

1 Reciprocal carbon subsidies between autotrophs and bacteria in
2 stream food webs under stoichiometric constraints

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4 Benoît O.L. Demars^{1,2}, Nikolai Friberg^{1,3,4}, Joanna L. Kemp¹, Barry Thornton²

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6 ¹ Norwegian Institute for Water Research, Gaustaallen 21, 0349 Oslo, Norway

7 ² The James Hutton Institute, Craigiebuckler, Aberdeen AB15 8QH, UK

8 ³ University of Copenhagen, Freshwater Biological Section, Universitetsparken 4, 3rd floor, 2100 Copenhagen,
9 Denmark

10 ⁴ University of Leeds, water@leeds, School of Geography, Leeds LS2 9JT UK

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15 Corresponding author:

16 Dr Benoît Demars, Norwegian Institute for Water Research, Gaustadalléen 21, 0349 Oslo, Norway

17 Tel.: + 47 98 227 757

18 e-mail: benoit.demars@niva.no

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20 running title: green and brown webs intermingle

21

22 Summary

23 1. Soils are currently leaching out their organic matter at an increasing pace and darkening aquatic
24 ecosystems due to climate and land use change, or recovery from acidification. The implications for
25 stream biogeochemistry and food webs remain largely unknown, notably the metabolic balance (biotic
26 CO₂ emissions), reciprocal subsidies between autotrophs and bacteria, and trophic transfer
27 efficiencies.

28

29 2. We use a flow food web approach to test how a small addition of labile dissolved organic matter
30 affects the strength and dynamics of the autotrophs-bacteria interaction in streams. Our paired streams
31 whole-ecosystem experimental approach combined with continuous whole-stream metabolism and
32 stable isotope probing allowed to unravel carbon fluxes in the control and treatment streams.

33

34 3. We increased the natural supply of dissolved organic matter for three weeks by only 12% by
35 continuously adding 0.5 mg L⁻¹ of sucrose with a δ¹³C signature different from the natural organic
36 matter. Both photosynthesis and heterotrophic respiration increased rapidly following C addition, but
37 this was short lived due to N and P stoichiometric constraints. The resulting peak in heterotrophic
38 respiration was of similar magnitude to natural peaks in the control observed when soils were
39 hydrologically connected to the streams and received soil derived carbon.

40

41 4. Carbon reciprocal subsidies between autotrophs and bacteria in the control stream accounted for
42 about 50% of net primary production and 75% of bacterial production, under low flow conditions
43 when stream water was hydrologically disconnected from soil water. The reciprocal subsidies were
44 weaker by 33% (autotrophs to bacteria) and 55% (bacteria to autotrophs) in the treatment relative to
45 the control. Net primary production relied partly (11% in the control) on natural allochthonous
46 dissolved organic carbon via the CO₂ produced by bacterial respiration.

47

48 5. Many large changes in ecosystem processes were observed in response to the sucrose addition. The
49 light use efficiency of the autotrophs increased by 37%. Ecosystem respiration intensified by 70%,
50 and the metabolic balance became relatively more negative, i.e. biotic CO₂ emissions increased by
51 125%. Heterotrophic respiration and production increased by 89%, and this was reflected by a shorter
52 (-40%) uptake length (S_{wOC}) and faster (+92%) mineralisation velocity of organic carbon. The
53 proportion of DOC flux respired and organic carbon use efficiency by bacteria increased by 112%.

54

55 6. Macroinvertebrate consumer density increased by 72% due to sucrose addition and consumer
56 production was 1.8 times higher in the treatment than in the control at the end of the experiment. The
57 trophic transfer efficiencies from resources to consumers were similar between the control and the
58 treatment (2-5%).

59

60 7. *Synthesis*. Part of the carbon derived from natural allochthonous organic matter can feed the
61 autotrophs via the CO₂ produced by stream bacterial respiration, intermingling the green and brown
62 webs. The interaction between autotrophs and bacteria shifted from mutualism to competition with
63 carbon addition under nutrient limitation (N, P) increasing biotic CO₂ emissions. Without nutrient
64 limitation, mutualism could be reinforced by a positive feedback loop, maintaining the same biotic
65 CO₂ emissions. A small increase in dissolved organic carbon supply from climate and land use change
66 could have large effects on stream food web and biogeochemistry with implications for the global C
67 cycle under stoichiometric constraints.

68

69

70 *Key-words*: reciprocal subsidies, microbial loop, dissolved organic matter, nutrient spiralling, whole-
71 stream metabolism, flow food web

72

73

74 **Introduction**

75 The global annual riverine flux of organic C ($0.26\text{-}0.53 \text{ Pg C year}^{-1}$) to the oceans is comparable to
76 the annual C sequestration in soil ($0.4 \text{ Pg C year}^{-1}$), suggesting that terrestrially derived aquatic losses
77 of organic C may contribute to regulating changes in soil organic carbon storage (Dawson 2013).
78 Soils are currently leaching out their organic matter at faster rates to aquatic ecosystems due to
79 climate and land use change, or recovery from acidification (e.g. Freeman et al. 2004, Monteith et al.
80 2007, Drake et al. 2015). The fate of this organic carbon in riverine systems remains poorly
81 understood at the global scale, notably the degassing of CO_2 back to the atmosphere (Drake et al.
82 2018), and remains highly debated regarding its contribution to aquatic food webs (e.g. Cole 2013,
83 Brett et al. 2017).

84

85 The bacterial processing rate of terrestrial organic matter is influenced by the strength and complexity
86 (aromaticity) of the carbon bonds of the organic matter and C:N:P ecological stoichiometry (Mineau
87 et al. 2016, Evans et al. 2017, Kominoski et al. 2018). Numerous labile dissolved organic carbon
88 (DOC) additions (from trace amount to 20 mg C L^{-1}) have repeatedly shown DOC use by bacteria and
89 its transfer through the food chain (e.g. Warren et al. 1964, Hall 1995, Hall and Meyer 1998, Parkyn
90 et al. 2005, Wilcox et al. 2005, Augspurger et al. 2008). Fewer studies used leaf leachate material
91 including less labile DOC (e.g. Cummins et al. 1972, Friberg and Winterbourn 1996, Wiegner et al.
92 2005, Wiegner et al. 2015). Many studies have also traced the flux of autochthonous DOC through to
93 the bacteria (e.g. Lyon and Ziegler 2009, Risse-Buhl et al. 2012, Hotchkiss et al. 2014, Kuehn et al.
94 2014).

95

96 The interactions between autotrophs and bacteria are however difficult to study (Amin et al. 2015),
97 because primary producers, decomposers and organic matter (allochthonous and autochthonous) are
98 intricately connected both in benthic biofilms (Kamjunke et al. 2015, Battin et al. 2016) and pelagic
99 aggregates (Grossart 2010). In theory, bacteria and autotrophs could compete for limiting nutrients
100 (Currie and Kalff 1984), notably when bacteria have lower C:nutrient biomass ratios than autotrophs
101 (Daufresne and Loreau 2001). However, bacteria can have high C:nutrient ratios similar to autotrophs
102 (Cotner et al. 2010) supporting the co-existence of autotrophs and heterotrophs (Daufresne and
103 Loreau 2001). Correlative analyses in streams seem to support the idea that positive interactions
104 between autotrophs and bacteria increase with nutrient (N, P) limitation (Carr et al. 2005, Scott et al.
105 2008). Whole-stream metabolism in open streams under low flow conditions (hydrologically
106 disconnected from catchment soils supplying allochthonous organic carbon) showed ecosystem
107 respiration to be tightly correlated to gross primary production (Demars et al. 2016), suggesting a
108 strong indirect mutualistic interaction between autotrophs and bacteria (Demars et al. 2011a), with
109 autotrophs providing detritus C, N, P to bacteria and bacteria regenerating N and P by mineralisation
110 (Cotner et al. 2010, Demars et al. 2011b), in agreement with theory (Daufresne and Loreau 2001).

111

112 The potential benefit of bacterial CO_2 for primary producers has been hypothesised to explain a small
113 increase (16-20%) in gross primary production following dissolved organic carbon addition and
114 associated increase in bacterial activities (Robbins et al. 2017). The reciprocal carbon subsidies
115 between autotrophs and bacteria in stream food webs has however, to our knowledge, not been
116 studied either empirically or theoretically. It raises the possibility that the carbon of the primary
117 producers may be partly derived from allochthonous organic matter processed by the bacteria, and
118 intermingle the green and brown webs (Zou et al. 2016), an overlooked issue in the autochthony-
119 allochthony debate (e.g. Cole 2013, Brett et al. 2017).

120

121 Here we test how dissolved organic carbon addition affects the strength and dynamics of the
122 autotrophs-bacteria interaction in streams – see Figure 1. More specifically, we characterised the
123 effect of enhancing or displacing natural DOC benthic uptake by adding a small flux of labile organic
124 carbon (Lutz et al 2012) and estimating in-stream ecosystem C fluxes under potential stoichiometric
125 constraints (C:N:P) using a flow food web approach (*sensu* Marcarelli et al. 2011).

126 The addition of labile organic carbon (sucrose) may reduce the mineralisation of N and P and increase
127 allochthonous nutrient demand by bacteria, reducing nutrient availability for autotrophs. If nutrients
128 are limiting, then C reciprocal subsidies between bacteria and autotrophs may be weakened by sucrose
129 addition, i.e. there could be a shift between mutualism and competition. With enough nutrients the C
130 microbial loop (to autotrophs) may be strengthened via an increase in CO₂ from bacterial respiration,
131 i.e. positive feedback loop. Allochthonous nutrients may also come from allochthonous OM so
132 bacteria may be under less nutrient constraints than algae, although the nutrient content of DOC
133 (C:N:P=3200:103:1, Stutter et al. 2013) is generally much poorer than benthic algae
134 (C:N:P=158:18:1, Kahlert 1998). The metabolic balance is likely to shift towards heterotrophy. A
135 large amount of sucrose is expected to be respired, so CO₂ emissions are predicted to increase.
136 Consumers and the trophic transfer efficiency should benefit from an increase in bacterial production.
137 The design also allows to quantify the potential of priming of allochthonous OM by sucrose via
138 bacterial activity (Kunc et al. 1976, Hotchkiss et al. 2014).

139 We combined whole-ecosystem stable isotope probing (using sucrose from sugar cane) and Bayesian
140 mixing model to characterise the carbon links (sources to mixtures). We converted the relative carbon
141 fluxes from the different sources, into carbon fluxes by estimating the production of the mixtures,
142 after taking into account assimilation efficiencies (e.g. carbon use efficiency, bacterial growth
143 efficiency). This approach relies on the integration of stream metabolism, nutrient cycling, ecological
144 stoichiometry, stable isotope probing of the food web and production estimates (see Welti et al. 2017).
145 In this study, we focus on the basal part of the food web, and notably on the elusive reciprocal C
146 subsidies between autotrophs and bacteria.

147

148 **Methods**

149 *Study area*

150 We studied two heather moorland catchments with soils rich in organic carbon, within the Glensauigh
151 research station of the James Hutton Institute in north-east Scotland (Long 2° 33' W, Lat 57° 55' N) –
152 Fig. 2. The streams were about 0.8-1.0 m wide in the studied sections and their channels significantly
153 undercut the banks by 30-46% of stream width. Brown trout (*Salmo trutta fario*, Salmonidae) was
154 present in both streams. The management of the land includes regular heather burning (10-12% of
155 surface area yearly target) for hill farming: mixed grazing of sheep and cattle. For further information
156 see Demars (2018).

