Understanding and predicting extreme wildfires in the contiguous United States

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Abstract

Wildfires are becoming more frequent in parts of the globe, but predicting where and when extreme events occur remains difficult. To explain and predict wildfire extremes across the contiguous United States, we integrate a 30 year wildfire occurrence record with meteorological and housing data in spatiotemporal Bayesian models with spatially varying nonlinear effects. We compared models with different distributions for the number and sizes of large fires. A zero-inflated negative binomial model for fire counts and a lognormal model for burn areas provided the best performance. We find that dryness and air temperature strongly regulate wildfire risk, with precipitation, and housing density playing weaker roles. Statistically, most of the variability in the chance of an extreme wildfire results from changes in fire frequency that influence sampling from the tails of fire size distributions, rather than changes in the underlying distributions of expected fire sizes. Meteorologically, extreme events occur when conditions are hot and dry, but the effect of air temperature is spatially variable. This model attains 98.9%interval coverage for the number of fires in a withheld data set over a five-year prediction time horizon. We argue that recent wildfire extremes need not be surprising, and future extremes might be predictable if accurate meteorological forecasts are available to drive models of fire occurrence and size.

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1 Introduction

Wildfire frequency and burned area has increased over the past couple decades in the United 2 States (Dennison et al. 2014; Westerling 2016), and elsewhere (Krawchuk et al. 2009; Pechony 3 and Shindell 2010). In addition to the ecological and smoke impacts associated with increased 4 burned area, there has been an increasing interest in extreme wildfires (Williams 2013) given 5 their impact on human lives and infrastructure (Kochi et al. 2010; Diaz 2012). While case 6 studies of particular extremes provide insight into what caused past events (Peterson et al. 7 2015), predictions of future extremes at a national level could inform disaster related resource 8 allocation. Here, we consider an extreme wildfire to be a fire with the largest burned area 9 over a bounded spatiotemporal domain, e.g., within a spatial region and a temporal interval. 10 Factors driving wildfire extremes vary in space and time (Barbero et al. 2014), but it is

11 unclear how best to account for this in a predictive model. Previous efforts have used year or 12 region-specific models, aggregating over space or time (Bermudez et al. 2009), temporally or 13 spatially explicit models (Mendes et al. 2010), and spatial models with year as a covariate 14 (Díaz-Avalos, Juan, and Serra-Saurina 2016). Recently, rich spatiotemporal models have been 15 described with linear, spatially constant covariate effects (Serra, Saez, Juan, et al. 2014; 16 Serra, Saez, Mateu, et al. 2014). However, linear, spatially constant effects are suboptimal 17 over large spatial domains with nonlinear drivers (Fosberg 1978, Goodrick (2002), Preisler et 18 al. (2004); Preisler and Westerling 2007; Balshi et al. 2009; Krawchuk et al. 2009; Pechony 19 and Shindell 2009; Vilar et al. 2010; Woolford et al. 2011; Woolford et al. 2014). For 20 example, global wildfire probability shows a hump-shaped relationship with temperature and 21 moisture (Moritz et al. 2012). Interactions among drivers also impose nonlinearity, e.g., in 22 hot and dry climates fuel sparsity limits fire spread (McLaughlin and Bowers 1982), but in 23 cold and wet climates, fires are energy limited (Krawchuk and Moritz 2011). 24

Prediction is also complicated by uncertainty in which distribution(s) to use to assign probabilities to extreme events. The generalized Pareto distribution (GPD) has frequently been used (Bermudez et al. 2009; Jiang and Zhuang 2011), but the GPD requires a threshold to delineate extreme events (Davison and Smith 1990, Coles (2014)). The utility and validity

of a threshold for extremes in a heterogeneous region is debatable (Tedim et al. 2018). 29 Recently proposed metastatistical extreme value (MEV) approaches do not require such a 30 threshold (Marani and Ignaccolo 2015; Zorzetto, Botter, and Marani 2016). In the MEV 31 framework, the occurrence and size of future events, and the parameters of their distributions 32 are treated as random variables which together imply a distribution for extremes. This 33 approach has roots in compound distributions (Dubey 1970; Wiitala 1999), doubly stochastic 34 processes (Cox and Isham 1980), superstatistics (Beck and Cohen 2003), and the Bayesian 35 posterior predictive distribution (Gelman et al. 2013). The link to Bayesian inference 36 is particularly useful, as it provides an easy way to propagate uncertainty forward to to 37 predictions of extremes (Coles, Pericchi, and Sisson 2003). 38

Here, we extend the MEV perspective to account for non-linear, spatially varying wildfire dynamics with the goal of predicting and explaining extreme wildfire events across the contiguous United States. We aim to predict occurrence (where and when), and magnitude (burn area) of large wildfires at a monthly time scale. Such predictions can be used to prioritize reactive fire suppression resources or inform proactive wildfire risk mitigation.

$_{44}$ Methods

45 Data description

We acquired wildfire event data for the contiguous United States from the Monitoring Trends in Burn Severity (MTBS, www.mtbs.gov) program (Eidenshink et al. 2007), which includes spatiotemporal information on the occurrence of wildfires in the United States from 1984 to 2015. The MTBS data contain fires greater than 1000 acres in the western U.S. and greater than 500 acres in the eastern U.S. For consistency across the U.S., we discarded all records in the MTBS data less than 1000 acres, retaining 10,315 fire events (Figure 1A). Each event in the MTBS data has a discovery date, spatial point location, and final size.

To explain fire size and occurrence, we used a combination of meteorological variables including humidity, air temperature, precipitation, and wind speed. These variables were bioRxiv preprint doi: https://doi.org/10.1101/384115; this version posted August 3, 2018. The copyright holder for this preprint (which was not certified by peer review) is the author/funder, who has granted bioRxiv a license to display the preprint in perpetuity. It is made available under a CC-BY 4.0 International license.

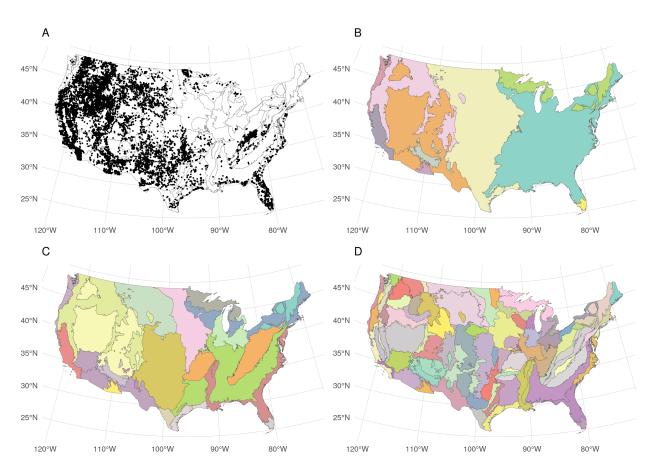


Figure 1. A. Fire ignition locations are shown as points across the study region. Colors in panels B, C, and D show level 1, 2, and 3 ecoregions respectively.

selected on the basis of previous work, and also with an aim to capture directly the effects 55 of easily interpretable meteorological quantities. Meteorological layers were acquired from 56 the gridMET data (Abatzoglou 2013) that blends monthly high-spatial resolution (~4-km) 57 climate data from the Parameter-elevation Relationships on Independent Slopes Model (Daly 58 et al. 2008) with high-temporal resolution (hourly) data from the National Land Data 59 Assimilation System (NLDAS2) using climatologically aided interpolation. The resultant 60 products are a suite of surface meteorological variables summarized at the daily time step 61 and at a 4-km pixel resolution. Daily total precipitation, minimum relative humidity, mean 62 wind speed, and maximum air temperature were averaged at a monthly time step for each of 63 84 Environmental Protection Agency level 3 (L3) ecoregions for each month from 1984 to 64 2015 (Omernik 1987; Omernik and Griffith 2014). We also computed cumulative monthly 65 precipitation over the previous 12 months for each ecoregion-month combination. We chose 66

to segment the U.S. with level 3 ecoregions as a compromise between the more numerous (computationally demanding) level 4 ecoregions, and the coarser level 2 ecoregions.

We used publicly available housing density estimates that were generated based on the U.S. 2000 decennial census as explanatory variables that may relate to human ignition pressure (Radeloff et al. 2010). These are provided at decadal time steps, and spatially at the level of census partial block groups. To generate approximate measures of housing density at monthly time intervals, we used a simple linear interpolation over time for each block group, then aggregated spatially across block groups to compute mean housing density for each ecoregion in each month.

76 Model development

⁷⁷ We built two types of models: one describing the occurrence of fires within each L3 ecoregion ⁷⁸ over time (i.e., the total number of fires occurring in each ecoregion for each month from ⁷⁹ 1984 - 2015), and another describing the size of each wildfire in each ecoregion and month. ⁸⁰ For occurrence models, the response variable was a count (number of fires), and for burn ⁸¹ area models, the response was a continuous positive quantity (size of each fire event). We ⁸² used the period from 1984 to 2009 for training, witholding the period from 2010 to 2015 to ⁸³ evaluate predictive performance.

