1 Carbon content, carbon fixation yield and dissolved organic carbon

2 release from diverse marine nitrifiers

- 3 Barbara Bayer^{1,2*}, Kelsey McBeain^{1,3}, Craig A. Carlson¹, and Alyson E. Santoro¹
- 4
- ⁵ ¹ Department of Ecology, Evolution and Marine Biology, University of California, Santa Barbara,
- 6 CA, USA
- ⁷²Current address: Department of Microbiology and Ecosystem Science, University of Vienna,
- 8 Vienna, Austria
- ⁹ ³Current address: Department of Oceanography, University of Hawai'i at Manoa, Honolulu, HI,
- 10 USA
- 11 *Correspondence: <u>bbayer@ucsb.edu</u> or <u>barbara.bayer@univie.ac.at</u>
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- 15 Author contribution: BB and AES designed the study. BB and KM conducted laboratory

16 experiments. BB analyzed the data and drafted the initial manuscript. AES and CAC contributed

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24 Abstract

25 Nitrifying microorganisms, including ammonia-oxidizing archaea, ammonia-oxidizing bacteria 26 and nitrite-oxidizing bacteria, are the most abundant chemoautotrophs in the ocean and play an 27 important role in the global carbon cycle by fixing dissolved inorganic carbon (DIC) into biomass. 28 The release of organic compounds by these microbes is less well known but may represent an 29 as-yet unaccounted source of dissolved organic carbon (DOC) available to heterotrophic marine 30 food webs. Here, we provide measurements of cellular carbon and nitrogen guotas, DIC fixation 31 yields and DOC release of ten phylogenetically diverse marine nitrifiers grown in multiple culture 32 conditions. All investigated strains released DOC during growth, making up on average 5-15% of 33 the fixed DIC. Neither substrate concentration nor temperature affected the proportion of fixed 34 DIC released as DOC, but release rates varied between closely related species. Our results also 35 indicate previous studies may have underestimated DIC fixation yields of marine nitrite oxidizers 36 due to partial decoupling of nitrite oxidation from CO₂ fixation, and due to lower observed yields 37 in artificial compared to natural seawater medium. The results of this study provide values for 38 biogeochemical models of the global carbon cycle, and help to further constrain the implications 39 of nitrification-fueled chemoautotrophy for marine food-web functioning and the biological 40 sequestration of carbon in the ocean.

41

42 Introduction

Marine microorganisms play a critical role in the global carbon cycle through their transformations of organic and inorganic carbon constituents. A fraction of the carbon dioxide (CO₂) that is captured by phytoplankton in the surface ocean sinks to depth as dead organic material, supporting a mesopelagic food web of both microbes and higher trophic levels (Hannides et al. 2013; Giering et al. 2014; Choy et al. 2015). Organic matter decomposition in the mesopelagic also releases ammonium, a reduced form of nitrogen that can be used as an energy source by

chemoautotrophic nitrifying archaea and bacteria to fuel dissolved inorganic carbon (DIC) fixation into biomass (Ward 2011). Chemoautotrophic production provides a new, labile, non-sinking source of particulate organic matter to the deep ocean which is otherwise dominated by refractory organic carbon (Reinthaler et al. 2010; Middelburg 2011), supporting a significant fraction of the heterotrophic microbial community in the mesopelagic (Hansman et al. 2009).

54 The main nitrifiers in the ocean are ammonia-oxidizing archaea (AOA), which oxidize 55 ammonia (NH₃) to nitrite (NO₂⁻), and nitrite-oxidizing bacteria (NOB), which further oxidize NO₂⁻ 56 to nitrate (NO₃⁻) (Ward 2011). These two steps are assumed to be tightly coupled, as NO₂⁻ 57 typically does not accumulate in oxic open ocean waters (with the exception of the primary nitrite 58 maximum at the base of the euphotic zone (Lomas and Lipschultz 2006; Santoro et al. 2013)). 59 Despite this tight coupling, AOA are approximately six times more abundant than NOB at a given 60 location and sampling depth (Santoro et al. 2019), possibly owing to their smaller cell size 61 compared to NOB (Watson and Waterbury 1971; Könneke et al. 2005; Santoro and Casciotti 62 2011; Bayer et al. 2016; Mueller et al. 2021), or as a result of the higher theoretical energy yield 63 from ammonia compared to nitrite oxidation (Bock and Wagner 2013). Ammonia-oxidizing 64 bacteria (AOB) are thought to play a minor role in global ocean nitrification due to their overall low 65 abundances (Santoro et al. 2010; Buchwald et al. 2015; Tolar et al. 2016).