157

158 *The control stream*

159 The control stream (Birnie Burn) is part of the long-term monitoring of the UK Environmental
160 Change Network (ECN, <http://data.ecn.ac.uk/>). There is a monitoring of soil temperature and moisture
161 on the hillslope of the Birnie Burn at 275 m elevation (Cooper et al. 2007). Volumetric soil moisture
162 content is recorded every 30 minutes at 10 and 45 cm depth, corresponding respectively to the base of
163 the O (organic layer) and B (subsoil) horizons of the humus iron podzol present. The stream is
164 equipped with a flume for continuous monitoring of discharge (catchment area 0.76 km²) and dip

165 water samples are collected weekly. This stream showed substantial increase in annual flow-weighted
166 mean concentrations of stream water DOC (+0.28 mg C L⁻¹ year⁻¹ during 1994-2007, Stutter et al.
167 2011). Water samples were analysed for DOC, pH, nutrients (N, P) and major ions. – see Cooper et al.
168 2007, Stutter et al. 2012 and Demars 2018 for further details.

169

170 *The paired stream*

171 The control stream was paired with a neighbouring stream (Cairn Burn) in 2005, also part of a long-
172 term monitoring scheme with samples collected every week or two for stream water quality. In the
173 late 1970s and early 80s two areas covering 33 ha were improved (reseeded, limed and fertilized) as
174 part of sheep grazing experiments (Hill Farming Research Organisation 1983). The added facilities at
175 the Cairn Burn (catchment area 0.9 km²) include a calibrated flume, water level, water electric
176 conductivity, water and air temperature and barometric pressure. Data were recorded every 5 minutes
177 (Campbell Scientific CR10x datalogger). Data logger, battery and barometric pressure were housed in
178 a weather resistant enclosure. Photosynthetic active radiations (PAR) were also recorded in air, one
179 metre above ground, at the same time intervals (LICOR, Lincoln, NE, USA). For more information,
180 see Demars (2018).

181

182 *Terrestrial DOC: main source of organic carbon*

183 DOC was the dominant flux of organic carbon (98%) under stable flows with average concentrations
184 of 9.3±1.7 mg C L⁻¹ in the two studied streams (Demars 2018). This DOC was of terrestrial origin as
185 shown by δ¹³C analyses of the natural DOC against terrestrial and aquatic plant material (Stutter et al.
186 2013). The median molar C:N:P stoichiometry of the DOC was 3201:103:1, with values ranging
187 between 978:38:1 to 12013:282:1 (Stutter et al. 2013). Chlorophyll a concentration in the water
188 column was extremely low (< 1 µg L⁻¹, Stutter et al. 2013). The pool of particulate organic carbon in
189 the sediment is very small (Demars 2018), and coarse particulate organic carbon was less than 10 g C
190 m⁻² (determined from Surber sampling of invertebrates, see below).

191

192 *DOC addition*

193 A carboy was refilled every two days with 6 kg of sucrose dissolved in over 60 L of stream water
194 filtered through muslin square in a large funnel. The carboy was set as a Mariotte bottle to ensure a
195 constant dripping rate of 22 mL min⁻¹ lasting 48 hours (15 mg C s⁻¹). The ventilation tube was netted
196 at the top to avoid insect contamination. The dripping rate was kept constant over the 22 days of
197 sucrose addition (23 August - 14 September) and was initially set to increase stream DOC
198 concentration by about 0.5 mg C L⁻¹ at 30 L s⁻¹. Sample were collected in washed bottles and filtered
199 on-site with pre-washed filters (0.45 µm Millipore PVDF membrane filter). DOC was determined
200 within 48 hours of collection with a Skalar San++ continuous flow analyzer (Breda, The Netherland),
201 using potassium hydrogen phthalate as standards and sodium benzoate for quality controls. The
202 detection limit was 0.1 mg C L⁻¹.

203

204 *Nutrient cycling studies*

205 Nutrient cycling studies were run in the control and manipulated streams before and during sucrose
206 addition. Nutrient cycling rates were derived from continuous *in-situ* nutrient addition experiments
207 where a conservative tracer is also included (Stream Solute Workshop 1990, Demars 2008). This

208 method tends to overestimate the nutrient uptake length S_w , average distance travelled by a nutrient
209 molecule in the water column before river bed uptake. This is due to the addition of nutrient
210 compared to isotopic tracer studies (e.g. Mulholland et al. 2002). However, preliminary tests in the
211 Cairn Burn showed that the bias can be kept small (10-15%) with small nutrient additions (Demars
212 2008). Nitrate (as KNO_3) and phosphate (as KH_2PO_4) were continuously added together with NaCl as
213 conservative tracer (*cf* Schade et al. 2011). When the plateau phase was reached, water samples were
214 collected at about 10 m interval along the reach and filtered on site (pre-washed 0.45 μm Millipore
215 PVDF membrane filter) (see Demars 2008). The samples were kept cool at 4-10°C. The nutrients
216 (PO_4 and NO_3) were determined within 48 hours by colorimetry using a Skalar San++ continuous
217 flow analyzer (Breda, The Netherland) and chloride by ion chromatography (Dionex DX600,
218 Sunnyvale, California). The limits of detections were 0.001 for NO_3 and PO_4 and 0.003 mg L^{-1} for Cl.
219 In order to provide a more comparable indicator of nutrient cycling for different hydrological
220 conditions, the uptake velocity v_f was also calculated as follows: $v_f = uz/S_w$, with u average water
221 velocity and z average depth. Short uptake lengths and fast uptake velocities indicate fast cycling rates
222 (high exchange rates between water and benthos).

223

224 *Whole-stream metabolism*

225 Whole-stream metabolism was estimated by the open channel two-station diel oxygen method of
226 Odum (1956) modified by Demars et al. 2011b, 2015, 2017 and Demars (2018). Many tracer studies
227 (using NaCl and propane) were carried out as detailed in Demars et al. (2011b) to estimate lateral
228 inflows, mean travel time and reaeration coefficient as a function of discharge within the range of
229 stable flows (up to 32 L s^{-1}). The relationships with discharge were very strong (Fig 3, Demars 2018)
230 allowing accurate parameterisation of metabolism calculations under varying flow conditions as in
231 Roberts et al. (2007) and Beaulieu et al. (2013). The high oxygen reaeration coefficient of those
232 streams (0.05-0.24 min^{-1}) required very accurate dissolved O_2 data. Oxygen concentrations were
233 measured with optic sensors fitted on multiparameter sondes TROLL9500 Professional (In-Situ Inc.,
234 Ft Collins, CO, USA) and Universal Controller Sc100 (Hach Lange GMBF), the latter powered with
235 two 12 V DC (75mA) car batteries per sensor kept charged with two 20 W solar panels (SP20
236 Campbell Scientific). The sensors were calibrated to within 1% dissolved oxygen saturation. Four
237 sondes were deployed in the Cairn Burn at 0, 84, 138, 212 m upstream of the flume to include an extra
238 control reach (138-212 m) upstream of the manipulated reach (0-84 m). Another two sondes were set
239 in the control stream Birnie Burn at 88 m interval (60-148 m upstream of the ECN flume). The
240 distances between oxygen stations corresponded to 80-90% of the oxygen sensor footprints ($3u/k_2$),
241 with u/k_2 entirely independent of discharge ($R^2=0.0005$), which allowed the manipulated reach to be
242 independent of the control reach. The DOC injection point was 28 m upstream of the top station of the
243 manipulated reach, and this distance corresponded to 69% of the oxygen sensor footprint of the top
244 station. All sondes were deployed from May to October 2007, logging at 5 min time step interval.

245 The net metabolism was only calculated for stable flow conditions (3-32 L s^{-1}), as follows (Demars
246 2018):

$$247 \quad NEP_t = \left(\frac{C_{AV\ t+\Delta t} - C_{AV\ t}}{\Delta t} - k_2(C_s - C_{AV\ t}) - \frac{\theta(C_g - C_{AV\ t})}{\Delta t} \right) z$$

248 with NEP_t net ecosystem production at time t ($\text{g O}_2 \text{ m}^{-2} \text{ min}^{-1}$), C_{AV} average dissolved oxygen ($\text{g O}_2 \text{ m}^{-3}$)
249 of the two stations at time $t+\Delta t$ and t (min), Δt time interval (min), k_2 oxygen exchange coefficient
250 (min^{-1}), C_s saturated oxygen concentration ($\text{g O}_2 \text{ m}^{-3}$), θ the proportion of lateral inflows, z average
251 stream depth (m), C_g oxygen concentration in lateral inflows ($\text{g O}_2 \text{ m}^{-3}$). The latter was calculated as
252 follows (from baseflow analysis of stream hydrographs):

$$253 \quad C_g = (1/(1 + \exp(a \ln(Q) - b))0.9C_s) + (1 - 1/(1 + \exp(a \ln(Q) - b))0.1C_s)$$

254 with Q discharge, a and b constants, permitting to correct for baseflow (first term of the equation) and
255 soil water (second term of the equation) lateral inflows, see Demars (2018). The proportion (\pm se) of
256 total lateral inflows relative to discharge (Q_g/Q) was $10.7\pm 0.6\%$, $6.6\pm 0.5\%$, and $2.3\pm 0.4\%$ for the
257 Birnie Burn control, Cairn Burn control and treated reach, respectively, independently of discharge in
258 the range $3.8\text{--}32.5\text{ L s}^{-1}$ (stable flows).

259 All calculations were run in Excel using a preformatted spreadsheet (Demars 2018). The overall
260 uncertainties in daily stream metabolism, including cross-calibration errors, individual parameter
261 uncertainties, spatial heterogeneity (through the average of diel O_2 curves) and correction for lateral
262 inflows, were propagated through all the calculations using Monte Carlo simulations (Demars 2018).
263 The corrections for lateral inflows amounted to about 6% of ER for the treated reach (Cairn Burn),
264 19% and 16% in the control reaches, Cairn Burn and Birnie Burn, respectively.

265

266 *Identification of carbon sources and pathways*

267 *Macrophytes.* The percentage cover of bryophytes and filamentous green algae was measured with a
268 ruler across transects taken every two metres along the stream reaches. Young shoots of filamentous
269 green algae and bryophytes were collected by hand along both studied reaches before and after
270 sucrose addition. All samples were freeze dried and milled prior to analyses for C, N, $\delta^{13}C$ and $\delta^{15}N$.
271 The main source of inorganic carbon for primary producers was assumed to be CO_2 because of the
272 low alkalinity (remaining below $0.5\text{ meq HCO}_3\text{ L}^{-1}$ under low flows). The fractionation factor for CO_2
273 assimilation into macrophyte tissue is known to vary with pCO_2 and growth rate, and was set at -
274 $25.5\pm 3.5\%$ within the range of Rubisco forms IA and IB in the absence of carbon concentrating
275 mechanism and transport limitation (22–29%, e.g. Raven et al. 1994, McNeven et al. 2007, Boller et
276 al. 2015). Net primary production was assumed to be driven by filamentous green algae and biofilm
277 autotrophs, and bryophytes were not used to calculate the flow food webs.

278

279 *Inorganic carbon.* The $\delta^{13}C$ of dissolved CO_2 is known to be variable (Findlay 2004, Billett and
280 Garnett 2010, Billett et al. 2012). The $\delta^{13}C$ of dissolved CO_2 reflects both the signature of terrestrial C
281 from groundwater and soil water inflows as well as in-stream processes (biotic respiration) and CO_2
282 exchange with the atmosphere ($\delta^{13}CO_2$ about -8‰, e.g. Billett et al. 2012). In the absence of direct
283 measurements, we considered estimating $\delta^{13}C$ of dissolved CO_2 from available data of $\delta^{13}C$ of
284 dissolved inorganic carbon (DIC) in the Brocky Burn, a small stream in the adjacent catchment
285 (Waldron et al. 2007). The equilibrium method used to derive the $\delta^{13}C$ of CO_2 from the $\delta^{13}C$ of DIC
286 (Zhang et al. 1995) was not applicable because the time to reach the carbonate equilibrium (20–200 s)
287 approached the average time spent by a CO_2 molecule in the stream before emission in the atmosphere
288 (300–1000 s, calculated from travel time and reaeration coefficient of CO_2 , see below). The alternative
289 evasion method (Zhang and Quay 1997) did not seem more appropriate (Billett and Garnett 2010).