⁸⁴ Fire occurrence

We constructed four models for fire occurrence and compared their predictive performance 85 based on test-set log likelihood and posterior predictive checks for the proportion of zeros, 86 maximum count, and total count. The models differed in the distributions used in the 87 likelihood, representing counts as a Poisson, negative binomial, zero-inflated Poisson, or 88 zero-inflated negative binomial random variable. The Poisson distribution is a common choice 89 for counts, and the negative binomial distribution provides an alternative that can account 90 for overdispersion. The zero-inflated versions of these distributions include a component to 91 represent extra zeros, which might be expected to work well if there are independent processes 92

⁹³ that determine whether nonzero counts are possible (Lambert 1992).

For spatial units (ecoregions) s = 1, ..., S and time steps (months) t = 1, ..., T, each model 94 defines a probability mass function for $n_{s,t}$: the number of fires over 1000 acres in ecoregion 95 s and time step t. For each of the four count distributions under consideration, location 96 parameters $\mu_{s,t}$ and (for zero-inflated models) structural zero inflation parameters $\pi_{s,t}$ were 97 allowed to vary in space and time. We used a log link function to ensure that $\mu_{s,t} > 0$, and 98 a logit link function to ensure that $\pi_{s,t} \in (0,1)$. Concatenating over spatial and temporal 99 units, so that $\mu = (\mu_{s=1,t=1}, \mu_{s=2,t=1}, ..., \mu_{s=S,t=1}, \mu_{s=S,t=2}, ..., \mu_{s=S,t=T})$, and similarly for π , 100 we modeled distributional location and (when applicable) zero inflation parameters as: 101

$$\log(\boldsymbol{\mu}) = \alpha^{(\boldsymbol{\mu})} + \mathbf{X}\boldsymbol{\beta}^{(\boldsymbol{\mu})} + \boldsymbol{\phi}^{(\boldsymbol{\mu})} + \log(\boldsymbol{a}),$$

$$\operatorname{logit}(\boldsymbol{\pi}) = \alpha^{(\pi)} + \mathbf{X}\boldsymbol{\beta}^{(\pi)} + \boldsymbol{\phi}^{(\pi)},$$

where $\alpha^{(\mu)}$ and $\alpha^{(\pi)}$ are scalar intercept parameters, **X** is a known $(S \times T) \times p$ design matrix, where p is the number of input features, $\beta^{(\mu)}$ and $\beta^{(\pi)}$ are column vector parameters of length p, $\phi^{(\mu)}$ and $\phi^{(\pi)}$ are column vector parameters of length $S \times T$ containing spatiotemporal adjustments, and \boldsymbol{a} is a known offset vector of areas for spatial unit s = 1, 2, ..., S, repeated T times.

107 Burn area

We developed five candidate models for burn area, each of which specified a different distribution for burn areas (Reed and McKelvey 2002; Hernandez et al. 2015), including the generalized Pareto (Hosking and Wallis 1987), tapered Pareto (Schoenberg, Peng, and Woods 2003), lognormal, gamma, and Weibull distributions. We evaluated each model in terms of test set log likelihood, and posterior predictive checks for burn area extremes. We defined the response y_i as the number of acres burned over 1000 for the i^{th} fire event, which occurred in spatial unit s_i and time step t_i .

Because each burn area distribution has a different parameterization, we included covariate ef-115 fects in a distribution-specific way. For the generalized Pareto distribution (GPD), we assumed 116 a positive shape parameter, leading to a Lomax distribution for exceedances (Bermudez et al. 117 2009). The GPD and Lomax shape parameters are related by $\kappa^{(GPD)} = 1/\kappa^{(L)}$, and the GPD 118 scale parameter is related to the Lomax scale and shape parameters by $\sigma^{(GPD)} = \sigma^{(L)} / \kappa^{(L)}$. 119 We introduced covariate dependence via the Lomax scale parameter using a log link. For 120 event i, $\log(\sigma_i^{(L)}) = \alpha + X_{(s_i,t_i)}\beta + \phi_{s_i,t_i}$, where α is an intercept parameter, β is a length 121 p vector of coefficients, $X_{(s_i,t_i)}$ is a row vector from X, and ϕ_{s_i,t_i} is a spatiotemporal ad-122 justment for s_i and t_i . For the tapered Pareto model, we modeled the shape parameter 123 as $\log(\kappa_i) = \alpha + X_{(s_i,t_i)}\beta + \phi_{s_i,t_i}$. The lognormal model included covariate dependence via 124 the location parameter: $\mu_i = \alpha + X_{(s_i,t_i)} \beta + \phi_{s_i,t_i}$. The gamma model used a log link for 125 the expected value: $\log(E(y_i)) = \alpha + \mathbf{X}_{(s_i,t_i)} \boldsymbol{\beta} + \phi_{s_i,t_i}$. Last, we modeled the Weibull scale 126 parameter as $\log(\sigma_i) = \alpha + X_{(s_i,t_i)}\beta + \phi_{s_i,t_i}$. More detail on the parameterization of each 127 burn area distribution is provided in the Supporting Information. 128

129 Accounting for nonlinear forcing

The design matrix **X** was constructed to allow for spatially varying nonlinear effects of housing density and meteorological drivers. We used B-splines to account for nonlinearity (Figure 2) and allowed the coefficients for each basis vector to vary spatially (Wood 2017). First, we constructed univariate B-splines for log housing density, wind speed, same month precipitation, previous 12 month precipitation, air temperature, and humidity, with five degrees of freedom (including an intercept) for each variable. This step generated 30 basis vectors (five for each of six variables).

To allow for spatial variation in these nonlinear effects, we added interaction effects between each of the basis vectors and ecoregions (Brezger and Lang 2006; Kneib, Hothorn, and Tutz 2009). The hierarchical nesting of ecoregion designations (Figure 1B-D) lends itself to such interactions. Conceptually, coefficients in a level 3 ecoregion may be related to coefficients in the level 2 ecoregion containing the level 3 region, the level 1 ecoregion containing the level 2 region, and a global effect. The coefficient associated with a basis vector for any level

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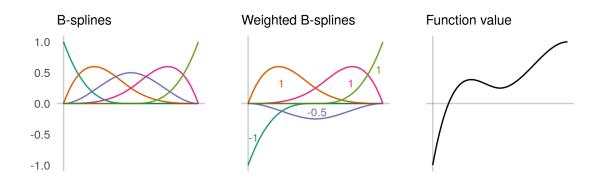


Figure 2. Conceptual figure to illustrate the use of B-splines to construct nonlinear functions. In the left panel, five B-spline vectors are shown, which map values of an input variable (on the x-axis) to a value on the y-axis. The middle panel shows the same B-spline vectors, but weighted (multiplied) by real numbers, with the weights illustrated as annotations. These weighted B-spline vectors are summed to produce the values of a nonlinear function (right panel).

3 ecoregion is treated as a sum of a global effect, a level 1 ecoregion adjustment, a level 2 143 ecoregion adjustment, and a level 3 ecogregion adjustment. Thus, for every univariate basis 144 vector, we included interaction effects with ecoregion at each of the three ecoregion levels. 145 This allows borrowing of information across space (level 3 ecoregions in a level 2 ecoregion are 146 often adjacent), and for regions that are ecologically similar. We also included adjustments 147 on the global intercept for each level 1, 2, and 3 ecoregion to account for spatial variation that 148 is unrelated to climate or housing density. This specification induces sparsity in X that we 149 exploit to increase the efficiency of computing μ and π . In total, X has p = 3,472 columns, 150 with 97% zero entries. 151

¹⁵² Prior specification

To avoid overfitting, we used a regularized horseshoe prior on the coefficients associated with the spatially varying nonlinear effects described above (Piironen, Vehtari, and others 2017). This prior places high probability close to zero, while retaining heavy enough tails that nonzero coefficients are not shrunk too strongly toward zero. This is consistent with

our prior expectation that most of the coefficients associated with the columns in X were 157 close to zero. For the zero inflated count models, we used a multivariate horseshoe to allow 158 information sharing between the zero inflated and distribution specific location parameters 159 (Peltola et al. 2014). For the remaining count models and all burn area models, this 160 was a univariate horseshoe prior. Spatiotemporal random effects were constructed using 161 a temporally autoregressive, spatially intrinsically autoregressive formulation (Besag and 162 Kooperberg 1995; Banerjee, Carlin, and Gelfand 2014). Details of these priors and the 163 resulting joint distributions are provided in the Supporting Information. 164

¹⁶⁵ Posterior predictive inference and extremes

We used the posterior predictive distribution to check each model and make inference on extremes. The posterior predictive distribution provides a distribution for replications of observed data (y^{rep}), and predictions of future data (Gelman et al. 2013). Conceptually, for a "good" model, y^{rep} should be similar to observed training data y, and future predictions should be similar to future data. Distributions over both quantities can be obtained by conditioning y and marginalizing over model parameters θ , e.g., $[y^{\text{rep}}|y] = \int [y^{\text{rep}}|\theta][\theta|y]d\theta$.

