Millipore, Darmstadt, Germany). CARD-FISH was
524 performed according to [62] including the following modifications: Cells were permeabilized
525 with a lysozyme solution (PBS; pH 7.4, 0.005 M EDTA pH 8.0, 0,02 M Tris-HCl pH 8.0, 10
526 mg ml^{-1} lysozyme; Sigma-Aldrich) at 37°C for 60 minutes followed by proteinase K solution
527 treatment (7.5 $\mu\text{g ml}^{-1}$ proteinase K; Merck, Darmstadt, Germany in PBS; pH=7.4, 0.005 M
528 EDTA pH 8.0, 0.02 M Tris-HCl pH 8.0,) at room temperature for 5 minutes. Endogenous
529 peroxidases were inactivated by incubation in a solution of 0.15% H_2O_2 in methanol for 30
530 min at room temperature. Horseradish-peroxidase (HRP)-labeled probes were purchased from
531 Biomers.net (Ulm, Germany). Tyramides were labeled with Alexa Fluor 594 or Alexa Fluor
532 488. All probes were applied as listed in Table S1. For double hybridization, the peroxidases
533 from the first hybridization were inactivated in 0.15% H_2O_2 in methanol for 30 min at room
534 temperature. Finally, the filters were counterstained with DAPI (4',6' -diamino-2-
535 phenylindole) and analyzed by epifluorescence microscopy (Axiophot II Imaging, Zeiss,
536 Germany). Selected filters were analyzed by Confocal Laser Scanning Microscopy (LSM 780,
537 Zeiss, Germany) including the Airyscan technology.

538 **Synthesis of authentic standards for metabolites.** To produce alkyl-CoM standards, 1 g of
539 coenzyme M was dissolved in 40 ml 30% (v/v) ammonium hydroxide solution and to this
540 solution 1.8 to 2 g of bromoethane, bromopropane or bromobutane was added. The mixture
541 was incubated for 5 h at room temperature under vigorous shaking and then acidified to pH 1
542 with HCl. The produced standard had a concentration of approx. 25 mg ml^{-1} which for mass
543 spectrometry measurements was diluted to 10 $\mu\text{g ml}^{-1}$.

544 **Extraction of metabolites from the Ethane50 culture.** In the anoxic chamber 20 ml of
545 Ethane50 culture was harvested into 50 ml centrifuge tubes. Tubes were centrifuged at 3000
546 rcf for 10 min and the supernatant was removed. The pellet was resuspended in 1 ml
547 acetonitrile:methanol:water (4:4:2; v/v/v) mixture in lysing matrix tubes (MP Biomedicals,
548 Eschwege, Germany) with glass beads. Afterwards the tubes were removed from the anoxic
549 chamber and the samples were mechanical lysed in a FastPrep homogenizer (MP Bio) with 5
550 cycles with 6 M/s for 50 sec, and cooling on ice for 5 min in between the homogenization
551 steps. Finally, the samples were centrifuged for 5 min at $13.000 \times g$ and the supernatant
552 transferred to a new tube and stored at -20°C .

553 **Solvents for LC-MS/MS.** All organic solvents were LC-MS grade, using acetonitrile (ACN;
554 BioSolve, Valkenswaard, The Netherlands), isopropanol (IPA; BioSolve, Valkenswaard, The
555 Netherlands), and formic acid (FA; BioSolve, Valkenswaard, The Netherlands). Water was
556 deionized by using the Astacus MembraPure system (MembraPure GmbH, Henningsdorf,
557 Berlin, Germany).

558 **High resolution LC-MS/MS.** The analysis was performed using a QExactive Plus Orbitrap
559 (Thermo Fisher Scientific) equipped with an HESI probe and a Vanquish Horizon UHPLC
560 System (Thermo Fisher Scientific). The metabolites from cell extracts were separated on an
561 Accucore C30 column (150×2.1 mm, $2.6 \mu\text{m}$, Thermo Fisher Scientific), at 40°C , using a
562 solvent gradient created from the mixture of the buffer A (5% Acetonitrile in water, 0.1%
563 formic acid) and buffer B (90/10 IPA/ACN, 0.1% formic acid). The solvent gradient was the
564 following: Fraction B (%) of 0, 0, 16, 45, 52, 58, 66, 70, 75, 97, 97,15, 0, at -2 min. (pre run
565 equilibration), 0, 2, 5.5, 9, 12, 14, 16, 18, 22, 25, 32.5, 33, 34.4 and 36 min. of each run, and
566 a constant flow rate of $350 \mu\text{l min}^{-1}$. The samples injection volume was $10 \mu\text{l}$. The MS
567 measurements were acquired in negative mode for a mass detection range of 70–1000 Da. In
568 alternation, a full MS and MS/MS scans of the eight most abundant precursor ions were

569 acquired in negative mode. Dynamic exclusion was enabled for 30 seconds. The settings for
570 full range MS1 were: mass resolution of 70,000 at 200 m/z , AGC target 5×10^5 , injection time
571 65 ms. Each MS1 was followed by MS2 scans with the settings: mass resolution 35,000 at
572 200 m/z , AGC target 1×10^6 , injection time 75 ms, loop count 8, isolation window 1 Da,
573 collision energy was set to 30 eV.

574 **Determination of carbon back flux into the ethane pool.** Aliquots of active AOM culture
575 (50 ml) were transferred into 70 ml serum bottles with $N_2:CO_2$ headspace. In the SIP
576 experiment addition of 99% ^{13}C -labeled inorganic carbon (1 ml, 350 mM) led to $\delta^{13}C$ -DIC
577 values of +25,000 ‰ as measured by cavity ringdown spectrometry. Ethane (2 atm = 1.8 mM)
578 was added to both experiments and cultures were stored at 50°C. To determine the overall
579 ethane oxidation activity sulfide concentrations were measured every few days as described
580 above and converted to ethane oxidation rates using ratios of eq. 1. To measure the
581 development of ethane $\delta^{13}C$ values 1 ml of the gas phase was samples every few days, as
582 stored it 10 ml Exetainer vials with 2 ml NaOH and measured ethane isotopic composition
583 using gas chromatography coupled via a combustion oven to isotope ratio mass spectrometry
584 (He as carrier gas, flow rate, column, temperature program).

585 **Net ethane production test.** To test for net ethane production, in 156 ml serum flasks
586 replicate incubations with about 0.5 g active Ethane50 culture (wet weight) in 100 sulfate-
587 free medium were prepared. Four different conditions were tested in three biological
588 replicates with the addition of (1) 1.5 atm H_2 ; (2) replicate condition to (1) but only 0.05 g
589 biomass (3) 1.5 atm H_2 plus 28 mM sulfate (4) an activity control with addition of sulfate
590 and 1.5 atm ethane. Cultures were incubated over 27 days at 50°C and sulfate and ethane
591 concentrations were monitored as described above.