290 We suggest another approach: $\delta^{13}C$ of dissolved CO_2 may be estimated indirectly under low flows
291 using the fractionation factor of Rubisco $-25.5\pm 3.5\%$ and $\delta^{13}C$ of bryophytes (strict CO_2 user and no
292 CO_2 transport limitation). In our study, the average $\delta^{13}C$ of bryophytes was -36.4‰, and assuming the
293 above fractionation coefficient of -25.5‰, Glensauigh $\delta^{13}C$ of dissolved CO_2 would be -10.9‰. This
294 is similar to the $\delta^{13}C$ of DIC (6–14‰ under low flows, Waldron et al. 2007) and $\delta^{13}C$ of bryophytes (-
295 33.3‰, Palmer et al. 2001) reported for the Brocky Burn. The proof of concept comes from another
296 Scottish stream (Dighty Burn), where Raven et al. (1994) reported the $\delta^{13}C$ of dissolved CO_2 as -
297 14.7‰ and bryophytes as -39.2‰, suggesting a fractionation coefficient of -24.5‰ by difference. In
298 our calculations under low flow conditions we therefore assumed $\delta^{13}C$ of dissolved CO_2 as $-11\pm 3\%$.

299 To quantify the reciprocal subsidies between autotrophs and bacteria, it remained to decompose the
300 overall stable isotope signature of stream dissolved CO₂ into the allochthonous (groundwater, soil
301 water and atmospheric exchange) and autochthonous (respiration by heterotrophs and autotrophs)
302 sources. The allochthonous signature, $\delta^{13}C_{CO_2\text{-allochthonous}}$, can be deduced from rearranging:

$$\begin{aligned} 303 \quad F_{CO_2} \times \delta^{13}C_{CO_2} &= F_{CO_2\text{-allochthonous}} \times \delta^{13}C_{CO_2\text{-allochthonous}} \\ 304 \quad &+ F_{CO_2\text{-heterotrophs}} \times \delta^{13}C_{CO_2\text{-heterotrophs}} \\ 305 \quad &+ F_{CO_2\text{-autotrophs}} \times \delta^{13}C_{CO_2\text{-autotrophs}} \end{aligned}$$

306 where F_{CO_2} represents CO₂ fluxes (g C m⁻² day⁻¹, with all fluxes expressed as positive values) and $\delta^{13}C$
307 the isotope signature (‰) of the different sources of CO₂. We averaged the $\delta^{13}C$ of the autotrophs
308 (filamentous green algae and biofilm primary producers). The $\delta^{13}C_{CO_2\text{-allochthonous}}$ was only calculated
309 for the control stream under low stable flows and assumed to apply to both streams. Uncertainties
310 were propagated in quadrature using standard deviation δx for sums, and relative uncertainties $\delta x/x$ for
311 the division.

312

313 *Sucrose*. The $\delta^{13}C$ of sucrose from sugar cane is similar to that of dissolved CO₂ (-12±1‰, Jähren et
314 al. 2006, Augspurger et al. 2008, Kankaala et al. 2010, de Castro et al. 2016), but sucrose uptake by
315 autotrophs was assumed to be without isotopic discrimination (Wright and Hobbie 1966). The
316 proportion of carbon derived from added sucrose (F_S) in resources and consumers was calculated from
317 their $\delta^{13}C$ in the control (C) and treatment (T) reaches, before (B) and after (A) sucrose addition as
318 follows:

$$319 \quad F_S = \frac{T_A - (T_B + (C_A - C_B))}{-12 - (T_B + (C_A - C_B))}$$

320 with all uncertainties propagated in quadrature using standard deviation δx for sums, and relative
321 uncertainties $\delta x/x$ for the division. The standard error of the mean was calculated as $sem = \delta x/\sqrt{n}$
322 with n average number of samples in C_B , C_A , T_B , T_A .

323

324 *Allochthonous organic carbon*. The $\delta^{13}C$ of the DOC (average ±SD) was available from a previous
325 study from the same catchment and showed it was of terrestrial origin, i.e. not autochthonous ($\delta^{13}C = -$
326 28.5 ± 0.3 ‰, Stutter et al. 2013, see above). Coarse particulate organic matter (CPOM) was also
327 collected by hand along both studied reaches before and after sucrose addition. Since there was hardly
328 any difference in $\delta^{13}C$ between DOC and CPOM ($\delta^{13}C = -27.4 \pm 0.7$ ‰, Table S1), we used the $\delta^{13}C$ of
329 CPOM determined in this study as the signature for allochthonous organic carbon.

330

331 *Periphyton*. Periphyton (or biofilm) samples represent a mixture of primary producers (algae and
332 cyanobacteria), bacteria and fine particulate organic matter. The samples were collected before and at
333 the end of the sucrose addition from the flumes and stones with a toothbrush, funnel and bottle. All
334 samples were freeze dried and milled.

335 We also placed six pairs (with/without Vaseline) of unglazed ceramic tiles (10 x 10 cm) fixed on
336 bricks and deployed along the studied reaches in both streams three weeks before the start of the
337 manipulation. Vaseline was applied around half the tiles to prevent grazing by invertebrates. After
338 three weeks, there was hardly any growth on the tiles, and so the tiles were left in the stream until the
339 end of the manipulation. One brick in the control stream was lost. At the end of the experiment, the

340 tiles were frozen at -20°C, later freeze dried and the biofilm was scraped with a razor blade. Since
341 there was little biomass per tile (about 1 g C m⁻²), the biofilm was pooled by stream and grazer
342 treatments (leaving two samples per stream). Very little grazing activity was observed on the tiles
343 during the six weeks and unsurprisingly no difference in biofilm dry mass emerged due to grazer
344 exclusion (paired t-tests on ln transformed mass; Birnie, t₄, p=0.13; Cairn, t₅, p=0.26).

345 Phospholipid fatty acids (PLFAs) were extracted from the biofilm samples from the tiles and
346 derivatised to their methyl esters (FAMES) following the procedure of Frostegård et al. (1993) using
347 the modified extraction method of Bligh and Dyer (1959), as detailed in Certini et al. (2004).
348 Quantification and δ¹³C values of the PLFAs were both determined by Gas Chromatography-
349 Combustion-Isotope Ratio Mass Spectrometry (GC-C-IRMS) as described by Main et al. (2015), and
350 averaged for each stream (δ¹³C, Table S2). Only PLFAs up to 19 carbon chain length were determined
351 (excluding some long-chain essential polyunsaturated fatty acids, e.g. Muller-Navarra et al. 2000,
352 Gladyshev et al. 2011).

353

354 *Bacterial carbon.* The bacterial PLFAs were identified as those most affected by sucrose addition in
355 the treated reach and information derived from the literature (supplementary information). This also
356 allowed to determine the δ¹³C of bacteria in the control reach. We used a fractionation factor of -3‰
357 for the δ¹³C of bacterial fatty acids relative to bulk tissue samples (Hayes 2001, Gladyshev et al.
358 2014). This is within the range of observed values in other studies (Boschker and Middelburg 2002,
359 Bec et al. 2011). We used the same fractionation factor of -25.5±3.5‰ for the assimilation of CO₂
360 coming from bacterial respiration into green algal tissue. Since we only had comparative data for the
361 treatment period, the fraction of sucrose within PLFAs was calculated as follows:

362
$$F_s = \frac{T_A - C_A}{-12 - 3 - C_A}$$

363

364 *Biofilm autotroph carbon.* The carbon from cyanobacteria and algae were identified from a specific
365 PLFA (α-linolenic acid, Risse-Buhl et al. 2012) using the same fractionation factor as above between
366 PLFA and bulk tissue (Table S2).

367

368 *Macroinvertebrate consumers.* Macroinvertebrate densities were estimated from twelve to thirteen
369 Surber samples (20 x 20 cm, mesh size 200 μm) collected randomly along the reaches (total 51
370 samples). The samples were stored in 70% alcohol, sorted and identified. Macroinvertebrate for CN
371 and stable isotope studies were collected by kick sampling and hand net. The animals were sorted and
372 identified live within a day into Petri dishes and after allowing time for gut evacuation, placed in
373 Eppendorf tubes and freeze dried. The average biomass of macroinvertebrate taxa was assessed by
374 weighing the freeze-dried mass of all individuals within a tube divided by the number of animals. The
375 whole macroinvertebrates were then crushed before insertion into a tin capsule.

376

377 *Carbon isotopic turnover.* Bacteria and algae were likely to be fully labelled within a week, and three
378 weeks of sucrose addition was thought to be sufficient to label invertebrates, albeit not fully for all
379 consumer species (especially predators), with carbon turnover ranging from about 10-35 days (Le
380 Cren and Lowe-McConnell 1980, Hall and Meyer 1998, Collins et al. 2016 and references therein),
381 and a time lag to reach equilibrium at each trophic level (e.g. the time for bacteria and algae to reach
382 full equilibrium, likely to be less than a week, will delay the time grazers may reach their full
383 equilibrium). The proportion of tissue turnover over 14 and 21 days was estimated from the consumer

384 individual body mass M (g) of the taxa (assuming freeze-dried mass = 0.2 fresh mass, Waters 1977
385 cited in Wetzel 2001, p. 718) with tissue isotopic turnover rate λ (day^{-1}) derived from the isotopic
386 half-life study of Vander Zanden et al. (2015), general equation for invertebrates:

$$387 \quad \lambda = \frac{\ln(2)}{25.8 M^{0.23}}$$

388 and from the remaining individual body mass M_t at time t :

$$389 \quad M_t = M \exp(-\lambda t)$$

390 This provided an indication to what extent the isotopic signature of consumers may have reached an
391 equilibrium. We used the proportion of macroinvertebrate tissue turnover over 21 days to provide the
392 fraction of sucrose at equilibrium for individual taxa and grouped by functional feeding groups
393 following Demars et al. (2012).

394

395 *Analytical methods.* The total carbon and total nitrogen concentrations and the $\delta^{13}\text{C}$ and $\delta^{15}\text{N}$ natural
396 abundance isotope ratios of the milled samples were determined using a Flash EA 1112 Series
397 Elemental Analyser connected via a ConFlo III to a Delta^{Plus} XP isotope ratio mass spectrometer (all
398 Thermo Finnigan, Bremen, Germany). The isotope ratios were traceable to reference materials
399 USGS40 and USGS41 (both L-glutamic acid); certified both for $\delta^{13}\text{C}$ (‰_{VPDB}) and $\delta^{15}\text{N}$ ($\text{‰}_{\text{air N}_2}$). The
400 carbon and nitrogen contents of the samples were calculated from the area output of the mass
401 spectrometer calibrated against National Institute of Standards and Technology standard reference
402 material 1547 peach leaves which was analysed with every batch of ten samples. Long term
403 precisions for a quality control standard (milled flour) were: total carbon $40.3 \pm 0.35 \%$, $\delta^{13}\text{C}$ $-25.4 \pm$
404 0.13‰ , total nitrogen $1.7 \pm 0.04 \%$ and ^{15}N $0.367 \pm 0.0001 \text{ atom } \%$ (mean \pm sd, $n = 200$). Data
405 processing was performed using Isodat NT software version 2.0 (Thermo Electron, Bremen,
406 Germany) and exported into Excel. Total P was determined after 30 min digestion in 50% nitric acid
407 at 120°C for CPOM, bryophytes, biofilm and green filamentous algae (see Demars and Edwards
408 2007).

409 *Data analyses.* Most studies use $\delta^{13}\text{C}$ and $\delta^{15}\text{N}$ to identify the flow path in the food web. Here the
410 BACI experimental design allowed to calculate the proportion of sucrose (Fs) in all parts of the food
411 web and was used as a tracer in addition to $\delta^{13}\text{C}$ to determine the sources of carbon for bacteria and
412 algae in the treatment reach after 21 days of sucrose addition. Thus, the carbon pathways were
413 identified with carbon tracers. End member mixing analyses were used to determine the proportion of
414 C sources and their uncertainties in primary producers and bacteria. We provided the numerical
415 solutions given by a Bayesian stable isotope mixing model SIAR 4.2.2 (Parnell et al. 2010) in R
416 version 3.1.3 (R Core Team 2015). The numerical solutions of SIAR were very similar to the
417 analytical solutions of IsoError 1.04 (Phillips and Gregg 2001), and suggested the results were not
418 biased (*cf* Brett 2014). Note, the relative importance of individual resources to consumers was outside
419 the scope of this study.