Posterior predictive distributions facilitate model checks that compare predicted and observed 172 test statistics (Gelman, Meng, and Stern 1996). To evaluate whether a model captures 173 tail behavior, one can compare an empirical maximum $(T(y) = \max(y))$ to the predicted 174 distribution of maxima $T(y^{rep})$. We also include predictive checks for the proportion of 175 zero counts, and totals for count and burn area models. Posterior predictive inference for 176 maxima is similar in spirit to the MEV approach. Both obtain a distribution over maxima 177 by marginalizing over unknowns including the number of events, size of each event, and 178 parameters of their distributions (Marani and Ignaccolo 2015). However, a Bayesian approach 179 explicitly conditions on the observed data to obtain a posterior distribution of parameters. 180 Seeing this connection is useful in the context of including priors and propagating uncertainty 181 to derived parameters, including total burn areas and probabilities of million acre wildfires. 182

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¹⁸³ Parameter estimation

We used a combination of variational approximations and Hamiltonian Monte Carlo methods 184 to sample from the posterior distributions of count and burn area models. A variational 185 approximation (Kucukelbir et al. 2015) was used for count models to quickly identify a 186 preferred model and avoid excessive multi-day model runs. Models were fit in the Stan 187 probabilistic programming language using the rstan package (Carpenter et al. 2016; Stan 188 Development Team 2018). The best performing count model and all burn area models were 189 fit using the No-U-Turn Sampler (Hoffman and Gelman 2014). We ran four chains for 1000 190 iterations each, and discarded the first 500 iterations as warmup. Convergence was assessed 191 using visual inspection of trace plots, with potential scale reduction statistic values $\hat{R} \ge 1.1$ 192 as an indicator convergence failure (Brooks and Gelman 1998). 193

¹⁹⁴ Implementation

All data processing, model fitting, and visualization were implemented with open source 195 software, primarily in the R programming language (R Core Team 2017), and wrapped in 196 a reproducible workflow via GNU Make and Docker (Stallman, McGrath, and Smith 2004; 197 Boettiger 2015). Data cleaning and transformation required the R packages assertthat (Wick-198 ham 2017a), lubridate (Grolemund and Wickham 2011), Matrix (Bates and Maechler 2018), 199 pbapply (Solymos and Zawadzki 2018), splines (R Core Team 2018), tidyverse (Wickham 200 2017b), and zoo (Zeileis and Grothendieck 2005). Spatial data were processed with raster 201 (Hijmans 2017), rgdal (Bivand, Keitt, and Rowlingson 2018), sf (Pebesma 2018), and spdep 202 (Bivand and Piras 2015). Finally, we used complot (Wilke 2017), ggrepel (Slowikowski 2018), 203 ggthemes (Arnold 2018), patchwork (Pedersen 2017), and RColorBrewer (Neuwirth 2014) for 204 visualization. The manuscript was written in R Markdown (Allaire et al. 2018). Analyses 205 were run on an Amazon Web Services m5.2xlarge EC2 instance with four physical cores and 206 32 GB of RAM, and the whole workflow requires \approx 72 hours. All code to reproduce the 207 analysis is available on GitHub at https://github.com/mbjoseph/wildfire-extremes (Joseph 208 2018). 209

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Model	Holdout log likelihood
ZI Negative binomial	-3068 (55)
Negative binomial	-3090 (60)
ZI Poisson	-3526 (71)
Poisson	-4194 (124)

 Table 1. Performance of count models on the test set in descending order. Posterior means are provided with standard deviations in parentheses.

$_{210}$ **Results**

²¹¹ Wildfire occurrence

The zero-inflated negative binomial distribution performed best on the held-out test set (Table 1), and was able to recover the proportion of zeros, count maxima, and count totals in posterior predictive checks for both the training and test data (Figure 3). All of the other count models that we considered exhibited lack of fit to at least one of these statistics in posterior predictive checks. Hereafter, we report results from the zero-inflated negative binomial model.

Minimum relative humidity and maximum air temperature had the strongest effects on both 218 the zero-inflation component and the expected value of the negative binomial component 219 (Figure 4, posterior median for ρ : 0.663, 95% credible interval (CI): 0.361 - 0.86). The 220 model uncovered unique effects of meteorological variables at level 1, 2, and 3 ecoregions 221 (Figure 5). For example, a positive interaction effect between the second air temperature 222 basis vector and the L1 Great Plains ecoregions indicates that the expected number of 223 wildfires in plains ecoregions with cold conditions is high relative to other ecoregions. The 224 Ozark/Ouachita-Appalachian forest and Ozark Highlands were also identified as having 225 region-specific temperature effects (Figure 5). Twelve month total precipitation also had 226 region specific effects in the Mississippi Alluvial and Southeast Coastal Plains ecoregion, 227 where it tended to reduce fire risk (Figure 5). In contrast, increasing cumulative twelve month 228 precipitation tended to increase risk in desert ecoregions (Figure 4). Housing density showed 229

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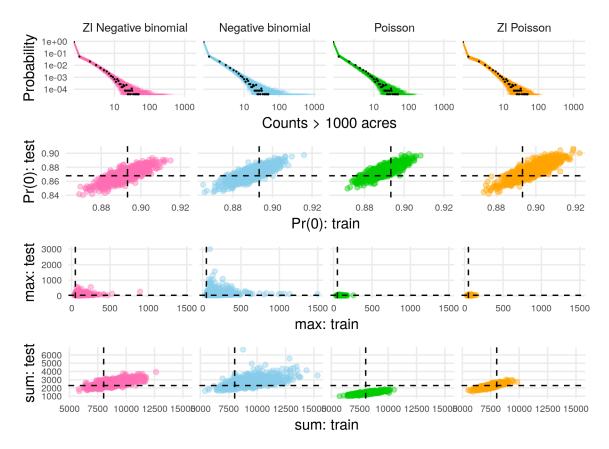


Figure 3. Count predictive checks. Row one shows observed count frequencies as black points and predicted frequencies as lines. Rows two, three, and four show predicted proportions of zeros, maxima, and sums (respectively) in the training and test data, with empirical values as dashed lines. Rows two through four facilitate comparison of performance on training and test sets. Ideally, model predictions cluster around the dashed lines for both the training (x-axis direction) and test (y-axis direction) sets, leading to a tight cluster of points at the intersection of the dashed lines.

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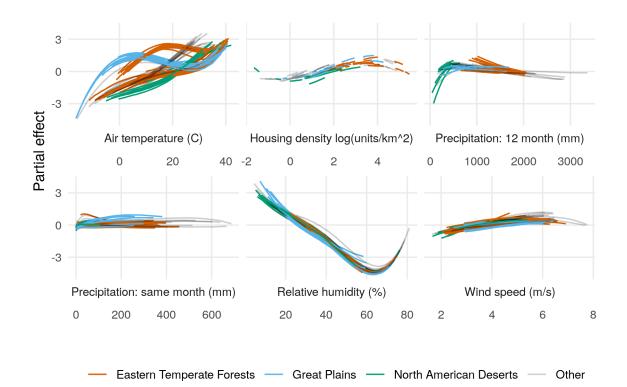


Figure 4. Partial effects on the log-transformed negative binomial mean component of the zero-inflated negative binomial model for each level 3 ecoregion, colored by level 1 ecoregion. Lines are posterior medians. Results are similar for the zero-inflation component.

a unimodal relationship to expected count (Figure 4), with lower expected counts in sparsely
populated ecoregions, and higher expected counts with moderately populated ecoregions.

Posterior 95% credible interval coverage for the number of fires over 1000 acres in the test 232 set was 98.9%. The lowest test set interval coverage was 91.7%, in the Cross Timbers L3 233 ecoregion. When observed counts fell outside the 95% prediction interval, counts were larger 234 than predicted 100% of the time. The largest difference between observed numbers and 235 predicted 97.5% posterior quantiles (the upper limit for the 95% credible interval) occurred 236 for the Columbia Mountains/Northern Rockies L3 ecoregion in August 2015, when 36 fires 237 over 1000 acres occurred and at most 23.025 were predicted. For nearly half of the level 238 3 ecoregions (44 of 85), accounting for 40.2% of the land area of the contiguous U.S., the 239 zero-inflated negative binomial model had 100% test set prediction interval coverage. 240

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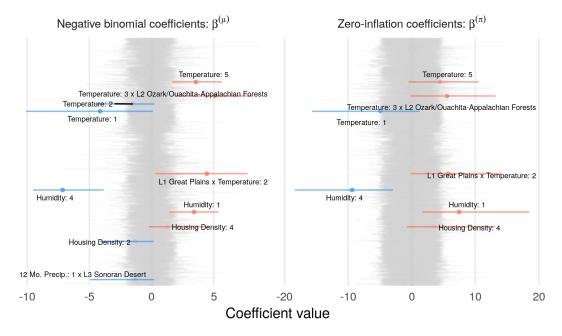


Figure 5. Caterpillar plots of zero inflated negative binomial model coefficients, $\beta^{(\mu)}$ (left) and $\beta^{(\pi)}$ (right). Horizontal line segments denote 95% credible intervals. Grey segments indicate coefficients with a less than 87% posterior probability of being positive or negative, and colored segments indicate coefficients that are probably positive (red) or negative (blue). B-spline vectors are indicated by colons, e.g., Humidity:1 indicates the first basis vector corresponding to humidity. Interactions between variables a and b are represented as Intxn(a x b). Level 1 ecoregions are represented by L1 ecoregion name, and L2 and L3 indicate level 2 and 3 ecoregions.