592 **Data availability.** All sequence data are archived in the ENA database under the INSDC
593 accession numbers PRJEB36446 and PRJEB36096. Sequence data from Loki's Castle is
594 archived under NCBI BioSample number SAMN13220465. The 16S rRNA gene amplicon
595 reads have been submitted to the NCBI Sequence Read Archive (SRA) database under the
596 accession number SRR8089822. All sequence information has been submitted using the data
597 brokerage service of the German Federation for Biological Data (GFBio) [63], in compliance
598 with the Minimal Information about any (X) Sequence (MIxS) standard [64], but some data is
599 still under ENA embargo. For reviewers, sequence information is stored under
600 <https://owncloud.mpi-bremen.de/index.php/s/QSMycWOBB38AunL>.

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618 retrieved additional samples. C.H. and G.W. performed the cultivation, physiology and
619 isotope experiments. C.H., F.V. K.-M. V. and K.K. performed fluorescence microscopy.
620 C.H., M.L. and G.W. performed metabolite analysis. C.H., R. L.-P. M., R.A. K.K. and G.W.
621 performed metagenomic and phylogenetic analyses and developed the metabolic model, C.H.,
622 and G.W. wrote the manuscript with contributions from all co-authors.

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771

772

773 **Fig. 1** Cultivation and stoichiometry test of the Ethane 50 culture. **A**, Rates of methane-dependent
774 (blue) and ethane-dependent (red) sulfide production in sediments of the Guaymas Basin incubated at
775 50°C.. **B**. Determination of activity doubling times in anaerobic ethane-oxidizing culture. Logarithmic

776 y-axis with sulfide production shows a decrease in activity at 3 mM sulfide and estimated activity
777 doubling times in low sulfide concentrations of 6-7 days. **C**, Development of ethane (diamond), sulfate
778 (triangles) and sulfide (squares) concentrations in the Ethane50 culture. Gray symbols show
779 corresponding concentrations measured in control incubations without ethane addition (data from 1 of
780 3 replicate incubations, for complete data see Table S6. The ratios of the slopes of sulfate and sulfide
781 to ethane (1.92 and 1.82) are close to the stoichiometric ratios of sulfate reduction and ethane
782 oxidation. The small offset may relate to biomass production and sampling artifacts.

783

784 **Fig. 2** Microbial composition of the Ethane50 culture. **A**, Relative abundance of phylogenetic clades
785 of archaea and **B**, bacteria based on 16S rRNA gene amplicon sequencing present in the inoculated
786 sediment, and in cultures with no substrate, with methane and ethane after 150 days of incubation. **C**,
787 relative abundance of active microbial groups based on 16 rRNA fragments recruited from the genome
788 of Ethane37 and Ethane50 after 2.5 years of incubation and the transcriptome of the Ethane50 culture
789 after one year of incubation with ethane. **D**, Laser-scanning micrograph and **E**, epifluorescence
790 micrograph of microbial consortia stained with probes specific for the GoM-Arc1 clade (red, Alexa-
791 594) and *Ca. Desulfofervidus* (green, Alexa-488) in the Ethane50 culture. Scale 10 μm .

792

793 **Fig. 3** Phylogenetic affiliation based on 32 marker genes and *mcrA* amino acid sequences of *Ca.*
794 *Ethanoperedens*. **A**, Phylogenetic affiliation of *Ca. Ethanoperedens* within the *Euryarchaeota* based on
795 32 aligned marker gene amino acid sequences; outgroup is *Thaumarchaeota*. The scale bar indicates
796 10% sequence divergence. **B**, Phylogenetic affiliation of *mcrA* amino acid sequences. The *mcrA*
797 sequences of GoM-Arc1 form a distinct branch within the non-canonical, potentially multi-carbon
798 alkane-activating MCRs. The *mcrA* genes of the GoM-Arc1 cluster can be further divided into those
799 from cold-adapted organisms, including *Ca. Argoarchaeum ethanivorans*, and the cluster including the
800 thermophiles of the genus *Ca. Ethanoperedens*. Sequences from the Ethane50 enrichment are depicted
801 in red, environmental sequences from metagenomes and single cell genomes from this study in grey
802 and *Ca. Argoarchaeum ethanivorans* in blue.

803

804 **Fig. 4** Metabolic model of anaerobic ethane oxidation in *Ca. Ethanoperedens thermophilum*. Ethane is
805 activated in the ethane specific MCR. The produced CoM-bound ethyl groups are consecutively
806 oxidized and transformed to CoA-bound acetyl units. Acetyl-CoA is cleaved using the ACDS of the
807 Wood-Ljungdahl pathway. The remaining methyl groups are fully oxidized on the reversed
808 methanogenesis pathway. Similar to ANME archaea and *Ca. Syntrophoarchaeum*, *Ca. Ethanoperedens*
809 does not contain a reductive pathway; hence electrons released during ethane oxidation are transferred

810 to the partner bacterium *Ca. D. auxilii*. Therefore, in both partners, cytochromes and pili are present
811 and expressed, similar as described in thermophilic consortia performing AOM [22] (for detailed
812 expression patterns, see Table S3)

813

814 **Fig. 5** Detection of coenzyme M bound intermediates in the Ethane50 culture. **A.** Upper 4 panels,
815 Total ion counts for HPLC peaks for authentic standards of methyl-, ethyl-, propyl and butyl CoM,
816 and chromatograms for ethane and mixed alkane gases (methane to butane). **B,C.** Mass spectra (m/z
817 168.5-171.5) for culture extracts after providing the Ethane50 culture with **(B)** non-labeled ethane and
818 **(C)** 30% ^{13}C - labeled ethane. Diagrams show the relative intensities for ethyl-CoM-H ($^{12}\text{C}_4\text{H}_9\text{S}_2\text{O}_3^-$)
819 calculate m/z 168.9988) and its isotopomers with one or two ^{13}C carbon or one ^{34}S isotope.

820

821 **Fig. 6** Abundance and exemplary micrographs of GoM-Arc1 archaea in sediments from cold seeps and
822 Guaymas Basin. **A.** Abundance estimations of archaeal cells detected by the GoM-Arc1 specific probe
823 GOM-ARCI-660 in a CARD-FISH survey. Detection limit approx. 5×10^4 cells per ml sediment. **B – F**
824 Epifluorescence **(B-E)** and laser-scanning micrographs **(F)** of environmental samples using CARD-
825 FISH with combining the GoM-Arc1 specific probe (red) and the general bacterial probe EUB-338
826 (green). Environmental samples originated from the seep sites Hydrate Ridge, Oregon **(B)**; Gulf of
827 Mexico **(C)**, Guaymas Basin **(D)**; Loki's Castle **(E)** and Katakolo Bay, Greece **(F)**. Scale 5 μm for D-F
828 and 2 μm for B and C.