66 Despite the known difference in theoretical energy yield, there are many uncertainties 67 regarding the organic carbon yield from ammonia versus nitrite oxidation (hereinafter referred to 68 as DIC fixation yield) and the contribution of both functional groups to chemoautotrophic DIC 69 fixation in the ocean. AOA cultures have recently been shown to release dissolved organic carbon 70 (DOC) during growth (Bayer et al. 2019a), pointing to a potential loss of cellular fixed carbon that 71 is not captured by conventional methods measuring DIC incorporation into biomass. The release 72 of DOC by nitrifiers might represent an as-yet unaccounted source of organic material in the deep 73 ocean potentially fueling the microbial loop, with important implications for the marine carbon 74 cycle. However, it remains unclear if DOC release is a phenomenon only observed under specific

culture conditions restricted to some AOA, or a common feature shared by diverse autotrophicnitrifiers under natural conditions.

Here, we report combined measurements of DIC fixation and DOC release of ten phylogenetically diverse marine nitrifiers comprising two AOA genera (*Nitrosopumilus* and *Ca.* Nitrosopelagicus), one AOB genus (*Nitrosomonas*) and three NOB genera (*Nitrospina, Nitrospira,* and *Nitrococcus*), and further explore the effect of substrate concentration, temperature, and different culture media on these measurements. The results of this study will inform ecological theoretical models to further constrain DIC fixation yields associated with nitrification in order to better understand the dynamics involved in the sequestration of carbon in the ocean.

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85 Methods

86 *Nitrifier culture sources*

87 The AOA cultures used in this study were three axenic *Nitrosopumilus* strains and one 88 Nitrosopelagicus enrichment culture. Ca. Nitrosopelagicus brevis U25 originates from a North 89 Pacific Ocean water sample (Santoro and Casciotti 2011; Carini et al. 2018). The level of 90 enrichment during the time of this study was ~90%. Nitrosopumilus sp. CCS1 is a novel AOA 91 strain isolated from a seawater sample collected from the California Current system in the North 92 Pacific Ocean (Santoro et al. unpublished). Nitrosopumilus adriaticus NF5 (=JCM 32270^T) =NCIMB 15114^T) and *Nitrosopumilus piranensis* D3C (=JCM 32271^T =DSM 106147^T =NCIMB 93 94 15115^T) were isolated from the Northern Adriatic Sea and have been described in detail (Bayer 95 et al. 2016, 2019).

The four axenic NOB strains, *Nitrospina gracilis* Nb-211, *Nitrospina* sp. Nb-3, *Nitrococcus mobilis* Nb-231 and *Nitrospira marina* Nb-295, were obtained from the culture collection of John B. Waterbury and Frederica Valois at the Woods Hole Oceanographic Institution (WHOI). *N. gracilis* Nb-211 was isolated from surface waters of the South Atlantic Ocean (Watson and

Waterbury 1971), *N. mobilis* Nb-231 was isolated from a surface water sample obtained from the South Pacific Ocean (Watson and Waterbury 1971) and *N. marina* Nb-295 was isolated from a water sample collected at a depth of 206 m from the Gulf of Maine in the Atlantic Ocean (Watson et al. 1986). *Nitrospina* sp. Nb-3 was isolated from the Pacific Ocean off the coast of Peru and has not yet been officially described (Watson and Waterbury, unpublished), however, it shares a high 16S rRNA gene sequence similarity with strain 3/211 (Lücker et al. 2013).

AOB strains used in this study, *Nitrosomonas marina* C-25 and *Nitrosomonas* sp. C-15 (also known as strain Nm51, (Koops et al. 1991)), were both obtained from the culture collection of John B. Waterbury and Frederica Valois at WHOI and were revived from 60-year old cryostocks. Strain C-15 was isolated from surface water (1 m depth) of the South Pacific Ocean off the Peruvian continental shelf (Watson and Mandel 1971) and strain C-25 was isolated from surface waters of the South Atlantic Ocean (200 miles off the Amazon River mouth) (Watson and Mandel 1971).