²⁴¹ Wildfire burned areas

The lognormal distribution performed best on the test set (Table 2), and captured tail-242 behavior better than other burn area distributions (Figure 6). The GPD model was too 243 heavy-tailed to adequately capture the pattern in the empirical data, predicting fires far 244 larger than those observed in the training and test sets (Figure 6). The tapered Pareto 245 distribution was too light-tailed (Figure 6). The gamma and Weibull models performed very 246 poorly overall on the test set (Table 2), apparently due to a lack of congruence between the 247 shapes of these distributions and the actual burn area distribution. Despite a poor fit to the 248 bulk of the wildfire burn area distribution, both performed adequately in the upper tails 249

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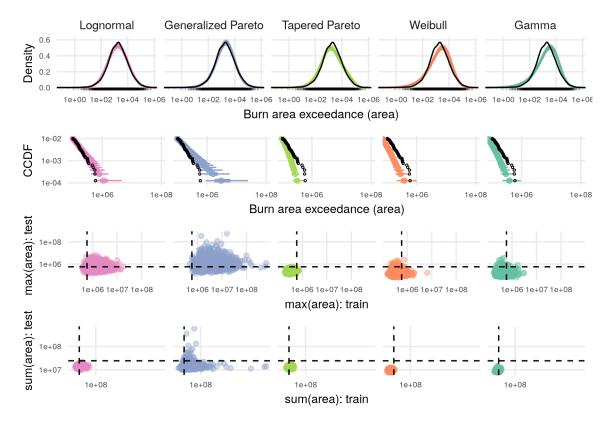


Figure 6. Predictive checks for burn area models. The top row shows predicted density in color and empirical density for the training set in black, which reveals overall lack of fit for the gamma and Weibull models. Row two shows the complementary cumulative distribution function (CCDF) at the tails, with 95% and 50% prediction intervals shown in color and observed data as black points, which shows that the Generalized Pareto distribution predicts values that are too extreme. The third and fourth rows show checks for maximum and total burn areas in the training and test set, with observed values as dashed lines and posterior draws as colored points. These final two rows facilitate checks for summary statistics on both the training and test set, with the ideal model generating predictions (colored points) clustered close to where the dashed lines intersect.

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Model	Holdout log likelihood
Lognormal	-22341 (39)
Generalized Pareto	-22375 (42)
Tapered Pareto	-22388 (48)
Weibull	-23499 (230)
Gamma	-26191 (908)

 Table 2. Performance of burn area models on the test set in descending order. Posterior

 means are provided with standard deviations in parentheses.

(Figure 6). Hereafter we present results for the lognormal model, which had the highest test
set log likelihood and captured tail behavior of the empirical fire size distribution.

Relative humidity was the primary driver of expected burn area for a fire event (Figure 7A). 252 The first basis vector for mean daily minimum relative humidity was the only coefficient with 253 a 95% credible interval that did not include zero (posterior median: 1.66, 95% CI: (0.57 -254 (2.28)). This nonlinear effect can be observed in Figure 7B as an increase in the expected burn 255 area below 20% mean daily minimum humidity. This leads to a seasonality gradient among 256 ecoregions of expected fire sizes, with little or no seasonal signal in typically humid ecoregions 257 such as Marine West Coast Forests of the Pacific Northwest, and seasonal oscillations in 258 ecoregions that have periodic fluctuations between dry and humid conditions such as the 259 Temperate Sierras (Figure 7C). There was not strong evidence that meteorological variables 260 had spatially variable effects on expected wildfire burn area. 261

Overall, 95% posterior predictive interval coverage in the test set for burn areas was 92.3%. 262 The lowest test set coverage was 0%, for the Eastern Great Lakes Lowlands L3 ecoregion, 263 followed by 50%, for the Central California Valley L3 ecoregion, though these ecoregions had 264 just 1 and 2 wildfire events in the test set. When observed fire sizes fell outside the 95%265 prediction interval, 23.3% of wildfires were smaller than predicted, and 76.7% of wildfires 266 were larger than predicted. The largest discrepancy between the actual size of a wildfire and 267 the predicted 97.5% posterior quantile was observed with the Wallow Fire in 2011 which 268 burned 563,655 acres, but the predicted upper limit for size was 50,136. We investigate this 269

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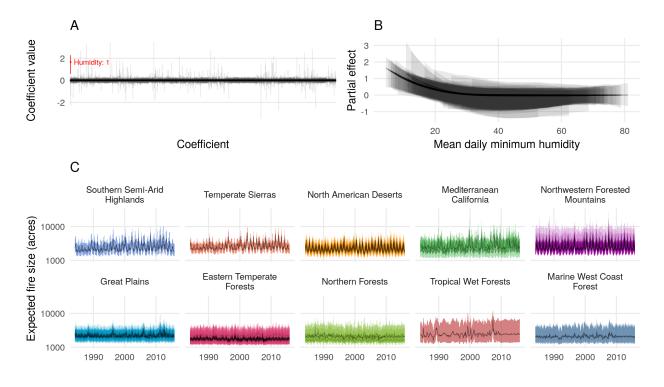


Figure 7. A. Estimated posterior medians and 95% credible intervals for each of the 3,473 coefficients associated with expected burn area. Only one coefficient - the first basis vector for humidity - had a 95% credible interval that excluded zero, shown in red. This effect is visualized in **B**. Partial effects of mean daily minimum humidity for each level 3 ecoregion, with posterior medians drawn as lines, and the 95% credible intervals as ribbons. **C**. Monthly time series of expected fire sizes for every level 3 ecoregion, faceted and colored by level 1 ecoregions sorted by mean humidity. Lines are posterior medians and ribbons are 95% credible intervals.

discrepancy further in the case study below. The lognormal burn area model achieved 100% interval coverage in 26 of 66 ecoregions that had wildfire events in the test set, accounting for 29% of the land area of the contiguous U.S.

²⁷³ Inference on extremes

By combining the output of the event count and burn area models, we derived posterior prediction intervals for the size of the largest fire in a month for each region (the "burn area

maximum"), integrating over uncertainty in the number of fires, as well as the lognormal 276 mean and standard deviation for burn area. In the holdout period from 2010 to 2015, a 99%277 prediction interval achieved 77.6% interval coverage, with 13.6% of the burn area maxima 278 (109 fire events) being larger than predicted (Figure 8). The model predicted the total area 279 burned over the entire contiguous United States in test period from 2010 to 2015 to be 280 28,407,396 (95% CI: (18,236,854 - 52,279,618) and the actual value was 26,639,835. While 281 fires over a million acres in size have happened historically, all fires in the training and 282 test sets were below one million acres. If we extrapolate, the probability that at least one 283 fire in the contiguous U.S. exceeded one million acres in the period from 2010 to 2015 was 284 estimated to be between 0.171 and 0.661 (95% CI), with a posterior median of 0.33. The 285 highest estimated probability for a million acre event was 0.014 (posterior median), with 286 a 95% CI of (0, 0.272) seen for the Southwestern Tablelands ecoregion in June 2011. The 287 second highest probability for a million acre event was 0.003 (posterior median), with a 95%288 CI of (0, 0.048) seen for the Arizona/New Mexico Mountains ecoregion in June 2011. 289

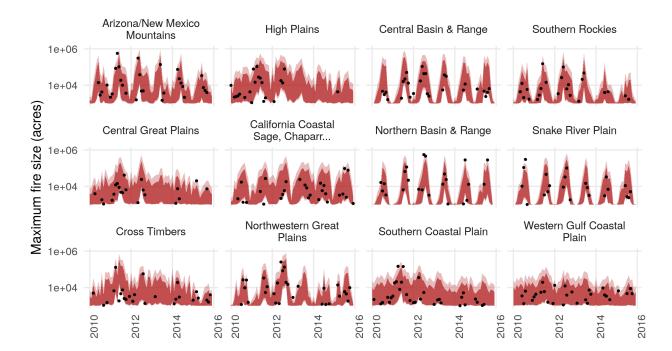


Figure 8. Posterior 99% (light red) and 95% (dark red) prediction intervals for the burn area of the largest fire event by month and level 3 ecoregion in the test set, shown for ecoregions with wildfires in more than 20 months. Empirical maxima are shown as black dots.

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²⁹⁰ Error analysis case study: the 2011 Wallow Fire

To better understand how well the model could or could not anticipate notable extreme 291 events, and why, we used the largest fire in the test set as a case study. The Wallow Fire 292 was accidentally ignited on May 29, 2011 by two campers in the L3 Arizona/New Mexico 293 Mountains ecoregion. It burned through the month of June and into early July. The model 294 underpredicted the total burn area of the Wallow Fire. Integrating over uncertainty in the 295 predicted number of fires and expected fire size, the 99% credible interval for the maximum 296 fire size for May 2011 was (1,820 - 270,304) acres, but the Wallow Fire is recorded as 563,655 297 acres in the MTBS data. 298

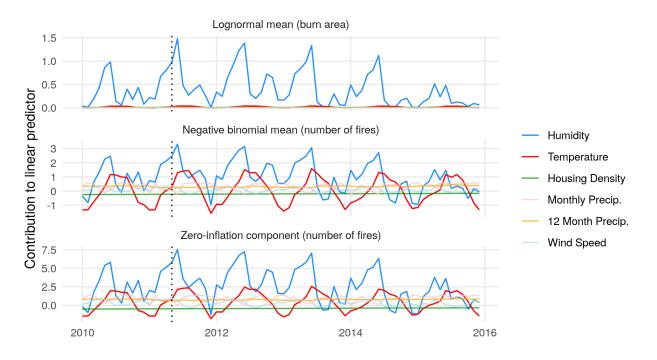


Figure 9. Posterior median contribution of each input variable to the linear predictor function of model components for the Arizona/New Mexico Mountains level 3 ecoregion from 2010-2016. A dotted vertical line marks May 2011, when the Wallow Fire ignited. Vertical positions of colored lines show contributions to the linear predictor function of each model component.