420

421 *Quantification of carbon fluxes*

422 To assess trophic transfer efficiency through the food web, our production estimates were all
423 standardized to $\text{g C m}^{-2} \text{ day}^{-1}$. Respiration and photosynthesis rates in oxygen were converted to
424 carbon using a respiratory and photosynthetic quotient of 1 (Williams and del Giorgio 2005, but see
425 Berggren et al. 2012).

426

427 *Total CO₂ emissions.* In the absence of direct measurements, the excess partial pressure of CO₂
428 (EpCO₂) of the streams was estimated from three measured parameters: pH, alkalinity and
429 temperature (Neal et al. 1998, as applied in Demars et al. 2016 with atmospheric CO₂=384 ppm,
430 ftp://ftp.cmdl.noaa.gov/ccg/co2/trends/co2_annmean_mlo.txt). Our EpCO₂ estimates have high
431 uncertainties (±50%, Demars 2018). EpCO₂ is the concentration of free CO₂ in the stream water (C_t at
432 time t) relative to the atmospheric equilibrium free CO₂ concentration (C_{SAT}):

$$433 \quad EpCO_2 = C_t / C_{SAT}$$

434 C_{SAT} was calculated from published CO₂ solubility in pure water at equilibrium with atmospheric CO₂
435 in the temperature range 0-90°C (Carroll et al. 1991) and Henry's law (Stumm and Morgan 1981,
436 Butler 1982). C_t was calculated as EpCO₂ × C_{SAT} . The flux of CO₂ (F_{CO_2} , g C m⁻² day⁻¹) at the
437 interface between water and the atmosphere was calculated as for O₂ following Young and Huryn
438 (1998):

$$439 \quad F_{CO_2} = k_{CO_2} (C_{SAT} - C_t) \tau \frac{Q}{A}$$

440 with k_{CO_2} reaeration coefficient of CO₂ (day⁻¹), C_{SAT} and C_t (mg C L⁻¹ or g C m⁻³), τ mean travel time
441 of the stream reach (day), Q average water discharge (m³ day⁻¹), A surface water area of the stream
442 reach (m²). The reaeration coefficients between CO₂ and O₂ can be simply related as follows (Demars
443 et al. 2015, Demars 2018):

$$444 \quad k_{CO_2} = \frac{Dm_{CO_2}}{Dm_{O_2}} k_{O_2} = 0.81 \pm 0.01 k_{O_2}$$

445 based on the molecular diffusivity (Dm) of CO₂ and O₂ measured at three different temperatures
446 within the same study (Davidson and Cullen 1957).

447 The flux of CO₂ was then related to discharge within the range of low stable flows for which stream
448 metabolism was processed (Cairn n=47, R²=0.81; Birnie n=65, R²=0.77) to provide daily estimates.
449 For more details, see Demars 2018.

450

451 *Biotic CO₂ emissions.* These were simply calculated as the net ecosystem production (NEP), gross
452 primary production (GPP) plus ecosystem respiration (ER, a negative flux) expressed in g C m⁻² day⁻¹.
453 Bacterial respiration of DOC was calculated as heterotrophic respiration (HR, a negative flux) from:

$$454 \quad HR = ER + \alpha GPP \text{ with } \alpha = AR / GPP$$

455 with AR, autotrophic respiration and ER, ecosystem respiration, both negative fluxes (oxygen
456 consumption) and GPP, positive flux (producing oxygen). We partitioned ER into auto and
457 heterotrophic respiration with $\alpha=0.5$ (see Demars et al. 2015) and calculated uncertainties using $\alpha=0.2$
458 and $\alpha=0.8$. Bacterial respiration of the added sucrose was calculated as the difference in heterotrophic
459 respiration between the treatment and a control reach during sucrose addition, after standardising for
460 site differences using the control period.

461

462 *Allochthonous organic matter.* The overall flux at the outlet of both streams was calculated as
463 discharge times DOC concentration using the weekly data from the long-term ECN monitoring
464 collected in 2007-2008 under low stable flows. The DOC flux was then related to discharge (Cairn
465 n=74, R²=0.85; Birnie n=100, R²=0.83), to provide an estimate for the days for which DOC was not
466 measured (within the same range of flows).

467 The organic carbon uptake length (Sw_{OC} , in m) and mineralisation velocity (v_{f-OC} , in $m\ day^{-1}$) were
468 calculated as in previous studies (Newbold et al. 1982, Hall et al. 2016), here neglecting POC (see
469 above):

$$470 \quad Sw_{OC} = \frac{Q \times [DOC]}{-HR \times w}$$

471 with [DOC] dissolved organic carbon concentration ($g\ C\ m^{-3}\ day^{-1}$), Q discharge (m^3), HR
472 heterotrophic respiration (a negative flux expressed in $g\ C\ m^{-2}\ day^{-1}$) and w width (m), and

$$473 \quad v_{f-oc} = \frac{-HR}{[DOC]}$$

474

475 *Light.* We derived a conversion factor for photosynthetically active radiation (PAR; $1\ mol\ photon\ m^{-2}$
476 $day^{-1} = 6.13\ g\ C\ m^{-2}\ day^{-1}$) by relating the ratio of total quanta to total energy within the PAR
477 spectrum ($2.5 \times 10^{21}\ photon\ s^{-1}\ kJ^{-1} = 4.15 \times 10^{-3}\ mol\ photon\ kJ^{-1}$; Morel and Smith 1974) with the
478 reciprocal of the energy content of glucose expressed in carbon units ($15.7\ kJ\ g^{-1}\ glucose = 25.4 \times 10^{-3}$
479 $g\ C\ kJ^{-1}$; Southgate and Durnin 1970).

480

481 *Sucrose.* The flux of added sucrose over the treatment reach was determined at the top of the reach
482 (84 m upstream of the flume of Cairn Burn) using the average observed increase in DOC
483 concentrations (28/08, 5/09, 11/09/2007) multiplied by discharge (Fig S1).

484

485 *Whole-stream production.* We calculated the net ecosystem primary production (NPP) by assuming a
486 50% carbon use efficiency ϵ , that is half GPP (reviewed in Demars et al. 2015, 2017) and computed
487 NPP for $\epsilon=0.2$ and $\epsilon=0.8$ to provide uncertainties in our estimates. Heterotrophic production (HP, $g\ C$
488 $m^{-2}\ day^{-1}$) was calculated from heterotrophic growth efficiency HGE and heterotrophic respiration
489 (HR, negative flux expressed in $g\ C\ m^{-2}\ day^{-1}$) as follows:

$$490 \quad HP = \frac{-HR \times HGE}{1 - HGE}$$

491 Heterotrophic production was estimated for low (5%) and moderate (20%) bacterial growth
492 efficiencies (e.g. Berggren et al. 2009, Fasching et al. 2014, Berggren and del Giorgio 2015).

493

494 *Macroinvertebrate consumers.* Secondary production was estimated from the samples collected at the
495 end of the treatment period from the observed standing biomass ($mg\ C\ m^{-2}$) of individual taxa and
496 macroinvertebrate daily growth rate (day^{-1}). The standing biomass was determined from the density
497 (individuals m^{-2}) and average individual biomass (mg C). The growth rate (G) was determined from:

$$498 \quad G = a\ exp(bT)$$

499 with T average water temperature ($10.5^\circ C$), a and b taxon specific constants derived from a global
500 compilation of published data (Golubkov 2000, Gladyshev et al. 2016 – similarly to Morin and
501 Dumont 1994).

502

503 *Ecosystem efficiencies*

504 With all fluxes expressed in $\text{g C m}^{-2} \text{ day}^{-1}$, we calculated the light use efficiency (net primary
505 production/photosynthetic radiations), organic carbon use efficiency (bacterial production/DOC
506 supply), resource use efficiency (consumers/ [net primary production + bacterial production]), and the
507 proportion of biotic CO_2 emissions (net ecosystem production / total CO_2 emissions).

508

509 *Data analyses*

510 We calculated an effect size (i.e. proportional changes) of sucrose addition on our response variables
511 (e.g. nutrient cycling, stoichiometric ratios, metabolic fluxes, ecosystem efficiencies) using the values
512 of the control (C) and treatment (T) reaches, before (B) and after (A) sucrose addition as follows:

$$513 \quad \text{Effect size} = ((C_B - C_A) - (T_B - T_A))/T_B$$

514 with all uncertainties propagated in quadrature using standard deviation δx for sums, and relative
515 uncertainties $\delta x/x$ for the division. The standard error (se) of the effect size was calculated as $se =$
516 $\delta x/\sqrt{n}$ with n average number of independent samples or measurements in C_B, C_A, T_B, T_A (although
517 strictly speaking the samples were pseudo-replicated because the random collection was within one
518 plot). Since the overall design was unreplicated we simply interpreted the effect size in relation to the
519 standard error (we do not report p values).

520 The control period for the metabolic parameters and trophic transfer efficiencies was limited to the
521 two days prior to sucrose addition (21-22 August), so the whole data series was within a single period
522 of low stable flows, i.e. not interrupted by peak flows, producing more comparable results across
523 sites. Daily metabolic parameters were temporally pseudo-replicated (especially during the control
524 period), so it was only possible to report relative uncertainties ($\delta x/x$).

525 All effect sizes and efficiencies were calculated at the reach scale.

526

527 **Results**

528 *DOC addition*

529 We were fortunate to have no rainfall and low stable flows during the whole three weeks of carbon
530 addition. The target concentration was achieved at the top of the studied reach, 28 m downstream of
531 the injection point, where the DOC addition averaged 0.52 mg C L^{-1} (Fig. S1), despite a lower than
532 expected discharge (from 20 L s^{-1} down to 8 L s^{-1} by the end of the addition), due to significant uptake
533 and mineralisation within the 28 m mixing zone (an average 55% loss of sucrose flux from the point
534 of injection to the top of the reach).

535

536

537 *Stream metabolism continuous monitoring*

538 Bryophytes covered 4, 11 and 17% of the river bed in the Birnie control, Cairn control and Cairn
539 treatment reach, respectively. Filamentous green algae (mostly *Microspora* sp., Microsporaceae)
540 percentage cover increased during the period of sucrose addition from 1 to 19%, 11 to 33% and 4 to
541 37% of channel width in the Birnie control, Cairn control and Cairn treatment reach, respectively.

542 Gross primary production (GPP) peaked to $7.6 \text{ g O}_2 \text{ m}^{-2} \text{ day}^{-1}$ ten / eleven days after the start of
543 sucrose addition in the treatment reach before decreasing sharply down to an average $2.4 \text{ g O}_2 \text{ m}^{-2}$
544 day^{-1} during the last four days of the carbon addition, and this despite high photosynthetic active
545 radiations. In contrast, GPP remained relatively constant in the control reaches Birnie Burn (about 1.2

546 g O₂ m⁻² day⁻¹) and Cairn Burn (3.2 g O₂ m⁻² day⁻¹ for the first two weeks declining to 2.1 g O₂ m⁻²
547 day⁻¹ during the last week g O₂ m⁻² day⁻¹) – Fig. 4.

548 Peaks in ecosystem respiration (ER) down to -20 g O₂ m⁻² day⁻¹ (Birnie Burn) and -35 g O₂ m⁻² day⁻¹
549 (Cairn Burn) were more visible for the controls than the treatment reach, with respiration activity
550 inversely related to soil hydrological connectivity, as recorded by soil moisture continuous monitoring
551 (Fig. 4). There was an increase in ER in the treatment reach at the start of the sucrose addition, despite
552 the continuing loss of hydrological connectivity. More specifically, heterotrophic respiration activity
553 associated to sucrose addition peaked sharply 15 days after the start of the addition, processing up to
554 59% of the daily sucrose flux (Fig. 5). On average 35±20% of the added sucrose was respired during
555 the addition over just 84 m (or 15 minutes mean travel time). Heterotrophic production ranged
556 between 2% and 10% of the sucrose flux, based on bacterial growth efficiencies of 0.05 and 0.2
557 respectively.