829

830 **Table S1** PCR primers used the amplification of archaeal and bacteria 16S rRNA genes and
831 oligonucleotide probes used for CARD-FISH.

832

833 **Table S2** Summary of single cell and metagenome assembled genomes presented in this study and
834 average nucleotide and amino acid identities. ANI and AAI values were calculated with publically
835 available genomes and genomes presented in this study. Enveomics tools were used for the calculation
836 [50].

837

838 **Table S3** Genomes and gene expression data of the Ethane50 culture and overview of genes
839 potentially involved in the ethane metabolism and electron cycling in the Ethane50 culture. Expression
840 values shown in triplicates for *Ca. E. thermophilum* and *Ca. D. auxilii*.

841

842 **Table S4** Overview of environmental sampling sites used for this study.

843

844 **Table S5** Marker genes used for calculation of genome tree based on archaeal marker genes presented
845 in Rinke, Schwientek [58].

846

847 **Table S6** Summary dataset of the development of substrates and products in the Ethane50 enrichment
848 culture. Development of ethane, sulfide and sulfate concentrations in E50 culture in triplicates.
849 Development of sulfide concentration in E50 culture in 10 replicates. Development of ethane and
850 sulfide concentrations in triplicates of the Ethane50 culture with hydrogen gas (1.5 atm) with and
851 without sulfate. The positive control contained 1.5 bar ethane and sulfate.

852

853 **Fig. S1** Phylogenetic affiliation of the GoM-Arc1 clade archaea with other archaea based on 16S
854 rRNA gene comparison. The tree was constructed using ARB [56] and the FastTree 2 package [57]
855 using a 50% similarity filter. 410 sequences with a length of at least 1100 bp, excluding partial
856 sequences retrieved from single cells, were used. Bar shows 10% sequence divergence.

857

858 **Fig. S2** Comparison of *mcrA* sequences from the GoM-Arc1 clade to described canonical and non-
859 canonical *mcrA* sequences. A, alignment of the active site of the *mcrA* from different representative
860 genomes. The four different *Ca. Syntrophoarchaeum* sequences belong to the same genome bin.

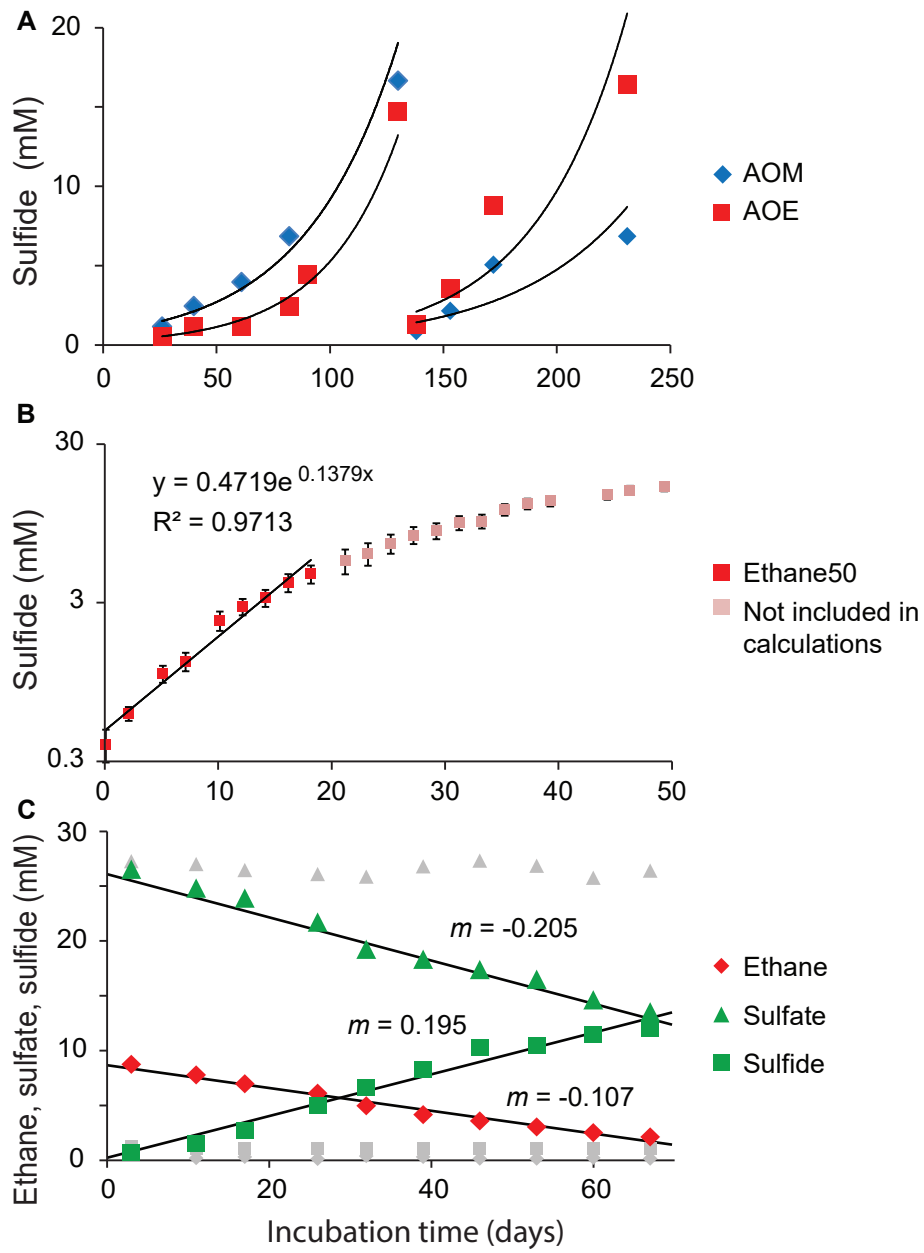
861 Amino acid positions refer to *Ca. E. thermophilum* E50 *mcrA* sequence. B, identity matrix of *mcrA*
862 sequences based on NCBI blastp alignment.