We evaluated the contribution of each covariate to the linear predictor functions of the three model components (lognormal mean for burn areas, negative binomial mean for counts, and

the logit probability of the zero-inflation component) to understand why these predictions 301 differed. We defined the contribution of a variable as the dot product of the elements in 302 the design matrix \mathbf{X} corresponding to a particular driver variable (e.g., humidity), and the 303 estimated coefficients in β corresponding to that variable. This provides a quantitative 304 measure of how each input variable contributes to the linear predictor for an ecoregion, and 305 incorporates the overall, level 1, level 2, and level 3 ecoregion adjustments on these effects. 306 Humidity is the primary driver of variation in the model's predictions overall, and June 2011 307 - the month after ignition - favored more large fires, with drier, hotter conditions (Figure 9). 308 The 99% credible interval for June 2011 was (10,053 - 1,005,827) acres, which contains the 309 true value. Evidently, conditions in May that drove (under)predictions of maximum burn 310 area were not representative of the conditions over most of the Wallow Fire's duration. The 311 failure of the model to correctly predict the size of the Wallow fire suggests potential avenues 312 for improvement, discussed below. 313

314 Discussion

Extreme wildfires are often devastating, but perhaps they need not be surprising. By allowing the non-linear effects of weather and housing density to vary across space, we were able to achieve good predictive accuracy for fire extremes over a five-year prediction window. We estimate a moderate chance of wildfires larger than what has been observed in recent decades, perhaps even over one million acres. Such predictions can support short to medium term wildfire management and probabilistic hazard assessment.

Driving a model with meteorological features raises challenges related to predictive uncertainty and covariate shift - a change in the underlying distribution of forcing variables, potentially outside of the historic range. Ideally, this uncertainty would be propagated forward in a predictive model, possibly through stacking of predictive distributions that are generated from multiple models of future climate dynamics (Yao et al. 2017). But, even if one had a perfect forecast, novel conditions present a challenge for predictive modeling (Quionero-Candela et al. 2009). For example, the High Plains ecoregion had its highest monthly precipitation, lowest 12 month running precipitation, driest, hottest, and windiest conditions in the test set period, so that the range of environmental conditions in the training data did not encompass the range of future conditions. Extrapolating beyond the range of training inputs is generally difficult, but the hierarchical spatial effect specification used here allows partial pooling among climatically similar ecoregions that can inform such predictions, unlike models fit separately to disjoint spatial regions.

Human-caused climate change is expected to increase fire activity in the western U.S. (Rogers 334 et al. 2011; Westerling et al. 2011; Moritz et al. 2012; Abatzoglou and Williams 2016) 335 and elsewhere (Flannigan et al. 2009), which when coupled with the nonlinear effect of 336 human-density provides a key inferential wrinkle. While most U.S. ecoregions are increasing 337 in human density over time, some of these ecoregions are in the range of values in which 338 this increases the expected number of large fires, while others are so populated that further 339 increases would reduce the chance of a large fire. The hump-shaped effect of human density 340 on the expected number of large fires is likely driven by ignition pressure and fire suppression 341 (Balch et al. 2017). As human density increases from zero, ignition pressure increases, but 342 eventually landscapes become so urbanized, fragmented, and/or fire-suppressed that wildfire 343 risk decreases (Syphard et al. 2007; Bowman et al. 2011; Bistinas et al. 2013; Knorr et al. 344 2013; Mcwethy et al. 2013; Syphard et al. 2017; Nagy et al. 2018). At intermediate density, 345 wildfire dynamics respond to human ignition and altered fuel distributions (Guyette, Muzika, 346 and Dey 2002), but these responses depend on environmental context and characteristics of 347 the human population (Marlon et al. 2008; Li et al. 2009). This model indicates that the 348 combination of moderate to high human density and dry conditions would nonlinearly increase 340 the chance of an extreme fire event. Both human density and dryness are expected to increase 350 in the future across large swaths of the U.S. (Llovd, Sorichetta, and Tatem 2017; Stavros 351 et al. 2014, Radeloff et al. (2010)), with potential implications for human mortality, health 352 risks from smoke and particulate emission, and the financial burden of wildfire management 353 (Reid et al. 2016; Radeloff et al. 2018). 354

This work points to promising directions for future predictive efforts. Default choices such as Poisson and GPD distributions should be checked against alternative distributions. Further,

the predictive skill of this model seems to suggest that ordinary events provide information on 357 extremes, which would not be the case if the generative distribution of extremes was completely 358 unique. Previous case studies have identified that extremes or anomalies in climatological 359 drivers play a role in the evolution of extreme wildfires (Peterson et al. 2015), but for this 360 work, monthly averages of climatological drivers over fairly large spatial regions were used, 361 which may smooth over anomalous or extreme conditions. Enhancing the spatiotemporal 362 resolution of predictive models could better represent climatic and social drivers of fire 363 dynamics and provide localized insights into fire dynamics to inform decision-making. This 364 raises computational challenges, but recent advances in distributed probabilistic computing 365 (Tran et al. 2017), efficient construction of spatiotemporal point processes (Shirota and 366 Banerjee 2018), and compact representations of nonlinear spatial interactions (Lee and 367 Durbán 2011) may provide solutions. 368

The Wallow Fire case study reveals at least one limitation of increasing the spatiotemporal 369 resolution. When the model predictions are driven by covariates that are summarized in space 370 and time (e.g. a mean across an ecoregion in a month), summary values may not represent 371 conditions that are most relevant to an event. With a discrete space-time segmentation, 372 events can occur at the boundary of a spatiotemporal unit, e.g., if a fire spreads into an 373 adjacent ecoregion or ignites on the last day of the month. Large wildfires can span months, 374 and a model that only uses conditions upon ignition to predict total burn area can fail to 375 account for conditions that change over the course of the event. Modeling ignitions as a point 376 process in continuous space and time (Brillinger, Preisler, and Benoit 2003), and explicitly 377 modeling subsequent fire duration and spatial dynamics could better separate conditions that 378 ignite fires from those that affect spread. Such an approach might be amenable to including 379 information on fuel continuity, which is likely to limit the size of extremely large fires and 380 did not factor into the current models predictions (Rollins, Morgan, and Swetnam 2002; 381 Hargrove et al. 2000). To the extent that a model reflects the generative process for extreme 382 events, the decomposition of contributions to the model's predictions may provide insight 383 into attribution for meteorological and anthropogenic drivers of extremes. 384

³⁸⁵ This paper presents and evaluates a statistical approach to explain and predict extreme

wildfires that incorporates spatially varying non-linear dynamics. The model reveals considerable differences in fire dynamics among ecoregions spanning the mountain west to the great plains, deserts, and eastern forests, and suggests a decent chance of very large fires exceeding one million acres in the contiguous U.S. Predictive approaches such as this can inform decision-making by placing probabilistic bounds on the number of wildfires and their sizes, while provide deeper insights into wildfire ecology.

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Supporting Information

Prior specifications

Prior distributions were chosen to regularize coefficients on the distribution specific means $\beta^{(\mu)}$ and structural zero parameters $\beta^{(\pi)}$. We used a regularized horseshoe prior on these coefficients, which shrinks irrelevant coefficients towards zero, while regularizing nonzero coefficients (Piironen, Vehtari, and others 2017). For zero-inflated models, we used a multivariate version of the regularized horseshoe (Peltola et al. 2014):

$$\begin{pmatrix} \beta_j^{(\mu)} \\ \beta_j^{(\pi)} \end{pmatrix} \sim \mathcal{N} \left(\mathbf{0}, \begin{vmatrix} \tau_1^2 \tilde{\lambda}_{1,j}^2 & \rho \tau_1 \tau_2 \tilde{\lambda}_{1,j} \tilde{\lambda}_{2,j} \\ \rho \tau_1 \tau_2 \tilde{\lambda}_{1,j} \tilde{\lambda}_{2,j} & \tau_2^2 \tilde{\lambda}_{2,j}^2 \end{vmatrix} \right) \right),$$
$$\tilde{\lambda}_{m,j}^2 = \frac{c_m^2 \lambda_j^2}{c_m^2 + \tau_m^2 \lambda_j^2},$$

for each response dimension m = 1, 2 and coefficient j = 1, ..., p. Here ρ is a correlation parameter, τ_1 and τ_2 are global variance hyperparameters, c_1 and c_2 are hyperparameters that determine the amount of shrinkage on the largest coefficients, and λ_j is a local scale parameter drawn from a half-Cauchy distribution that control the amount of shrinkage applied to coefficient j (Piironen, Vehtari, and others 2017). With this prior specification, information can be shared across the two response dimensions through the correlation parameter ρ , and/or through the local scale parameters λ_j . For count models without structural zeros (the Poisson and negative binomial models), this multivariate prior simplifies to a univariate regularized horseshoe prior.