558

559

560

561 *Nutrient cycling studies and stoichiometry*

562 The background concentrations of nitrate and phosphate were 180 and 90 µg N L⁻¹ and 2 and 4 µg P
563 L⁻¹ in the Birnie control and Cairn treatment reach, respectively. The added geometric mean of N and
564 P were on average 471 µg N L⁻¹ and 24 µg P L⁻¹. The addition of sucrose had no effect on nitrate and
565 phosphate nutrient uptake length and uptake velocity (Fig. 6, Table S3). The phosphate uptake length
566 was highly related to discharge and became very short, down to 31 m in the treatment reach towards
567 the end of the sucrose addition. Phosphate uptake velocity was about 0.2 mm s⁻¹ and an order of
568 magnitude faster than nitrate with uptake lengths in the kilometre range.

569 The molar C:N:P stoichiometric ratios of coarse particulate organic matter and bryophytes remained
570 stable throughout the experiment. Sucrose addition exerted strong effects on filamentous green algae
571 and periphyton stoichiometry (Fig. 7, Table S4). While the molar C:N:P stoichiometric ratios
572 decreased in the control stream, they increased sharply in the treatment reach following sucrose
573 addition: from 330:29:1 to 632:49:1 in filamentous green algae and from 262:29:1 to 428:38:1 in
574 periphyton.

575

576

577 *Fate of added sucrose*

578 The fraction of carbon derived from sucrose (F_S) among the food web resources filamentous green
579 algae, periphyton and its autotrophs and heterotrophs (bacteria) was relatively high at 24, 23, 36 and
580 68%, respectively (Fig. 8, Table S1). The average (range) tissue turnover of consumers was 76% (53-
581 92%) to 88% (67-98%) over 14 to 21 days. The proportion of carbon derived from sucrose in
582 macroinvertebrates, after correcting for tissue turnover, varied among taxa but averaged around 23%
583 independently of the functional groups, except for filter feeders with F_S=64% (Fig. 9, Table S5).

584

585 *Identification of carbon sources and pathways under low flows (end of experiment)*

586 The autotrophs (filamentous green algae and periphyton autotrophs) derived 54±11% of their CO₂
587 from allochthonous sources (groundwater and atmosphere) and 46±11% from bacteria in the control
588 reach. In the treatment reach, autotrophs derived a similar proportion of CO₂ from allochthonous
589 sources (59±9%), some osmotrophic uptake of sucrose (10±7%), and relatively less from bacteria
590 (31±11%). The bacteria used mostly autotrophic carbon (76±12%) relative to allochthonous organic
591 carbon (24±12%) in the control, but preferred sucrose (51±7%) to autotrophic (34±12) and
592 allochthonous organic matter (15±12%) in the treatment reach – see Fig. S2.

593

594 *Quantification of carbon fluxes and efficiencies*

595 The flow food web of the control and treatment reach over the three weeks of sucrose addition were
596 quantified according to our conceptual model (see Fig. 1). Figure 10 illustrates the C fluxes of net
597 primary production, bacterial production and secondary production as well as bacterial CO₂ flux and
598 overall net ecosystem production (emission of CO₂ to the atmosphere). Photosynthetic active radiation
599 (light) decreased slightly from 106 to 80 and 118 to 90 g C m⁻² day⁻¹ in the control and treatment
600 respectively. Allochthonous organic matter (DOC) was more than halved from 177 to 85 and 110 to
601 50 g C m⁻² day⁻¹ in the control and treatment. This was reflected by a general reduction in the organic
602 carbon uptake length 3214 to 2531 m and 4257 to 1886 m in the control and treatment, respectively,
603 independently of the carbon addition. The organic carbon uptake velocity decreased in the control
604 from 0.82 to 0.55 m day⁻¹ but increased in the treatment from 0.53 to 0.76 m day⁻¹.

605

606 Our estimates suggest that a large part of net primary production (0.23, range 0.09-0.38 g C m⁻² day⁻¹)
607 was used for bacterial production (0.15-0.69 g C m⁻² day⁻¹) in the control, and in turn nearly half of
608 the CO₂ fixed by autotrophs was derived from bacterial CO₂, although it represented a small fraction
609 of bacterial respiration driving net ecosystem production (biotic CO₂ emissions). These reciprocal C
610 subsidies between autotrophs and bacteria were not as strong relative to net primary production (0.91,
611 range 0.36-1.46 g C m⁻² day⁻¹) and bacterial production (0.14-0.66 g C m⁻² day⁻¹) in the treatment
612 during sucrose addition. The estimated flux of allochthonous organic matter assimilated by bacteria
613 was similar in the control (0.04-0.17 g C m⁻² day⁻¹) and treatment (0.02-0.10 g C m⁻² day⁻¹). Part of the
614 allochthonous organic matter respired by bacteria was recycled by primary producers and accounted
615 for 11±6% of net primary production in the control.

616

617 We also derived, from the BACI design, the effect size of sucrose addition on selected whole-
618 ecosystem metabolic properties and efficiencies (see Fig. 11, Table S6). All estimated ecosystem
619 properties and efficiencies were summarised in Table S7 for Birnie control and Cairn treatment before
620 and after sucrose addition. We present some key highlights below.

621

622 While GPP increased marginally (12%), there was a small relative increase in light use efficiency
623 (37%). ER intensified by 70%, and the net ecosystem production (NEP) became relatively more
624 negative by 125%, i.e. 125% relative increase in biotic CO₂ emissions. Heterotrophic respiration and
625 production increased by 89%, and this was reflected by a shorter (-40%) uptake length (Sw_{OC}) and
626 faster mineralisation velocity (92%) of organic carbon. The proportion of DOC flux respired (range
627 2.2-5.3%) and organic carbon use efficiency by bacteria (range 0.1-0.3%) increased by 112%. While
628 there was a relative decrease (-20%) in total CO₂ emissions, the proportion of biotic CO₂ emission
629 increased by 88%. The reciprocal subsidies between autotrophs and bacteria were weaker by 33%
630 (autotrophs to bacteria) and 55% (bacteria to autotrophs) in the treatment relative to the control.

631

632 The average consumer biomass per individual was similar between the control (0.20 mg C ind⁻¹) and
633 the treatment reach (0.21 mg C ind⁻¹) and average production per individual was slightly higher in the
634 treatment reach (6 µg C ind⁻¹ day⁻¹) than in the control (5 µg C ind⁻¹ day⁻¹), at the end of the sucrose
635 addition. Consumer density (range 1300-6000 ind m⁻²) increased by 72% due to sucrose addition and
636 consumer production was higher in the treatment (36 mg C m⁻² day⁻¹) than in the control (20 mg C m⁻²
637 day⁻¹) at the end of the experiment. The resource use efficiency (trophic transfer efficiency) by
638 consumers was similar between the control and the treatment reach (2-5%), with a size effect of
639 sucrose addition ranging from -33% to +8% depending on the heterotrophic growth efficiency (0.05
640 and 0.2, respectively) used to calculate heterotrophic production.

641

642

643 Discussion

644

645 Our experiment showed that a small continuous addition of labile DOC (0.52 mg C L⁻¹ as sucrose,
646 12% of total DOC, Fig. S1) can profoundly alter whole-ecosystem behaviour. The use of a before and
647 after control experiment together with the addition of a deliberate tracer with a distinctive $\delta^{13}\text{C}$
648 signature allowed not only to trace the fate of the added carbon into the treated reach but also to build
649 the flow food web of the control reach, unravel C reciprocal subsidies between autotrophs and
650 bacteria, and demonstrate the potential for some natural allochthonous organic matter to feed the
651 primary producers via bacterial respiration.

652

653

654 *Reciprocal subsidies between autotrophs and bacteria*

655 The use of autotroph carbon by bacteria has been shown before using a photosystem II inhibitor in
656 biofilm (Neely and Wetzel 1995) or $\delta^{13}\text{C}_{\text{DIC}}$ additions (e.g. Lyon and Ziegler 2009, Risse-Buhl et al.
657 2012, Kuehn et al. 2014, Hotchkiss and Hall 2015). The degradation of natural DOC by bacteria has
658 been inferred from bacterial respiration (e.g. Cory et al. 2014, Mineau et al. 2016, Demars 2018), and
659 many bacterial production estimates have been published (e.g. Fischer and Pusch 2001, Fukuda et al.
660 2006). The interaction strength between autotrophs and bacteria has been inferred statistically along
661 nutrient gradients (e.g. Carr et al. 2005, Scott et al. 2008), but this is the first study, to our knowledge,
662 quantifying the reciprocal carbon subsidies between autotrophs and bacteria in streams, notably
663 bacterial CO₂ use by autotrophs. This recycling of CO₂ in turbulent headwater streams is possible
664 within the intricate matrix of the biofilm (e.g. Kamjunke et al. 2015, Battin et al. 2016) or an algal
665 mat, as when the CO₂ is released in the water column it is very quickly degassed to the atmosphere
666 (here 5-15 min, Demars 2018). It is important to note that the reciprocal carbon subsidies were
667 identified under low flow conditions when stream water was hydrologically disconnected from soil
668 water. When the soils are hydrologically connected to the stream water, and supply high fluxes of
669 DOC, bacterial respiration and production was shown to be virtually independent of primary
670 production (Demars 2018).

671

672 Our flow food web under low flows (end of experiment) suggested a strong microbial loop (to
673 autotrophs) in the control stream, as expected at the low end of the nutrient gradient (Scott et al. 2008,
674 Lyon and Ziegler 2009). The addition of labile DOC weakened this microbial loop, as indicated by
675 the difference in reciprocal subsidies between algae and bacteria in the control and treatment (Fig. 10,
676 11), as predicted under nutrient limitation (Fig. 1). Bacterial production in the treatment reach relied
677 more on sucrose than autotrophs and allochthonous organic matter, assuming a similar bacterial
678 growth efficiency for all sources (range 5-20%), which is plausible under nutrient limitation (del
679 Giorgio and Cole 1998) and corresponded to previous studies (e.g. Berggren et al. 2009, Fasching et
680 al. 2014, Berggren and del Giorgio 2015).

681 The bacterial CO₂ flux to autotrophs assumed its $\delta^{13}\text{C}$ signature was the same as that of the bacteria,
682 but the proportion of C sources used in bacterial production may be different to the proportion of C
683 sources respired by bacteria. The controlled experiment allowed to calculate the bacterial respiration
684 of the added sucrose as 2.16 g C m⁻² day⁻¹ over the three weeks of sucrose addition. This represented
685 82% (53-184%) of the estimated treatment bacterial respiration (2.64 ± 1.45 g C m⁻² day⁻¹, Table S7)
686 and the proportion of sucrose in bacteria was estimated at 51 ± 7 % (Fig. S2). So the estimates were
687 still within measurement errors.

688

689 The autotrophs did not include bryophytes in our flow food web calculations because their
690 contribution to primary production was thought to be negligible over the few weeks of the

691 experiment, and particularly towards the end of the experiment (on which data the flow food webs
692 were based) when filamentous green algae were covering bryophytes. The lack of bryophyte growth
693 under very low phosphorus concentrations (here 2-4 $\mu\text{g P L}^{-1}$ of soluble reactive P) combined with
694 shading by epiphytes has been well documented (e.g. Finlay and Bowden 1994). This may also
695 explain the lack of changes in bryophyte C:N:P stoichiometry (Fig. 7).