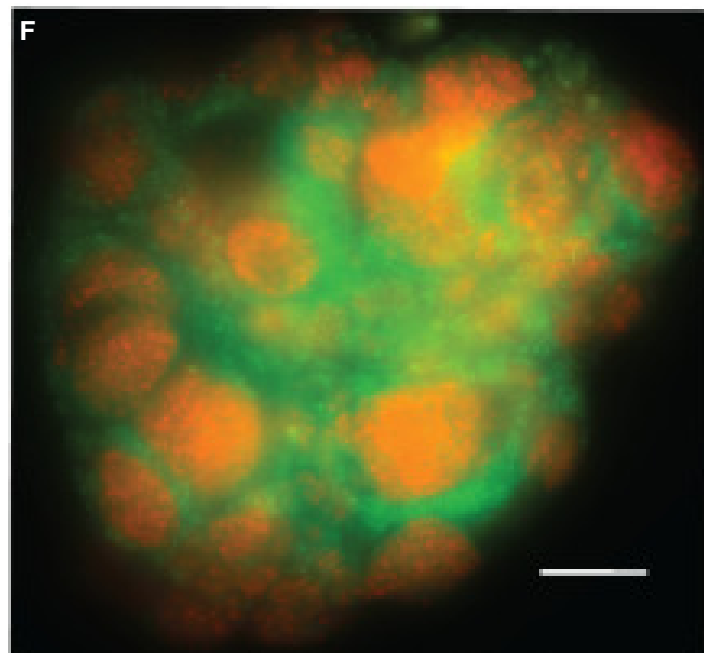
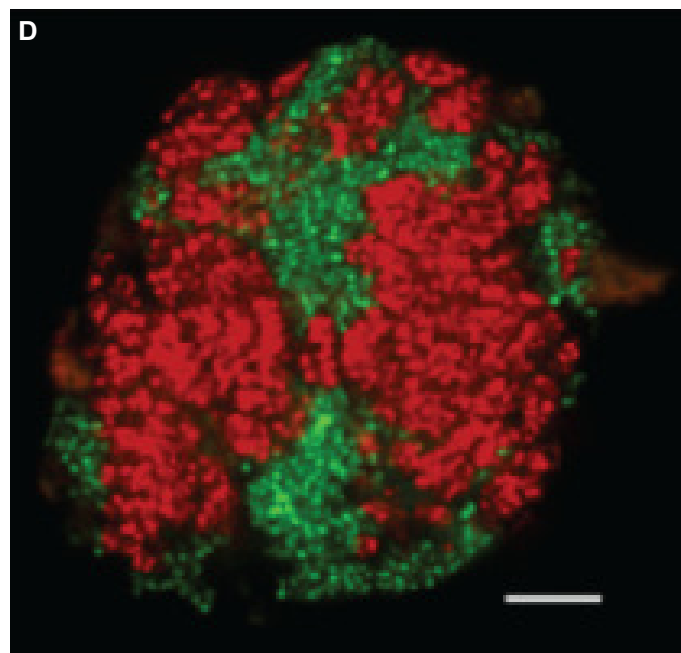
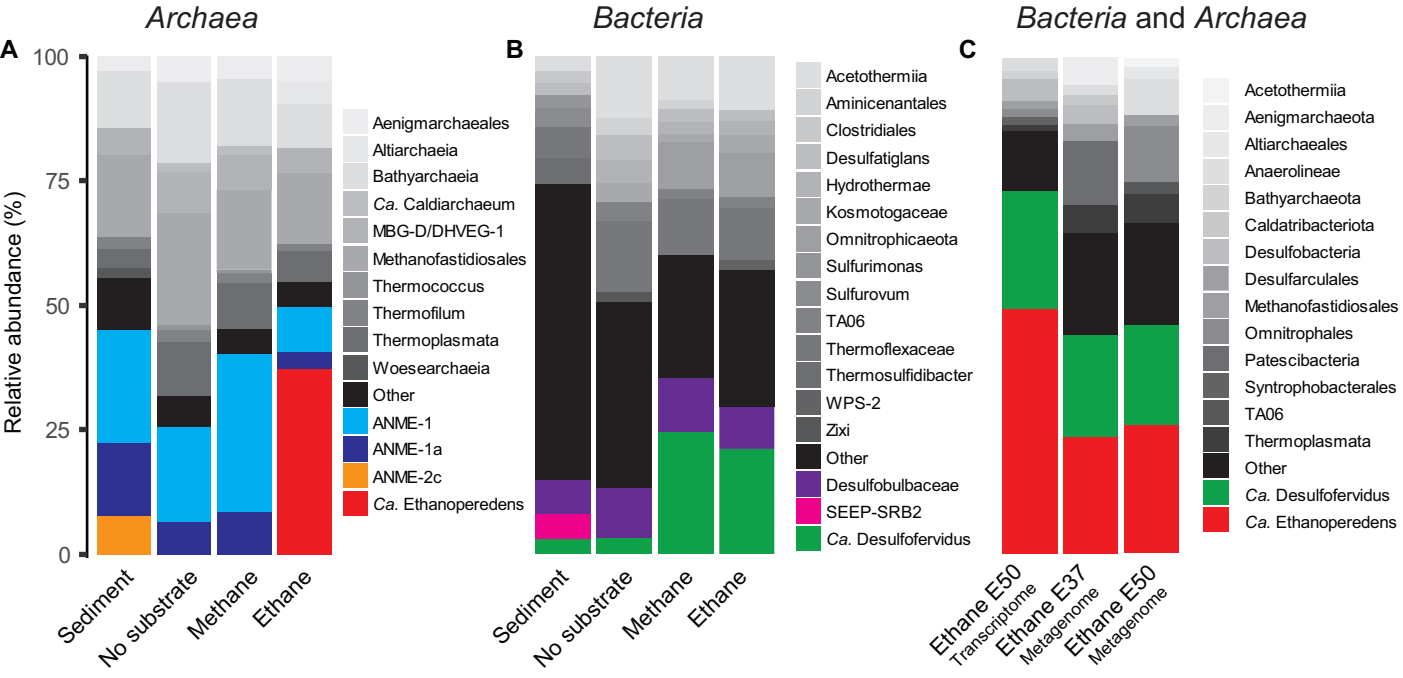
863

864 **Fig. S3** Development of sulfide concentrations in substrate experiments in replicate incubations of the
865 Ethane50 culture supplied with (A) ethane, alternative alkanes or a mix of these substrates; and (B)
866 ethane and sulfate compared to ethane with elemental sulfur or only elemental sulfur. Results show
867 that that ethane is the only alkane used as electron donor in *Ca. E. thermophilum*, and sulfate is the
868 only used electron acceptor. Further, elemental sulfur is not disproportionated.