Spatiotemporal random effects were constructed using a temporally autoregressive, spatially intrinsically autoregressive formulation (Besag and Kooperberg 1995; Banerjee, Carlin, and Gelfand 2014). Temporarily suppressing the superscript that indicates whether the effects are on μ or π , and denoting column t from an $S \times T \Phi$ as ϕ_t we have:

$$\boldsymbol{\phi}_{t=1} \sim \mathrm{N}(\mathbf{0}, (\tau^{(\phi)}(\mathbf{D} - \mathbf{W}))^{-1})$$

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$$\boldsymbol{\phi}_t \sim \mathcal{N}(\eta \boldsymbol{\phi}_{t-1}, (\tau^{(\phi)}(\mathbf{D} - \mathbf{W}))^{-1}), \quad t = 2, ..., T$$

where η is a temporal dependence parameter, $\tau^{(\phi)}$ is a precision parameter, **D** is an $S \times S$ diagonal matrix with entries corresponding to the number of spatial neighbors for each spatial unit, and **W** is an $S \times S$ spatial adjacency matrix with nonzero elements only when spatial unit *i* is a neighbor of spatial unit *j* ($w_{i,j} = 1$ if *i* is a neighbor of *j*, and $w_{i,j} = 0$ otherwise, including $w_{i,i} = 0$ for all *i*). $\tau^{(\phi)}$ is a precision parameter. We imposed a soft identifiability constraint that places high prior mass near $\sum_{s=1}^{S} \phi_{t,s}^* = 0$ for all *t*.

We applied a univariate regularized horseshoe prior to all β coefficients in burn area models (Piironen, Vehtari, and others 2017):

$$\beta_j \sim \mathrm{N}\left(0, \tau^2 \tilde{\lambda}_j^2\right), \quad \tilde{\lambda}_j^2 = \frac{c^2 \lambda_j^2}{c^2 + \tau^2 \lambda_j^2};$$

Spatiotemporal random effects were constructed in the same way as for the count models.

Joint distributions

Here we provide the unnormalized posterior densities for each model. Square brackets represent a probability mass or density function. Parameterizations for model likelihoods are provided first, followed by the factorization of the joint distribution, with explicit priors.

Poisson wildfire count model

We used the following parameterization of the Poisson distribution:

$$[n|\mu] = \frac{\mu^n e^{-\mu}}{n!},$$

where μ is the mean and variance.

$$[\boldsymbol{\beta}^{(\mu)}, \boldsymbol{\alpha}^{(\mu)}, \boldsymbol{\phi}, \boldsymbol{\sigma}^{(\phi)}, \boldsymbol{\eta}, \boldsymbol{\lambda}, c, \tau \mid \mathbf{N}] \propto \prod_{s=1}^{S} \prod_{t=1}^{T} [n_{s,t} | \boldsymbol{\beta}^{(\mu)}, \boldsymbol{\alpha}^{(\mu)}, \boldsymbol{\phi}_{s,t}] \times [\boldsymbol{\phi}_{1} | \boldsymbol{\sigma}^{(\phi)}] \prod_{t=2}^{T} [\boldsymbol{\phi}_{t} | \boldsymbol{\phi}_{t-1}, \boldsymbol{\sigma}^{(\phi)}, \boldsymbol{\eta}] \times \prod_{j=1}^{p} [\beta_{j}^{(\mu)} | \lambda_{j}, c, \tau] [\lambda_{j}] \times [\boldsymbol{\sigma}^{(\phi)}] [\boldsymbol{\eta}] [c] [\tau] [\boldsymbol{\alpha}^{(\mu)}]$$

$$= \prod_{s=1}^{S} \prod_{t=1}^{T} \operatorname{Poisson}(n_{s,t} | \exp(\alpha^{(\mu)} + \mathbf{X}_{(s,t)} \boldsymbol{\beta}^{(\mu)} + \phi_{s,t})) \times \operatorname{Normal}(\boldsymbol{\phi}_{1} | \mathbf{0}, ((\boldsymbol{\sigma}^{(\phi)})^{-2} (\mathbf{D} - \mathbf{W}))^{-1}) \times \prod_{t=2}^{T} \operatorname{Normal}(\boldsymbol{\phi}_{t} | \eta \boldsymbol{\phi}_{t-1}, ((\boldsymbol{\sigma}^{(\phi)})^{-2} (\mathbf{D} - \mathbf{W}))^{-1}) \times \prod_{j=1}^{p} \operatorname{Normal}\left(\beta_{j}^{(\mu)} | \mathbf{0}, \frac{\tau^{2} c^{2} \lambda_{j}^{2}}{c^{2} + \tau^{2} \lambda_{j}^{2}}\right) \times \operatorname{Cauchy}^{+}(\lambda_{j} | \mathbf{0}, 1) \times \left((\boldsymbol{\phi}^{(\phi)}) | \mathbf{0}, \frac{\tau^{2} c^{2} \lambda_{j}^{2}}{c^{2} + \tau^{2} \lambda_{j}^{2}} \right) \times \operatorname{Cauchy}^{+}(\lambda_{j} | \mathbf{0}, 1) \times \left((\boldsymbol{\phi}^{(\phi)}) | \mathbf{0}, \frac{\tau^{2} c^{2} \lambda_{j}^{2}}{c^{2} + \tau^{2} \lambda_{j}^{2}} \right) \times \operatorname{Cauchy}^{+}(\lambda_{j} | \mathbf{0}, 1) \times \left((\boldsymbol{\phi}^{(\phi)}) | \mathbf{0}, \frac{\tau^{2} c^{2} \lambda_{j}^{2}}{c^{2} + \tau^{2} \lambda_{j}^{2}} \right) \times \operatorname{Cauchy}^{+}(\lambda_{j} | \mathbf{0}, 1) \times \operatorname{C$$

Normal⁺($\sigma^{(\phi)}|0, 1^2$) × Beta($\eta|1, 1$) × Inv-Gamma($c^2|2.5, 10$) ×

Normal⁺ $(\tau|0, 5^2) \times \text{Normal}(\alpha^{(\mu)}|0, 5^2).$

Negative binomial wildfire count model

We used the following parameterization of the negative binomial distribution:

$$[n|\mu,\delta] = \binom{n+\delta-1}{n} \left(\frac{\mu}{\mu+\delta}\right)^n \left(\frac{\delta}{\mu+\delta}\right)^{\delta},$$

where μ is the mean, and δ is a dispersion parameter.

$$[\boldsymbol{\beta}^{(\mu)}, \boldsymbol{\alpha}^{(\mu)}, \boldsymbol{\phi}, \boldsymbol{\sigma}^{(\phi)}, \boldsymbol{\eta}, \boldsymbol{\lambda}, c, \tau, \delta \mid \mathbf{N}] \propto \prod_{s=1}^{S} \prod_{t=1}^{T} [n_{s,t} | \boldsymbol{\beta}^{(\mu)}, \boldsymbol{\alpha}^{(\mu)}, \phi_{s,t}, \delta] \times [\boldsymbol{\phi}_{1} | \boldsymbol{\sigma}^{(\phi)}] \prod_{t=2}^{T} [\boldsymbol{\phi}_{t} | \boldsymbol{\phi}_{t-1}, \boldsymbol{\sigma}^{(\phi)}, \boldsymbol{\eta}] \times \prod_{j=1}^{p} [\beta_{j}^{(\mu)} | \lambda_{j}, c, \tau] [\lambda_{j}] \times [\boldsymbol{\sigma}^{(\phi)}] [\boldsymbol{\eta}] [c] [\tau] [\boldsymbol{\alpha}^{(\mu)}] [\delta]$$

$$= \prod_{s=1}^{S} \prod_{t=1}^{T} \text{Negative Binomial}(n_{s,t} | \exp(\alpha^{(\mu)} + \mathbf{X}_{(s,t)} \boldsymbol{\beta}^{(\mu)} + \phi_{s,t}), \delta) \times \\ \text{Normal}(\phi_1 | \mathbf{0}, ((\sigma^{(\phi)})^{-2} (\mathbf{D} - \mathbf{W}))^{-1}) \times \\ \prod_{t=2}^{T} \text{Normal}(\phi_t | \eta \phi_{t-1}, ((\sigma^{(\phi)})^{-2} (\mathbf{D} - \mathbf{W}))^{-1}) \times \\ \prod_{j=1}^{p} \text{Normal}\left(\beta_j^{(\mu)} | 0, \frac{\tau^2 c^2 \lambda_j^2}{c^2 + \tau^2 \lambda_j^2}\right) \times \text{Cauchy}^+(\lambda_j | 0, 1) \times \\ \text{Normal}^+(\sigma^{(\phi)} | 0, 1^2) \times \text{Beta}(\eta | 1, 1) \times \text{Inv-Gamma}(c^2 | 2.5, 10) \times \\ \text{Normal}^+(\tau | 0, 5^2) \times \text{Normal}(\alpha^{(\mu)} | 0, 5^2) \times \text{Normal}^+(\delta | 0, 5^2).$$