696

697

698 *Boom and bust: role of nutrients*

699 Gross primary productivity (GPP) appeared to be stimulated by the addition of sucrose but this was
700 short lived, despite sustained light availability during the addition period (Fig. 4). Heterotrophic
701 respiration of sucrose peaked after two weeks (three days after GPP) but crashed within days while
702 the supply of sucrose was continuously flowing through the reach (with sucrose concentration
703 increasing from 0.22 to 0.88 mg C L^{-1} with falling discharge). This was in sharp contrast to the peaks
704 in respiration followed by a more sustained response of ecosystem (mostly heterotrophic) respiration
705 to hydrological connectivity with soil water in the control reaches (Fig. 4, see Demars 2018). This
706 boom and bust in the treated reach was likely due to nutrient limitation, mostly P according to the
707 changes in filamentous green algae and periphyton C:N:P stoichiometry. Surprisingly, the shortfall of
708 nutrients (N, P) was not compensated by faster nutrient cycling rates, as observed in previous studies
709 with much higher labile DOC additions (e.g. Bernhardt and Likens 2002). This may be due to
710 limitation of phosphate uptake by NH_4 availability (long term median in both streams 7 $\mu\text{g N L}^{-1}$) in
711 streams relatively rich in nitrate (see Oviedo-Vargas et al. 2013). The higher primary productivity in
712 the Cairn burn may partly result from higher P availability (4 $\mu\text{g P L}^{-1}$) than in the control Birnie Burn
713 (2 $\mu\text{g P L}^{-1}$), despite lower nitrate availability (90 $\mu\text{g N L}^{-1}$ versus 180 $\mu\text{g N L}^{-1}$, respectively) – Table
714 S3. These different nutrient supply rates may reflect the legacy of past experimental amendments (Ca,
715 N, P, K) in 33 ha of the Cairn burn catchment aimed to increase grassland productivity in the late
716 1970s and early 1980s (Hill Farming Research Organisation 1983).

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719 *Metabolic balance*

720 The metabolic balance (or net ecosystem production) was responsible for a quarter of CO_2 emissions
721 in the control and increased from 6 to 12 % in the treatment reach during sucrose addition. CO_2
722 emissions from these streams were therefore largely dominated by soil CO_2 derived from the
723 mineralisation of soil organic matter, rather than rock weathering of Dalradian acid schist drifts
724 (Demars 2018). The proportion of in-stream biotic emissions were comparable to continental scale
725 studies (<14% in African rivers, Borges et al. 2015; 28% in north American rivers, Hotchkiss et al.
726 2015). It is important to note that ecosystem photosynthesis and community respiration can be
727 decoupled at individual sites in long term studies with temporal changes in C supply (e.g. Roberts et
728 al. 2007, Beaulieu et al. 2013, Demars 2018). Under summer low flows respiration activity is more
729 constrained by gross primary production explaining the strong interaction between photosynthesis and
730 respiration (e.g. Demars et al. 2016).

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733 *Fate of carbon*

734 Sucrose is very labile and is well known to promote the growth of filamentous bacteria *Sphaerotilus*
735 *natans* (“sewage fungus”), even at relatively low concentrations (0.25-1.00 mg L^{-1}) in a stable flow
736 forested stream (Warren et al. 1964). This was not observed in this study, likely because the moorland
737 streams studied here were more open and colonised by *Microspora*, a common genus of filamentous
738 green algae in Scottish streams (Kinross et al. 1993). *Microspora* was able to uptake sucrose by
739 osmotrophy (Wright and Hobbie 1966) but this accounted to only 10% of C uptake by *Microspora*.
740 Overall daily uptake of sucrose by autotrophs represented 1.2% of the average daily flux of sucrose.

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The proportion of added labile carbon (sucrose) in the consumers varied widely between species (Table S5), as observed in previous studies (e.g. Hall 1995, Hall and Meyer 1998, Collins et al. 2016), but not across functional feeding groups (19-27%), except for filter feeders (64%). This is questioning the usefulness of functional feeding groups, as defined here, to construct food webs. Filter feeders and notably blackflies (Simuliidae) have been shown to directly assimilate DOC, extracellular polysaccharides and colloidal particles (Couch et al. 1996, Hershey et al. 1996, Ciborowski et al. 1997, Wotton 2009). This can explain the high proportion of added sucrose in blackflies (81%). Since the densities of blackflies were low (average 88-231 individuals m⁻²), C flux from direct uptake by consumers (filter feeders) were extremely small. The mass of sugar retained by all consumers was only 292 ±107 mg C m⁻², or about 25 g C for the treated stream reach, representing 0.2% of the sucrose flux over the three-week addition.

The treatment reach did not show the large peaks in respiration (outside the period of sucrose addition) that the control reaches showed when the catchment was hydrologically connected, as indicated by soil moisture (Fig. 4, Demars 2018). This is likely because the treatment reach is a more constrained reach largely disconnected from the land. It was initially chosen to avoid lateral inflows which were very small (2.3%) and mostly from a spring fed flush (i.e. groundwater), rather than seepage from organic and riparian soils known to stimulate bacterial activity (e.g. Brunke and Gonsler 1997, Pusch et al. 1998). Interestingly the average bacterial respiration was similar in the control and treatment reaches over the three weeks of sucrose addition (Fig. 10), suggesting part of the organic matter respired in the control was relatively labile and comparable to the 0.5 mg C L⁻¹ of added sucrose. The dynamic of this respiration within the three weeks was very different however, decreasing with the progressive loss of hydrological connectivity with soils in the control, and peaking after two weeks in the treatment (Fig. 4).

By and far, the largest quantity of carbon processed by bacteria was lost as CO₂ emission. Heterotrophic respiration over the treated reach respired (on average) 35% of the added sucrose. Bacterial production averaged 2-10% of the sucrose flux. Heterotrophic respiration of natural allochthonous DOC was about 3% in the control stream and 2.2 % prior to sucrose addition in the treatment. The fact that sucrose was processed 10 times faster than natural DOC was well reflected by the shortening of the organic uptake length (S_{wOC}) and increased mineralisation velocity (v_fOC). This was not surprising (see e.g. Marcarelli et al. 2011, Mineau et al. 2016), but the rate of mineralisation of natural DOC was relatively high and, scaled up to the first order catchment, represented 23±11% of the annual DOC flux in the control reach of the Cairn burn (Demars 2018). The changes in organic uptake lengths in the control and the differences between control and treatment prior to sucrose addition can be explained by the loss of hydrological connectivity with soil water, as indicated by the changes in soil moisture, and the difference in lateral inflows between the control (10.7%) and the treatment (2.3%). The shortening of the organic uptake length (S_{wOC}) reflected more hydrological changes as indicated by the different direction of change in organic carbon uptake (v_fOC) between the control and the treatment. In the control v_fOC declined by 32% against a 52% decline in DOC supply. In contrast, v_fOC increased by 43% in the treatment with the addition of labile carbon, despite a 55% fall in DOC supply (similar to the control). Overall the uptake lengths of natural DOC (i.e. excluding those influenced by sucrose addition) were longer (2.5-4.3 km) than the length of the first order streams studied here (1 km), and so a large part of the carbon is released to downstream ecosystems as previously observed (e.g. Wiegner et al. 2005), especially during time of loss of hydrological connectivity with the soil of the catchment (Demars 2018). The simple comparison between flow food webs (Fig. 10) suggest that the addition of labile carbon did not prime the bacterial use of natural allochthonous DOC.

792 *Choice of carbon for DOC addition*

793 We initially considered to add natural DOC to the stream after isolating DOC using reverse osmosis
794 (Sun et al. 1995; RealSoft Pro2S, US) but the product was too salty with high pH and high nutrient
795 concentrations (Stutter and Cains 2016). While reverse osmosis may be combined with electro dialysis
796 to avoid co-concentration of salt (Koprivnjak et al. 2006), the quantities needed for our whole-
797 ecosystem experiment were simply too large. We also considered several commercial humate sources
798 but rejected their use because of pH, solubility and nutrient issues. Sucrose derived from sugarcane
799 (C4 carbon fixation) has a very distinctive carbon stable isotope signature compared to the autotrophs
800 in temperate ecosystems (C3 carbon fixation). It offered the possibility to trace its fate through the
801 food web, and even to identify the bacterial carbon pathways in the control streams. It turned out that
802 sucrose was a more judicious choice than first thought because labile DOC (polysaccharide, amino
803 acids) is likely driving the respiration of the studied streams at the land-water interface (Demars
804 2018), as found in bioreactors (e.g. Drake et al. 2015).

805

806 **Conclusions**

807 Part of the carbon derived from allochthonous organic matter can feed the autotrophs via the CO₂
808 produced by stream bacterial respiration, intermingling the green and brown webs. The interaction
809 between autotrophs and bacteria shifted from mutualism to competition with carbon addition under
810 nutrient limitation (N, P) increasing biotic CO₂ emissions. Without nutrient limitation, mutualism
811 could be reinforced by a positive feedback loop, maintaining the same biotic CO₂ emissions. Even a
812 small increase in labile dissolved organic carbon supply due to climate and land use change could
813 have large effects on stream food web and biogeochemistry with implications for the global C cycle
814 under stoichiometric constraints.

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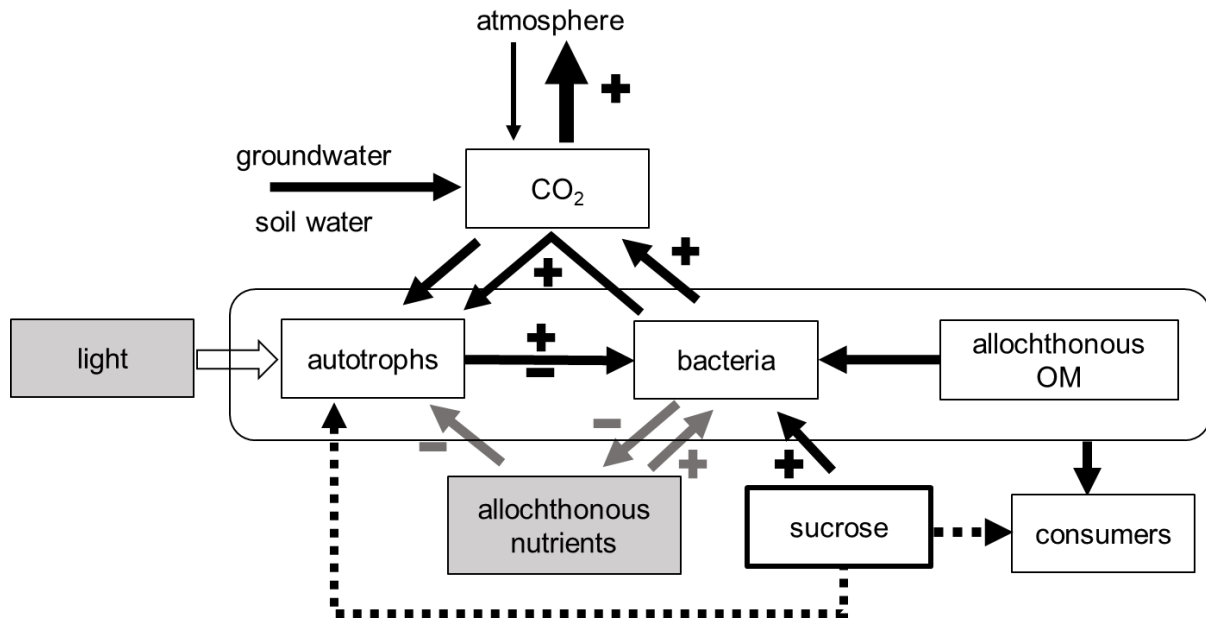
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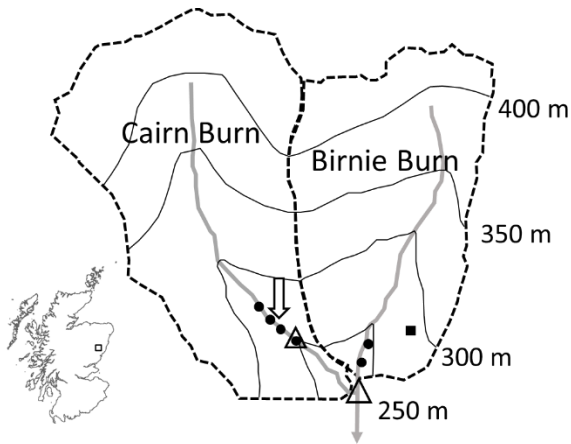
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1186 **Fig 1.** Possible mutual benefits of algae and bacteria and expected changes (+ and – symbols) in
1187 carbon and nutrient (N, P) fluxes due to sucrose addition. If nutrients are limiting, then C reciprocal
1188 subsidies between bacteria and autotrophs may be weakened by sucrose addition, i.e. there could be a
1189 shift between mutualism and competition. With enough nutrients the C microbial loop (to autotrophs)
1190 may be strengthened via an increase in CO₂ from bacterial respiration, i.e. positive feedback loop.