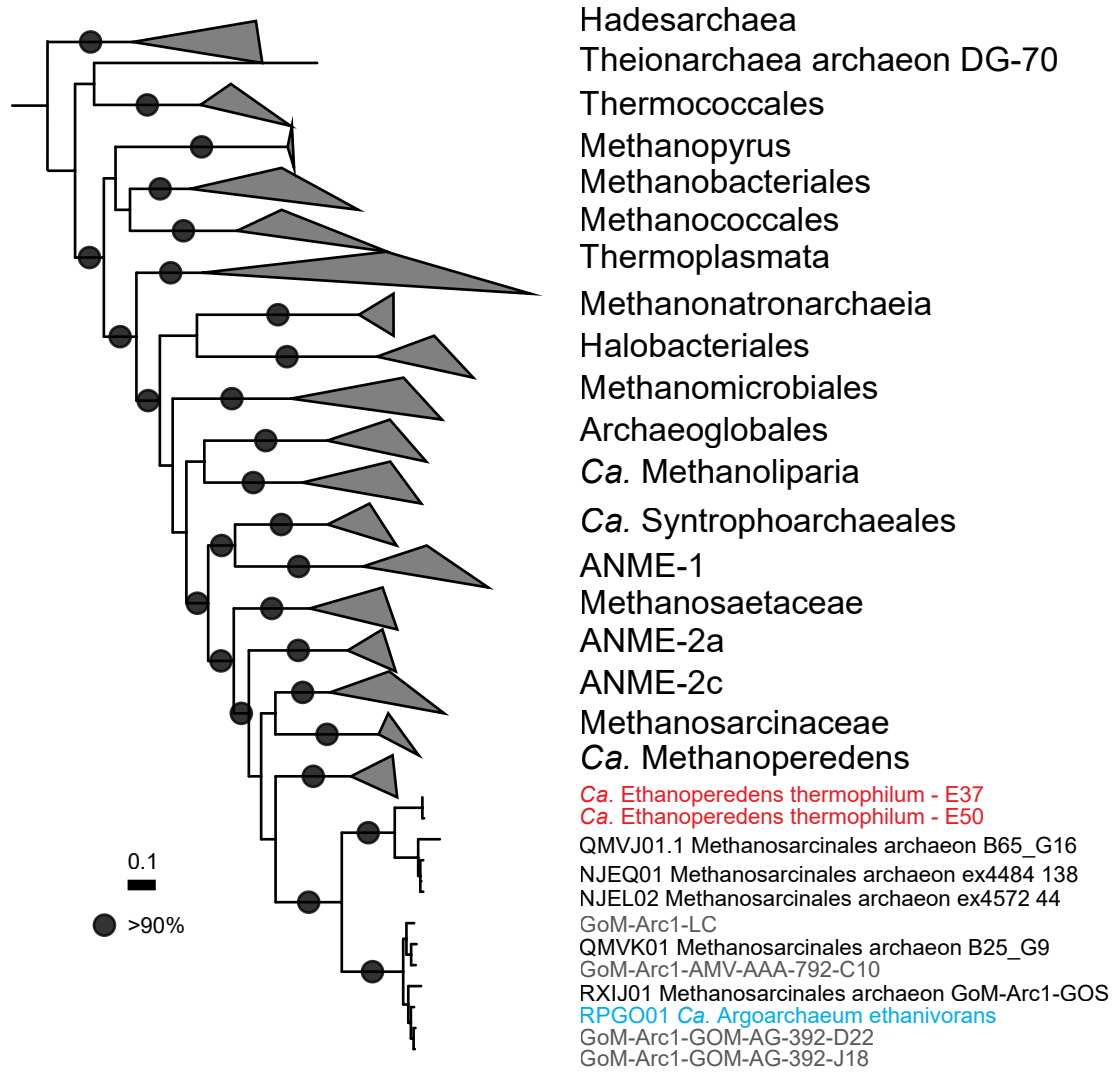
869

870 **Fig. S4** Test for the transfer of inorganic carbon into ethane in the ethane 50 culture. A, development
871 of sulfide concentrations in the culture with ethane as energy source and sulfate as electron acceptor B,
872 development of $\delta^{13}\text{C}$ values in ethane in the two cultures (controls $\delta^{13}\text{C}_{\text{DIC}} -35\%$) ^{13}C -DIC amended
873 culture with $\delta^{13}\text{C}_{\text{DIC}} =+25994 \%$. Based on simple mass balance calculations on the development of
874 fractions we infer that sulfate-dependent anaerobic AOM in these enrichments is accompanied by a
875 back flow of inorganic carbon amounting to 1-3% of the forward rate. This back reaction indicates a
876 general reversibility of ethane oxidation.