Zero-inflated Poisson wildfire count model

We used the following parameterization of the zero-inflated Poisson distribution:

$$[n|\mu,\pi] = I_{n=0}(1-\pi+\pi e^{-\mu}) + I_{n>0}\pi\frac{\mu^n e^{-\mu}}{n!},$$

where μ is the Poisson mean, and $1 - \pi$ is the probability of an extra zero.

$$\begin{split} [\boldsymbol{\beta}^{(\mu)}, \alpha^{(\mu)}, \boldsymbol{\beta}^{(\pi)}, \alpha^{(\pi)}, \boldsymbol{\phi}^{(\mu)}, \sigma^{(\phi,\mu)}, \eta^{(\mu)}, \boldsymbol{\phi}^{(\pi)}, \sigma^{(\phi,\pi)}, \eta^{(\pi)}, \boldsymbol{\lambda}, c, \tau, \rho \mid \mathbf{N}] \propto \\ \prod_{s=1}^{S} \prod_{t=1}^{T} [n_{s,t} | \boldsymbol{\beta}^{(\mu)}, \alpha^{(\mu)}, \boldsymbol{\beta}^{(\pi)}, \alpha^{(\pi)}, \boldsymbol{\phi}^{(\mu)}_{s,t}, \boldsymbol{\phi}^{(\pi)}_{s,t}] \times \\ [\boldsymbol{\phi}_{1}^{(\mu)} | \sigma^{(\phi,\mu)}] \prod_{t=2}^{T} [\boldsymbol{\phi}_{t}^{(\mu)} | \boldsymbol{\phi}^{(\mu)}_{t-1}, \sigma^{(\phi,\mu)}, \eta^{(\mu)}] \times \\ [\boldsymbol{\phi}_{1}^{(\pi)} | \sigma^{(\phi,\pi)}] \prod_{t=2}^{T} [\boldsymbol{\phi}_{t}^{(\pi)} | \boldsymbol{\phi}^{(\pi)}_{t-1}, \sigma^{(\phi,\pi)}, \eta^{(\pi)}] \times \\ \prod_{j=1}^{p} [\beta_{j}^{(\mu)}, \beta_{j}^{(\pi)} | \lambda_{j}, c, \tau, \rho] [\lambda_{j}] \times \\ [\sigma^{(\phi,\mu)}] [\sigma^{(\phi,\pi)}] [\eta^{(\mu)}] [\eta^{(\pi)}] [\alpha^{(\mu)}] [\alpha^{(\pi)}] [\rho] \prod_{m=1}^{2} [c_{m}] [\tau_{m}] \end{split}$$

Zero-inflated negative binomial wildfire count model

We used the following parameterization of the zero-inflated negative binomial distribution:

$$[n|\mu,\delta,\pi] = I_{n=0}(1-\pi+\pi\left(\frac{\delta}{\mu+\delta}\right)^{\delta}) + I_{n>0}\binom{n+\delta-1}{n}\left(\frac{\mu}{\mu+\delta}\right)^n\left(\frac{\delta}{\mu+\delta}\right)^{\delta}$$

where μ is the negative binomial mean, δ is the negative binomial dispersion, and , and $1 - \pi$ is the probability of an extra zero.

$$\begin{split} [\boldsymbol{\beta}^{(\mu)}, \alpha^{(\mu)}, \boldsymbol{\beta}^{(\pi)}, \alpha^{(\pi)}, \boldsymbol{\phi}^{(\mu)}, \sigma^{(\phi,\mu)}, \eta^{(\mu)}, \boldsymbol{\phi}^{(\pi)}, \sigma^{(\phi,\pi)}, \eta^{(\pi)}, \boldsymbol{\lambda}, c, \tau, \rho, \delta \mid \mathbf{N}] \propto \\ & \prod_{s=1}^{S} \prod_{t=1}^{T} [n_{s,t} | \boldsymbol{\beta}^{(\mu)}, \alpha^{(\mu)}, \boldsymbol{\beta}^{(\pi)}, \alpha^{(\pi)}, \phi^{(\mu)}_{s,t}, \phi^{(\pi)}_{s,t}, \delta] \times \\ & [\boldsymbol{\phi}_{1}^{(\mu)} | \sigma^{(\phi,\mu)}] \prod_{t=2}^{T} [\boldsymbol{\phi}_{t}^{(\mu)} | \boldsymbol{\phi}_{t-1}^{(\mu)}, \sigma^{(\phi,\mu)}, \eta^{(\mu)}] \times \\ & [\boldsymbol{\phi}_{1}^{(\pi)} | \sigma^{(\phi,\pi)}] \prod_{t=2}^{T} [\boldsymbol{\phi}_{t}^{(\pi)} | \boldsymbol{\phi}_{t-1}^{(\pi)}, \sigma^{(\phi,\pi)}, \eta^{(\pi)}] \times \\ & \prod_{j=1}^{p} [\beta_{j}^{(\mu)}, \beta_{j}^{(\pi)} | \lambda_{j}, c, \tau, \rho] [\lambda_{j}] \times \\ & [\sigma^{(\phi,\mu)}] [\sigma^{(\phi,\pi)}] [\eta^{(\mu)}] [\eta^{(\pi)}] [\alpha^{(\mu)}] [\alpha^{(\pi)}] [\rho] [\delta] \prod_{m=1}^{2} [c_{m}] [\tau_{m}]. \end{split}$$

Generalized Pareto/Lomax burn area model

We used the following parameterization of the GPD/Lomax distribution:

$$[y|\sigma,\kappa] = \frac{1}{\sigma} \left(\frac{\kappa y}{\sigma} + 1\right)^{-(\kappa+1)\kappa^{-1}},$$

where κ is a shape parameter and σ is a scale parameter.

$$[\boldsymbol{\beta}, \alpha, \boldsymbol{\phi}, \sigma^{(\phi)}, \eta, \kappa^{(L)}, \boldsymbol{\lambda}, c, \tau \mid \boldsymbol{y}] \propto \prod_{i=1}^{n_{\text{tot}}} [y_i | \boldsymbol{\beta}, \alpha, \phi_{s_i, t_i}, \kappa^{(L)}] \times [\boldsymbol{\phi}_1 | \sigma^{(\phi)}] \prod_{t=2}^{T} [\boldsymbol{\phi}_t | \boldsymbol{\phi}_{t-1}, \sigma^{(\phi)}, \eta] \times \prod_{j=1}^{p} [\beta_j | \lambda_j, c, \tau] [\lambda_j] \times [\alpha] [c] [\tau] [\kappa^{(L)}] [\eta] [\sigma^{(\phi)}]$$

$$= \prod_{i=1}^{n_{\text{tot}}} \text{Lomax}(y_i|\kappa^{(L)}, e^{\alpha + \mathbf{X}_{(s_i,t_i)}\beta + \phi_{s_i,t_i}}) \times \\ \text{Normal}(\phi_1|\mathbf{0}, ((\sigma^{(\phi)})^{-2}(\mathbf{D} - \mathbf{W}))^{-1}) \times \\ \prod_{t=2}^{T} \text{Normal}(\phi_t|\eta\phi_{t-1}, ((\sigma^{(\phi)})^{-2}(\mathbf{D} - \mathbf{W}))^{-1}) \times \\ \prod_{j=1}^{p} \text{Normal}\left(\beta_j|0, \frac{\tau^2 c^2 \lambda_j^2}{c^2 + \tau^2 \lambda_j^2}\right) \times \text{Cauchy}^+(\lambda_j|0, 1) \times \\ \text{Normal}(\alpha|0, 5^2) \times \text{Inv-Gamma}(c^2|2.5, 10) \times \text{Normal}^+(\tau|0, 5^2) \\ \text{Normal}^+(\kappa^{(L)}|0, 5^2) \times \text{Beta}(\eta|1, 1) \times \text{Normal}^+(\sigma^{(\phi)}|0, 1^2). \end{cases}$$

Tapered Pareto burn area model

We used the following parameterization of the tapered Pareto distribution:

$$[y|\kappa,\nu] = \left(\frac{\kappa}{y} + \frac{1}{\nu}\right) \exp(-x/\nu),$$

where κ is a shape parameter and ν a taper parameter.