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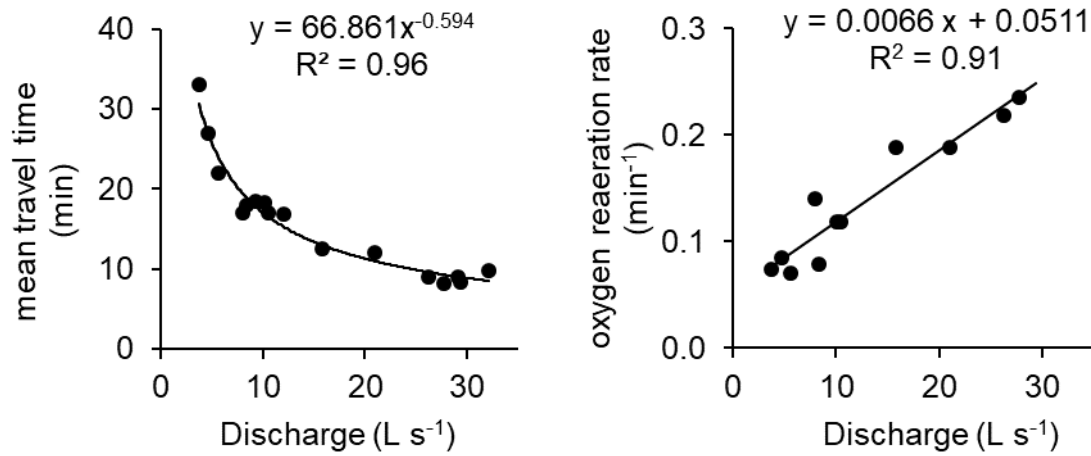
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1196 **Fig 2.** Paired stream experiment at Glensaugh research station. Birnie Burn is the control stream and
1197 Cairn Burn the manipulated stream with DOC addition indicated by the arrow. The Cairn Burn also
1198 has a control upstream of the treatment reach. The symbols refer to flumes (open triangles), dissolved
1199 oxygen stations (filled circles) and soil moisture instrumentation (filled square). The 50 m elevation
1200 contour lines are indicated. The catchment area is 0.99 km² (0.90 km² at the flume) for Cairn Burn
1201 and 0.76 km² for Birnie Burn. Inset shows the location of Glensaugh in Scotland.

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1207 **Fig 3.** Discharge was an excellent predictor of mean travel time and oxygen reaeration measured from
1208 NaCl and propane tracer studies (data from Cairn treatment reach, data of the other two reaches were
1209 presented in Demars 2018).

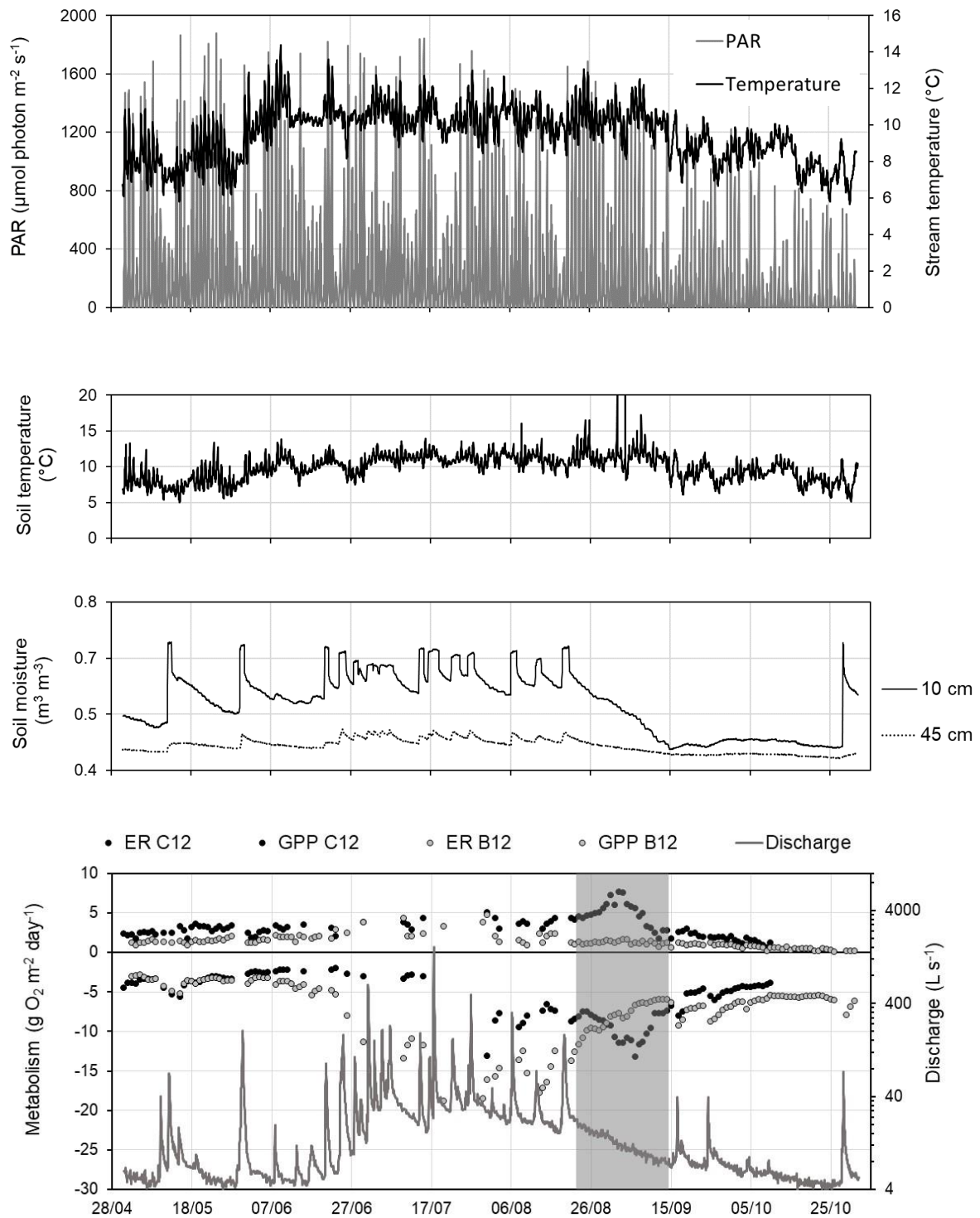
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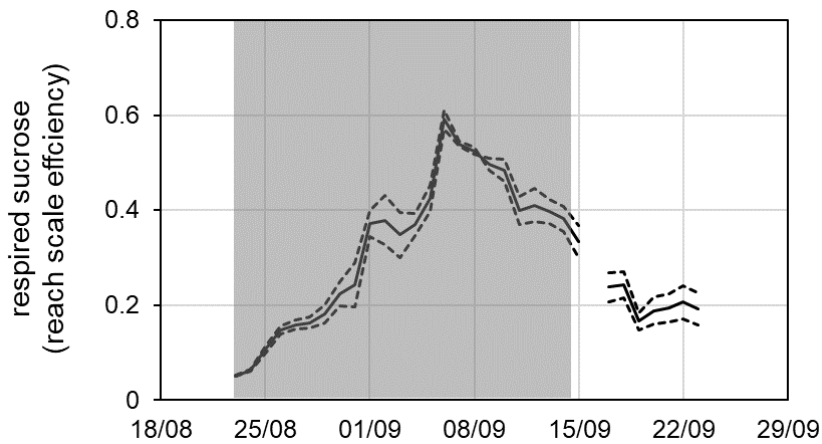
1219 **Fig. 4.** Continuous data monitoring and stream metabolism in the control reach (Birnie Burn, B12)
1220 and treatment reach (Cairn Burn, C12). Ecosystem respiration (ER) is negative (consuming O₂) and
1221 gross primary production (GPP) is positive (producing O₂). The period of DOC (sucrose) addition
1222 (23/08-14/09) is indicated by grey shading. The depths for soil moisture correspond to the depths of
1223 the organic soil (10 cm) and subsoil (45 cm).

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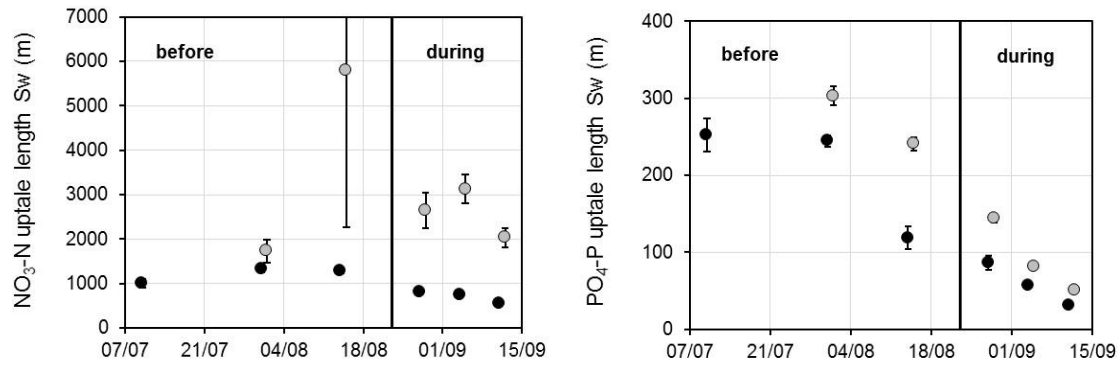
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1230 **Fig 5.** Efficiency of respired sucrose in the 84 m long treatment reach (with mean travel time of 15
1231 minutes) during (shaded area, 23/08-14/09) and shortly after sucrose addition, calculations based on
1232 heterotrophic respiration in the treatment reach relative to the Birnie control (same results with Cairn
1233 control, not shown). The black line was calculated with an autotrophic respiration (AR) of $0.5 \times \text{GPP}$
1234 and dashed lines with $\text{AR} = 0.2 \times \text{GPP}$ and $0.8 \times \text{GPP}$ (see method). On average $35 \pm 20\%$ of the daily flux
1235 of sucrose was respired within that reach during the sucrose addition.