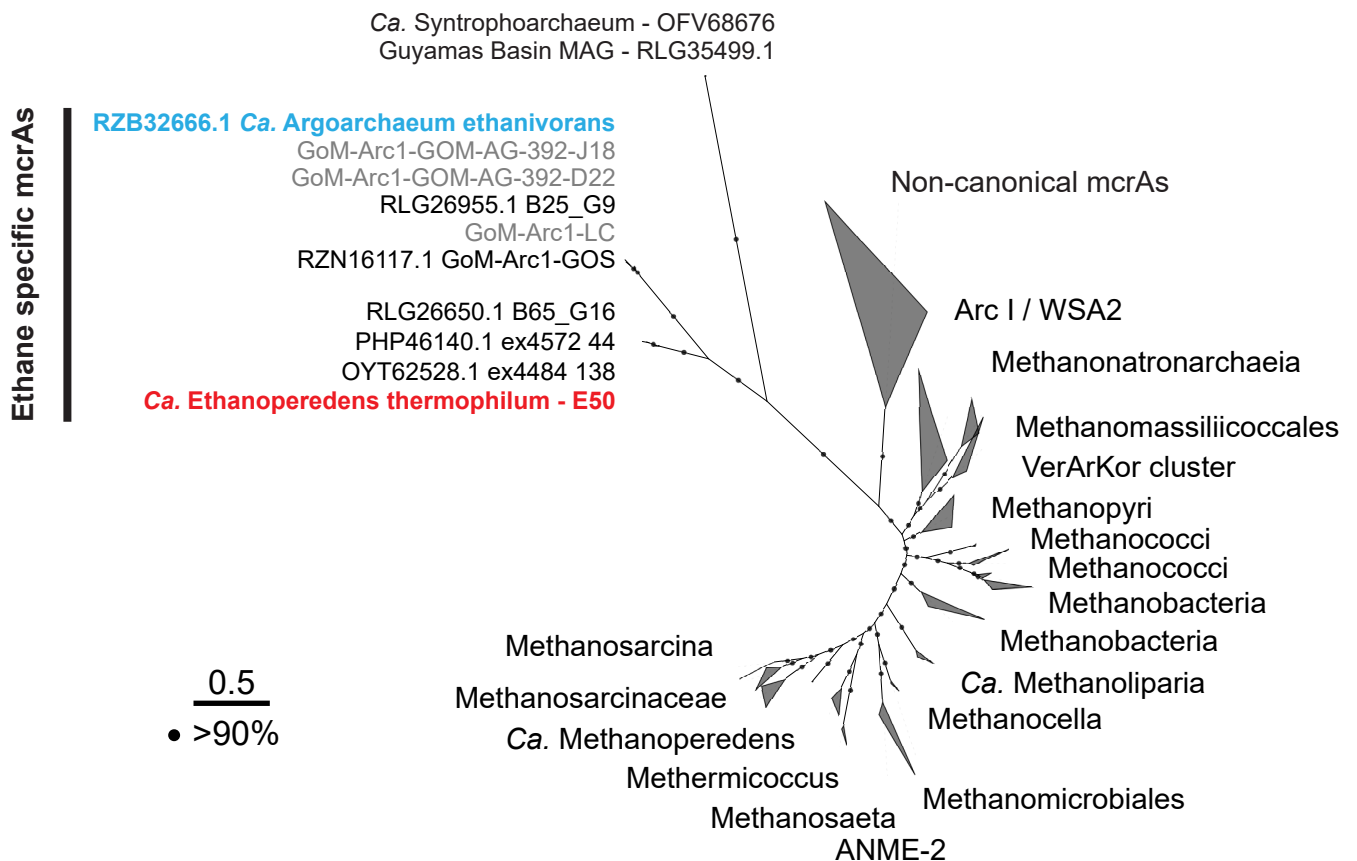


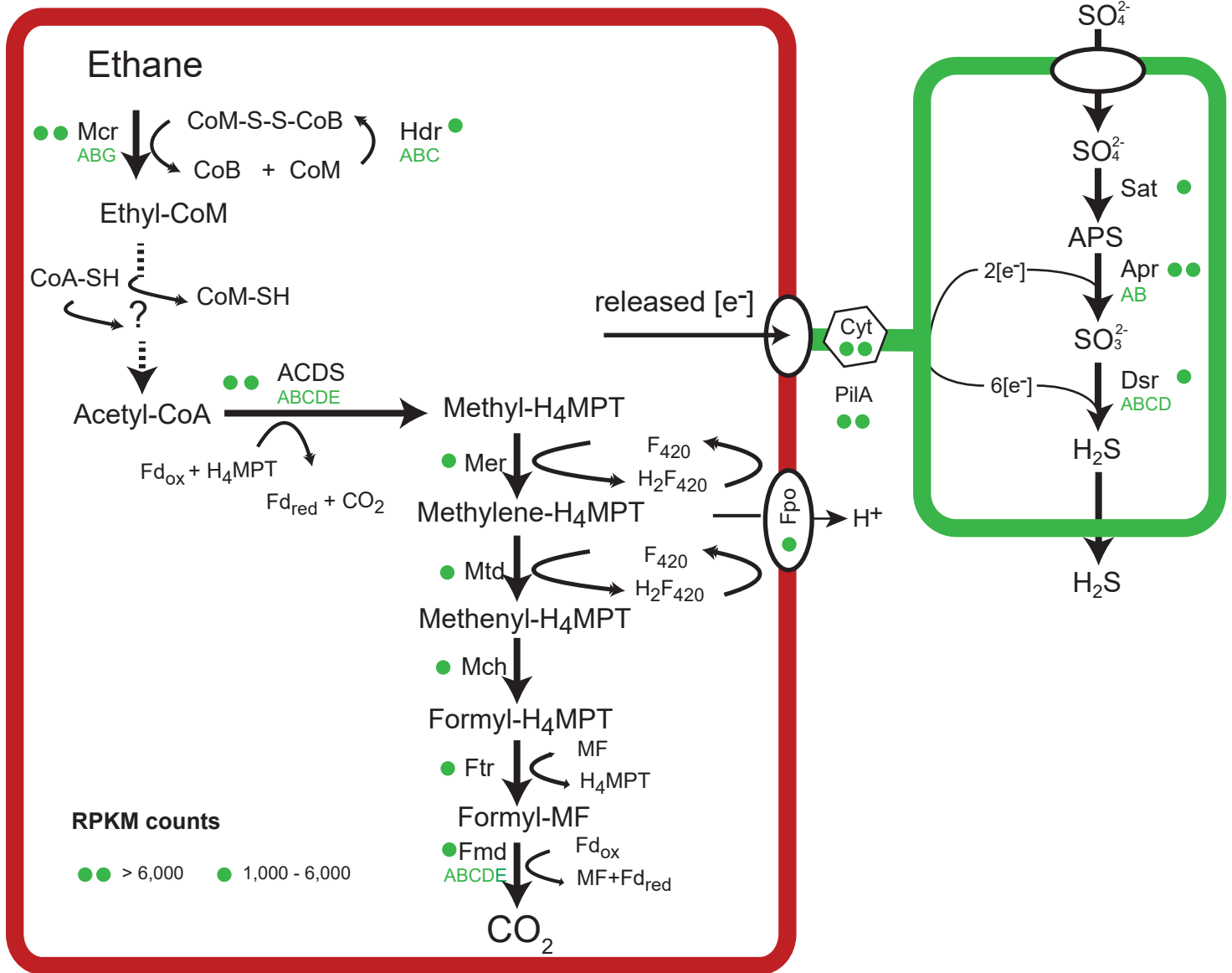


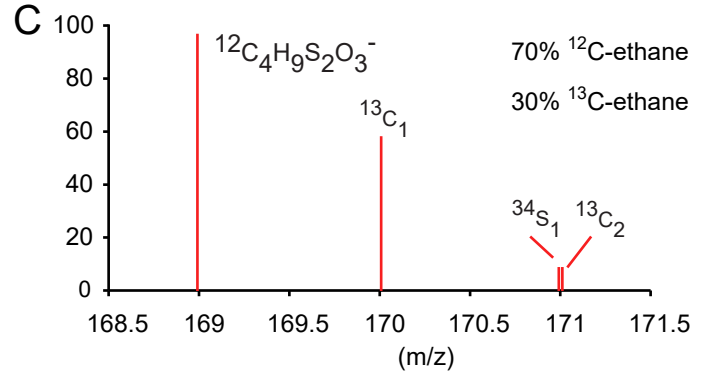
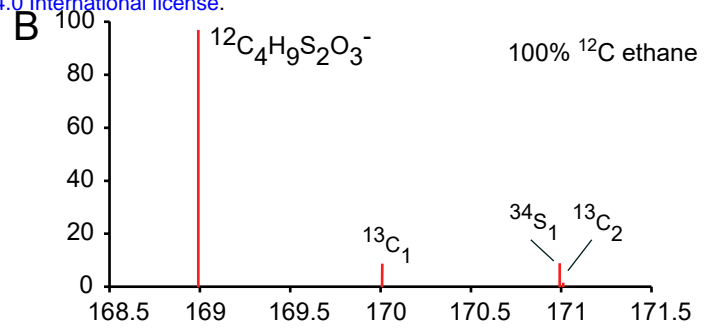