$$[\boldsymbol{\beta}, \alpha, \boldsymbol{\phi}, \sigma^{(\phi)}, \eta, \nu, \boldsymbol{\lambda}, c, \tau \mid \boldsymbol{y}] \propto \prod_{i=1}^{n_{\text{tot}}} [y_i | \boldsymbol{\beta}, \alpha, \phi_{s_i, t_i}, \nu] \times [\boldsymbol{\phi}_1 | \sigma^{(\phi)}] \prod_{t=2}^{T} [\boldsymbol{\phi}_t | \boldsymbol{\phi}_{t-1}, \sigma^{(\phi)}, \eta] \times \prod_{j=1}^{p} [\beta_j | \lambda_j, c, \tau] [\lambda_j] \times [\alpha] [c] [\tau] [\nu] [\eta] [\sigma^{(\phi)}]$$

$$= \prod_{i=1}^{n_{\text{tot}}} \text{Tapered Pareto}(y_i | e^{\alpha + \mathbf{X}_{(s_i, t_i)} \beta + \phi_{s_i, t_i}}, \nu) \times \\ \text{Normal}(\phi_1 | \mathbf{0}, ((\sigma^{(\phi)})^{-2} (\mathbf{D} - \mathbf{W}))^{-1}) \times \\ \prod_{t=2}^{T} \text{Normal}(\phi_t | \eta \phi_{t-1}, ((\sigma^{(\phi)})^{-2} (\mathbf{D} - \mathbf{W}))^{-1}) \times \\ \prod_{j=1}^{p} \text{Normal}\left(\beta_j | 0, \frac{\tau^2 c^2 \lambda_j^2}{c^2 + \tau^2 \lambda_j^2}\right) \times \text{Cauchy}^+(\lambda_j | 0, 1) \times \\ \text{Normal}(\alpha | 0, 5^2) \times \text{Inv-Gamma}(c^2 | 2.5, 10) \times \text{Normal}^+(\tau | 0, 5^2) \times \\ \text{Cauchy}^+(\nu | 0, 1) \times \text{Beta}(\eta | 1, 1) \times \text{Normal}^+(\sigma^{(\phi)} | 0, 1^2). \end{cases}$$

Lognormal burn area model

We used the following parameterization of the lognormal distribution:

$$[y|\mu,\sigma] = \frac{1}{y} \frac{1}{\sigma \sqrt{2\pi}} \exp\left(-\frac{(\log(y)-\mu)^2}{2\sigma^2}\right),$$

where μ and σ are location and scale parameters, respectively.

The unnormalized posterior density of this model is:

$$\begin{bmatrix} \boldsymbol{\beta}, \alpha, \boldsymbol{\phi}, \sigma^{(\phi)}, \eta, \sigma, \boldsymbol{\lambda}, c, \tau \mid \boldsymbol{y} \end{bmatrix} \propto \\ \prod_{i=1}^{n_{\text{tot}}} [y_i | \boldsymbol{\beta}, \alpha, \phi_{s_i, t_i}, \sigma] \times \\ [\boldsymbol{\phi}_1 | \sigma^{(\phi)}] \prod_{t=2}^{T} [\boldsymbol{\phi}_t | \boldsymbol{\phi}_{t-1}, \sigma^{(\phi)}, \eta] \times \\ \prod_{j=1}^{p} [\beta_j | \lambda_j, c, \tau] [\lambda_j] \times \\ [\alpha] [c] [\tau] [\sigma] [\eta] [\sigma^{(\phi)}] \end{bmatrix}$$

$$= \prod_{i=1}^{n_{\text{tot}}} \text{Lognormal}(y_i | \alpha + \mathbf{X}_{(s_i,t_i)} \boldsymbol{\beta} + \phi_{s_i,t_i}, \sigma) \times \\ \text{Normal}(\boldsymbol{\phi}_1 | \mathbf{0}, ((\sigma^{(\phi)})^{-2} (\mathbf{D} - \mathbf{W}))^{-1}) \times \\ \prod_{t=2}^{T} \text{Normal}(\boldsymbol{\phi}_t | \eta \boldsymbol{\phi}_{t-1}, ((\sigma^{(\phi)})^{-2} (\mathbf{D} - \mathbf{W}))^{-1}) \times \\ \prod_{j=1}^{p} \text{Normal}\left(\beta_j | 0, \frac{\tau^2 c^2 \lambda_j^2}{c^2 + \tau^2 \lambda_j^2}\right) \times \text{Cauchy}^+(\lambda_j | 0, 1) \times \\ \text{Normal}(\alpha | 0, 5^2) \times \text{Inv-Gamma}(c^2 | 2.5, 10) \times \text{Normal}^+(\tau | 0, 5^2) \times \\ \text{Normal}^+(\sigma | 0, 5^2) \times \text{Beta}(\eta | 1, 1) \times \text{Normal}^+(\sigma^{(\phi)} | 0, 1^2).$$

Gamma burn area model

We used the following parameterization of the gamma distribution:

$$[y|\kappa,\sigma] = \frac{1}{\Gamma(\kappa)\sigma^{\kappa}} y^{\kappa-1} \exp(-y/\sigma),$$

where κ is a shape parameter and σ a scale parameter.

$$[\boldsymbol{\phi}, \alpha, \boldsymbol{\phi}, \sigma^{(\phi)}, \eta, \kappa, \boldsymbol{\lambda}, c, \tau \mid \boldsymbol{y}] \propto \prod_{i=1}^{n_{\text{tot}}} [y_i | \beta, \alpha, \phi_{s_i, t_i}, \kappa] \times [\boldsymbol{\phi}_1 | \sigma^{(\phi)}] \prod_{t=2}^{T} [\boldsymbol{\phi}_t | \boldsymbol{\phi}_{t-1}, \sigma^{(\phi)}, \eta] \times \prod_{j=1}^{p} [\beta_j | \lambda_j, c, \tau] [\lambda_j] \times [\alpha] [c] [\tau] [\kappa] [\eta] [\sigma^{(\phi)}]$$

$$= \prod_{i=1}^{n_{\text{tot}}} \text{Gamma}(y_i|\kappa, \kappa/\exp(\alpha + \mathbf{X}_{(s_i,t_i)}\boldsymbol{\beta} + \phi_{s_i,t_i})) \times \\ \text{Normal}(\boldsymbol{\phi}_1|\mathbf{0}, ((\sigma^{(\phi)})^{-2}(\mathbf{D} - \mathbf{W}))^{-1}) \times \\ \prod_{t=2}^{T} \text{Normal}(\boldsymbol{\phi}_t|\eta\boldsymbol{\phi}_{t-1}, ((\sigma^{(\phi)})^{-2}(\mathbf{D} - \mathbf{W}))^{-1}) \times \\ \prod_{j=1}^{p} \text{Normal}\left(\beta_j|0, \frac{\tau^2 c^2 \lambda_j^2}{c^2 + \tau^2 \lambda_j^2}\right) \times \text{Cauchy}^+(\lambda_j|0, 1) \times \\ \text{Normal}(\alpha|0, 5^2) \times \text{Inv-Gamma}(c^2|2.5, 10) \times \text{Normal}^+(\tau|0, 5^2) \times \\ \text{Normal}^+(\kappa|0, 5^2) \times \text{Beta}(\eta|1, 1) \times \text{Normal}^+(\sigma^{(\phi)}|0, 1^2).$$

Weibull burn area model

We used the following parameterization of the Weibull distribution:

$$[y|\kappa,\sigma] = \frac{\kappa}{\sigma} \left(\frac{y}{\sigma}\right)^{\kappa-1} \exp\left(-\left(\frac{y}{\sigma}\right)^{\alpha}\right),$$

where κ is a shape parameter and σ is a scale parameter.

$$[\boldsymbol{\beta}, \alpha, \boldsymbol{\phi}, \sigma^{(\phi)}, \eta, \kappa, \lambda, c, \tau \mid \boldsymbol{y}] \propto \prod_{i=1}^{n_{\text{tot}}} [y_i | \beta, \alpha, \phi_{s_i, t_i}, \kappa] \times [\boldsymbol{\phi}_1 | \sigma^{(\phi)}] \prod_{t=2}^{T} [\boldsymbol{\phi}_t | \boldsymbol{\phi}_{t-1}, \sigma^{(\phi)}, \eta] \times \prod_{j=1}^{p} [\beta_j | \lambda_j, c, \tau] [\lambda_j] \times [\alpha] [c] [\tau] [\kappa] [\eta] [\sigma^{(\phi)}]$$

$$= \prod_{i=1}^{n_{\text{tot}}} \text{Weibull}(y_i|\kappa, \exp(\alpha + \mathbf{X}_{(s_i,t_i)}\boldsymbol{\beta} + \phi_{s_i,t_i})) \times \\ \text{Normal}(\boldsymbol{\phi}_1|\mathbf{0}, ((\sigma^{(\phi)})^{-2}(\mathbf{D} - \mathbf{W}))^{-1}) \times \\ \prod_{t=2}^{T} \text{Normal}(\boldsymbol{\phi}_t|\eta\boldsymbol{\phi}_{t-1}, ((\sigma^{(\phi)})^{-2}(\mathbf{D} - \mathbf{W}))^{-1}) \times \\ \prod_{j=1}^{p} \text{Normal}\left(\beta_j|0, \frac{\tau^2 c^2 \lambda_j^2}{c^2 + \tau^2 \lambda_j^2}\right) \times \text{Cauchy}^+(\lambda_j|0, 1) \times \\ \text{Normal}(\alpha|0, 5^2) \times \text{Inv-Gamma}(c^2|2.5, 10) \times \text{Normal}^+(\tau|0, 5^2) \times \\ \text{Normal}^+(\kappa|0, 5^2) \times \text{Beta}(\eta|1, 1) \times \text{Normal}^+(\sigma^{(\phi)}|0, 1^2).$$