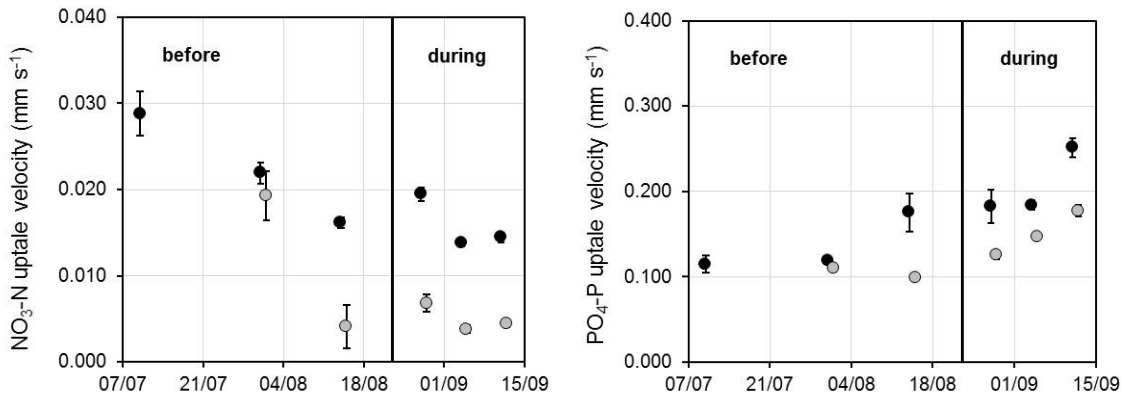
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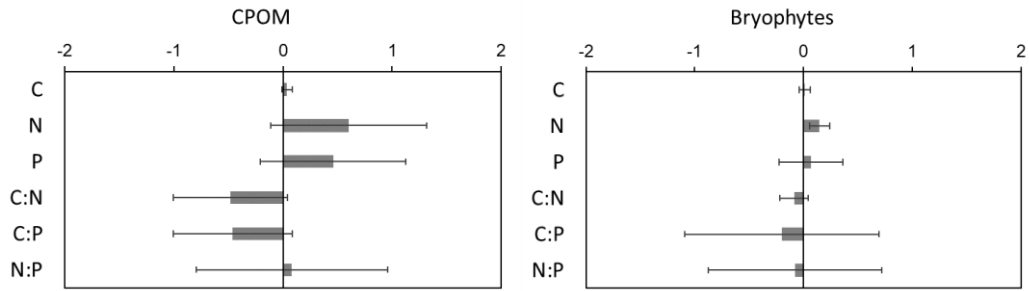
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1240 **Fig. 6.** Stream nutrient uptake length (S_w) and uptake velocity (v_i) before and during sucrose addition
1241 in the control (grey symbols) and treatment (black symbols). Error bars represent sem, with some
1242 error bars smaller than the symbols. Note the different magnitudes on the y axes between $\text{NO}_3\text{-N}$ and
1243 $\text{PO}_4\text{-P}$.

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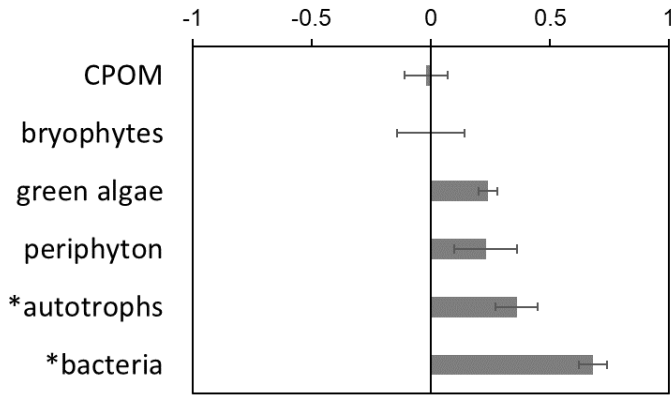
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1247 **Fig. 7.** Proportional changes in C, N, P (% w/w) and molar C:N:P stoichiometric ratios in the basal
1248 food web resources due to sucrose addition based on the before and after control impact experimental
1249 design. Error bars represent sem.

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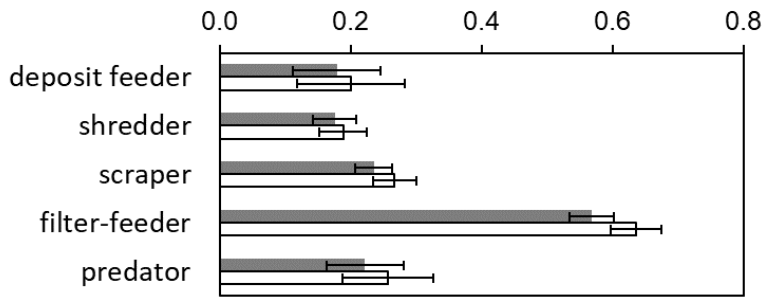
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1252 **Fig. 8.** Proportion of carbon derived from added sucrose in food web basal resources based on $\delta^{13}\text{C}$
1253 (see TABLE S1) and PLFA $\delta^{13}\text{C}$ for autotrophs and bacteria in periphyton (indicated by *, See TABLE
1254 S2). Error bars represent se. Size effect calculated from BACI design, except for periphyton
1255 autotrophs and bacteria (periphyton collected at the end of the experiment in the control and treatment
1256 reach, see method).

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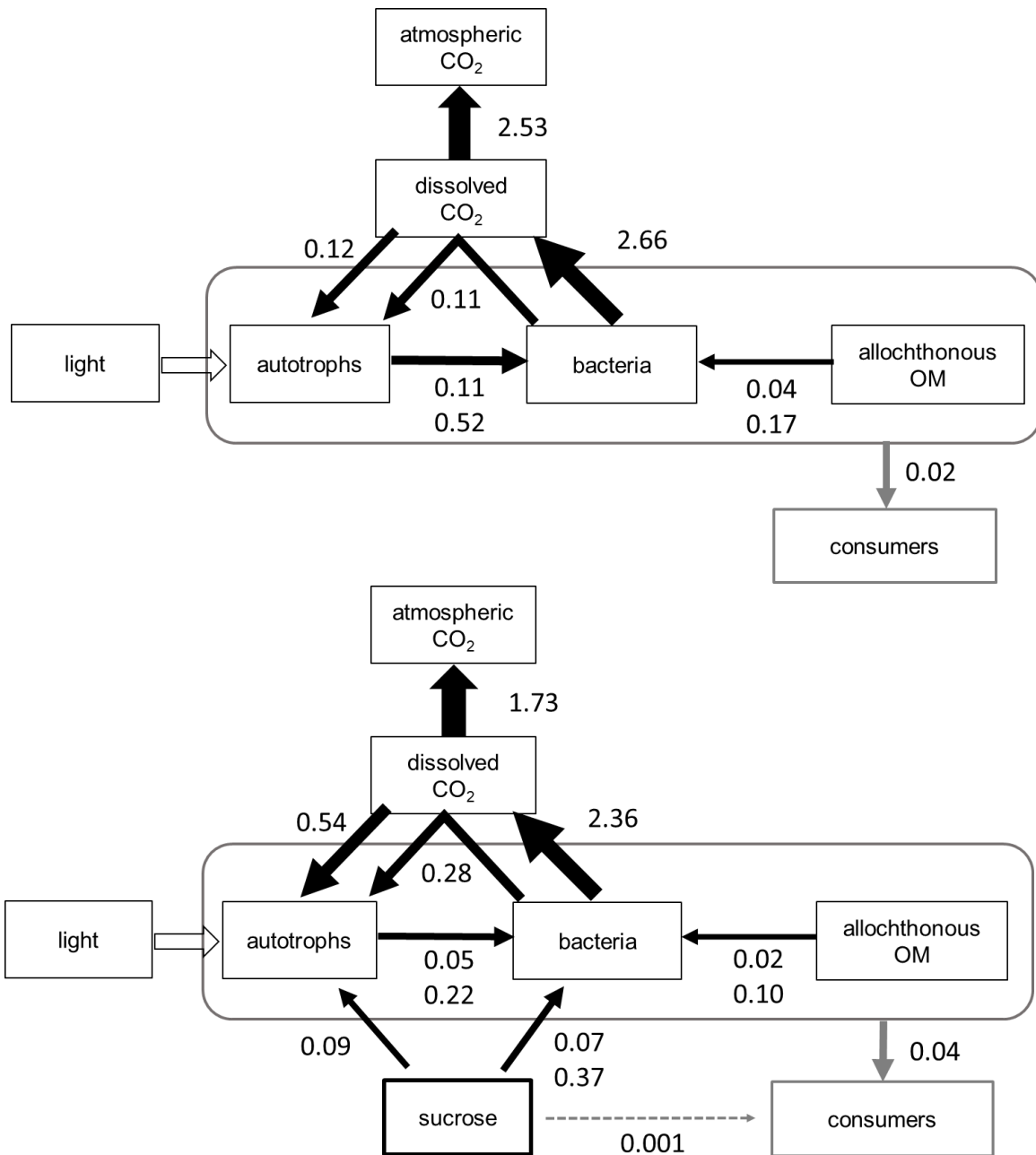
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1261 **Fig. 9.** Proportion of carbon derived from added sucrose in macroinvertebrate consumers by
1262 functional feeding groups based on the before and after control impact experimental design and
1263 associated $\delta^{13}\text{C}$ changes (see TABLE S5); observed (grey bars) and at equilibrium (open bars, see
1264 method); error bars represent the standard error.

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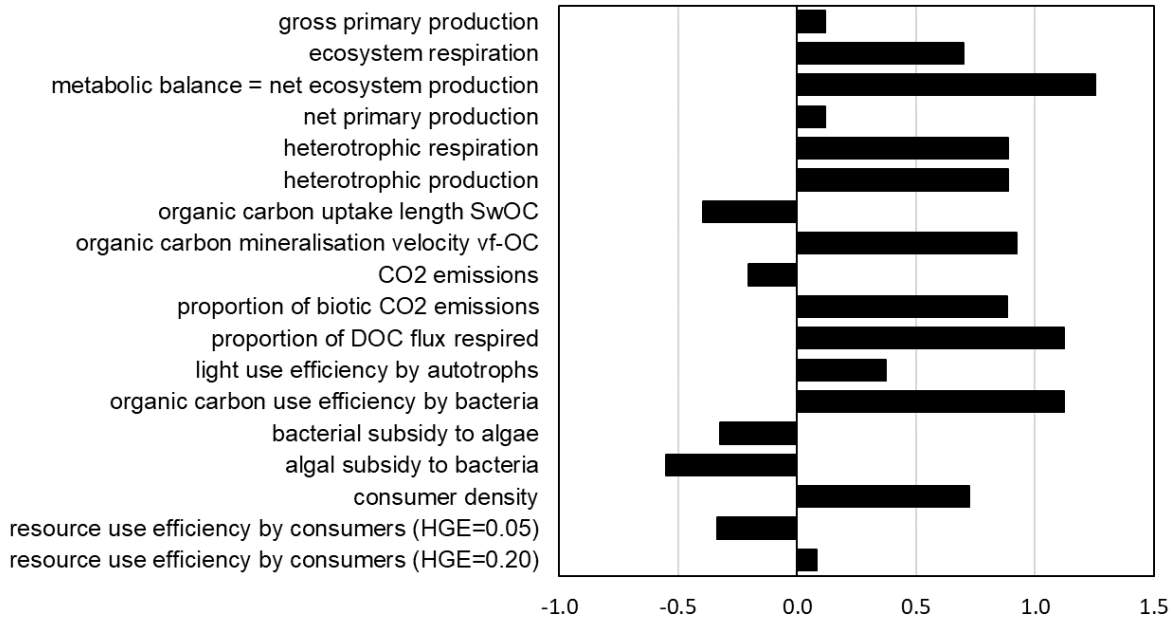
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1269 **Fig. 10. Flow food webs under low flows when stream water hydrologically disconnected from**
 1270 **soils:** in-stream biotic carbon fluxes ($\text{g C m}^{-2} \text{ day}^{-1}$, black and grey arrows) and energy flow (light) in
 1271 the control (top) and treatment reach (bottom) after three weeks of sucrose addition, based on source
 1272 partitioning using stable isotopes and production estimates. Autotrophs represent net primary
 1273 production, bacteria represent heterotrophic production, and consumers represent secondary
 1274 production. Bacterial respiration and net ecosystem production (biotic CO_2 emissions) are also
 1275 represented. Two figures were given for bacterial production based on heterotrophic growth
 1276 efficiencies (HGE) of 0.05 (low) and 0.2 (moderate). Note: CO_2 emissions to the atmosphere
 1277 represented 26% and 12% of total CO_2 emissions, in the control and treatment respectively with most
 1278 CO_2 derived from soil water and groundwater.

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1282 **Fig. 11.** Effect size (1=100%) of sucrose addition on selected ecosystem properties at reach scale
1283 based on the before and after control impact experimental design, except for resource use efficiency
1284 by consumers and algae-bacteria reciprocal subsidies relying on a simple comparison between the
1285 control and the treatment reach. See Table S6 for uncertainties (most very large) and Table S7 for the
1286 individual values. The resource use efficiency by consumers was estimated for two heterotrophic
1287 growth efficiency (HGE).

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