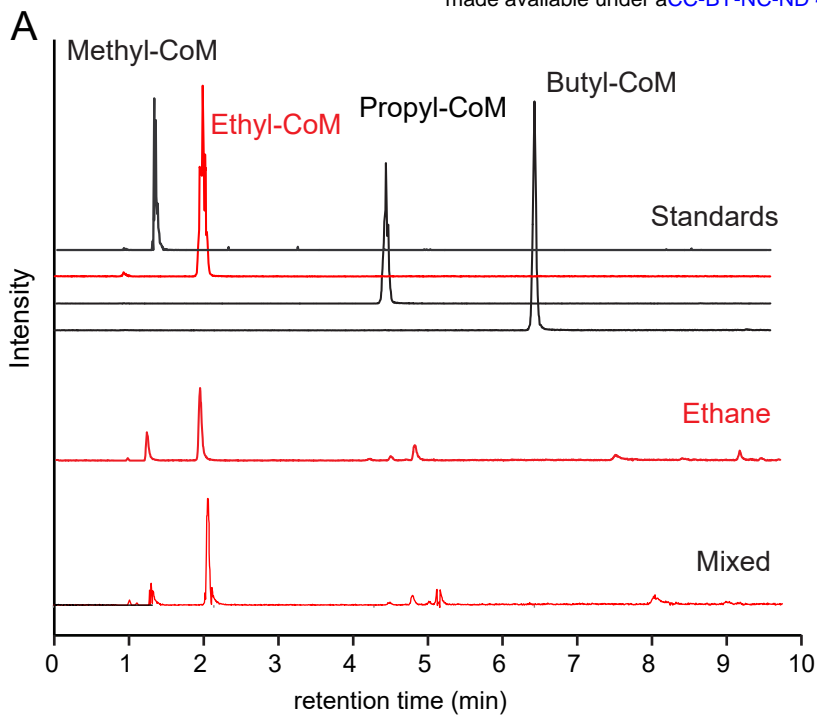
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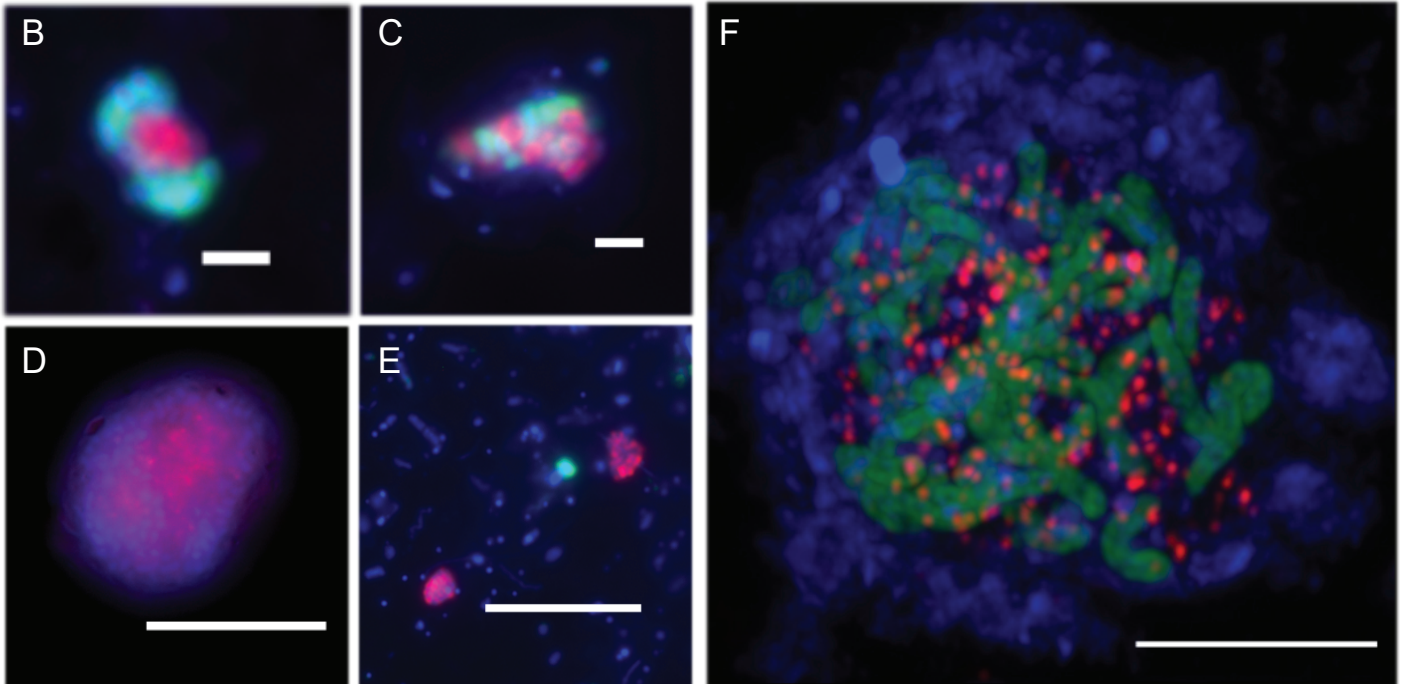
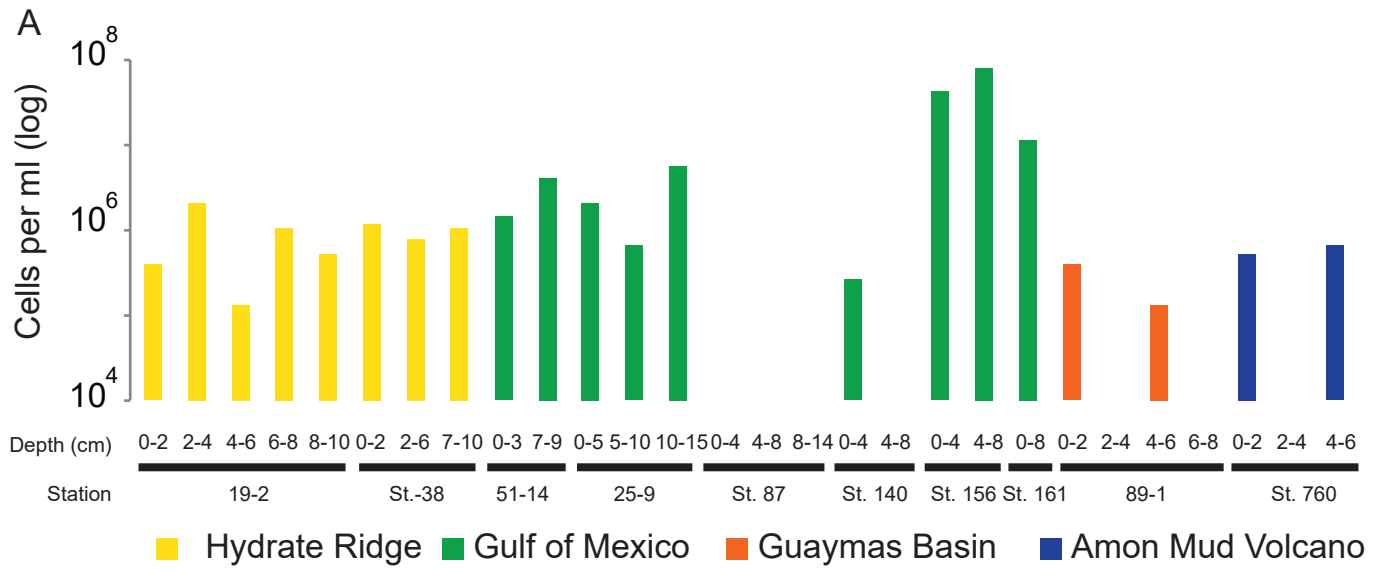


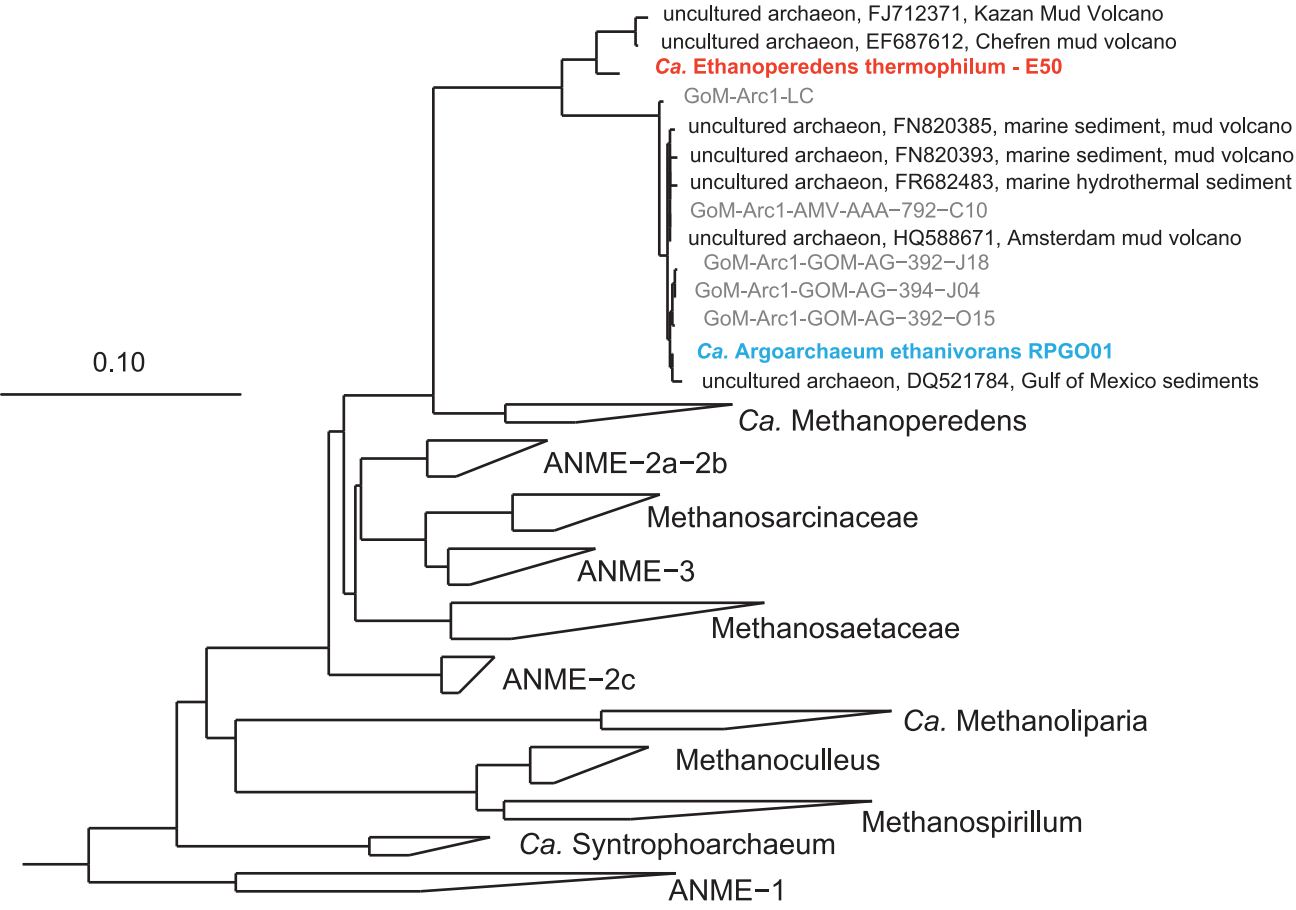
B











A

PRJEB36446 - *Ca. Ethanoperedens thermophilum*
RZB32666.1 - *Ca. Argoarchaeum ethanivorans*

RJS73045.1 - *Ca. Syntrophoarchaeum* sp.
 RJS73394.1 - *Ca. Syntrophoarchaeum* sp.
 RJS72636.1 - *Ca. Syntrophoarchaeum* sp.
 RJS71678.1 - *Ca. Syntrophoarchaeum* sp.

JMIY01000002 - ANME-2d

NJEM01000011 - ANME-1

WP_048108547.1 - *Methanosarcina barkeri*

360 370 380 389

T D Q G W L H N Y L A G G S S G W S N Y I I S V Y T D E V L E D Y
 T D Q G W L H N Y L A G G S S S W A S Y V M S V Y T D E V L E D P

F F Q L Y Y G Q M M I G T - V G A T - V A P A T F Y V N D L Y E S R
 Y L N L W V N G Y V I G G - I G G V - N A G A V I F T N N V L D E F
 Y G L I W F D S Y V L G G - I G G A - N T F S S P F T N N L L D D F
 V D A V W Y G I Y M S G G - V G F P - A F A - Y L Y L G N I L D D W

Y D Q I W L G G Y M S G G - V G F T M Y A T P A Y T N D I L D D Y
 Y D Q L W F G T Y M S G G - V G F T Q Y A S A T Y T D N I L E D F
 Y D Q I W L G S Y M S G G V G F T Q Y A T A A Y T D D I L D N N

B

PRJEB36446 - *Ca. Ethanoperedens thermophilum*
RZB32666.1 - *Ca. Argoarchaeum ethanivorans*

RJS73045.1 - *Ca. Syntrophoarchaeum* sp.
 RJS73394.1 - *Ca. Syntrophoarchaeum* sp.
 RJS72636.1 - *Ca. Syntrophoarchaeum* sp.
 RJS71678.1 - *Ca. Syntrophoarchaeum* sp.

JMIY01000002 - ANME-2d

NJEM01000011 - ANME-1

WP_048108547.1 - *Methanosarcina barkeri*

100									
69.3	100								
36.7	35.4	100							
34.4	32.1	45.3	100						
38.0	37.2	45.5	59.2	100					
30.4	30.1	39.3	37.5	36.9	100				
42.6	39.5	41.8	40.3	39.0	35.2	100			
40.3	37.7	37.2	37.6	38.3	35.8	46.4	100		
43.0	40.2	41.4	41.7	40.6	38.2	69.1	48.5	100	