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114 *Culture conditions*

115 Nitrosopumilus adriaticus NF5, Nitrosopumilus piranensis D3C, Nitrosomonas marina C-25 and 116 Nitrosomonas sp. C-15 were grown in HEPES-buffered artificial seawater medium containing 1 117 mM NH₄Cl, and Ca. Nitrosopelagicus brevis U25 was grown in natural seawater medium 118 containing 50 µM NH₄Cl. *Nitrospina gracilis* Nb-211, *Nitrospira marina* Nb-295 and *Nitrococcus* 119 mobilis Nb-231 were grown in artificial seawater medium supplemented with 1 mM NaNO₂. 120 Nitrosopumilus sp. CCS1 and Nitrospina sp. Nb-3 were grown under multiple culture conditions 121 as indicated in the Results and Discussion. All strains were routinely grown in 60 mL 122 polycarbonate bottles (Nalgene) containing 50 mL culture medium, and bottles were incubated at 123 either 15°C or 25°C (with the exception of Ca. Nitrosopelagicus brevis which was always 124 incubated at 22°C) in the dark without agitation.

125 The artificial seawater medium contained 18.54 g L⁻¹ NaCl, 4.7 g L⁻¹ MgSO₄ × 7H₂O, 3.55 g L⁻¹ MgCl₂ × 6H₂O, 1.03 g L⁻¹ CaCl₂ × 2H₂O, 0.51 g L⁻¹ KCl, 0.14 g L⁻¹ NaHCO₃. The natural 126 127 seawater medium consisted of aged seawater collected from the Santa Barbara Channel (approx. 128 10 m depth at, 0.2 µm pore size filtered). Artificial and natural seawater were supplemented with 2.6 mg L⁻¹ K₂HPO₄, 250 µg L⁻¹ FeNaEDTA, 30 µg L⁻¹ H₃BO₃, 20µg L⁻¹ MnCl₂ × 4H₂O, 20 µg L⁻¹ 129 CoCl₂ × 6H₂O, 24 µg L⁻¹NiCl₂ × 6H₂O, 20 µg L⁻¹ CuCl₂ × 2H₂O, 144 µg L⁻¹ ZnSO₄ × 7H₂O, 24 µg 130 131 L^{-1} Na₂MoO₄ × 2H₂O. The pH was adjusted to 7.8-8.0 with NaOH or HCI. Due to the pH decrease 132 associated with ammonia oxidation, culture medium with high initial NH4⁺ concentrations (>250 133 µM) was buffered by addition of 10 mM HEPES (pH 7.8). AOA cultures were supplemented with 50 U L⁻¹ catalase (Sigma-Aldrich, Cat. Nr. C9322) to reduce oxidative stress and NOB cultures 134 were supplemented with 50 ng L⁻¹ cyanocobalamin. To test the effect of reduced and organic 135 136 nitrogen compounds on *Nitrospina sp.* Nb-3, NH₄CI (50 µM) or tryptone (150 mg L⁻¹) were added 137 to the culture medium.

NO₂⁻ concentrations were measured using the Griess-Ilosvay colorimetric method (Strickland and Parsons 1972) and enumeration of cells was performed on an Easy-Cyte flow cytometer (Guava Technologies) following SYBR Green staining as previously described (Bayer et al. 2021).

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143 Cellular carbon and nitrogen content measurements

To determine C : N ratios, ca. 100-500 mL of culture was filtered onto combusted (450°C, 4h) glass fiber filters (Advantec, GF-75, 25mm). Filters were acidified with HCI (10% v/v), dried (60°C, 24 h), and packed into tin capsules prior to being analyzed on a CHN elemental analyzer (Exeter Analytical, CEC 440HA). The instrument was calibrated with acetanilide following manufacturer protocols.

149 Cellular carbon (C) content was calculated using both, CHN elemental analyzer (only for large cells) and ¹⁴C-DIC incorporation measurements (see below), divided by the number of newly 150 151 produced cells. Additionally, C content of the AOA strain Nitrosopumilus sp. CCS1 was calculated 152 from a dilution series of concentrated cells as described in (White et al. 2019). Cells were concentrated using tangential flow filtration (Pellicon) and a dilution series of 1.1 to 5.6 x10¹¹ cells 153 154 L⁻¹ was constructed by resuspending cell concentrates in culture medium (Fig. S1). The total organic 155 C content for each vial of the dilution series was directly measured by high temperature combustion 156 using a modified Shimadzu TOC-V as described in (Carlson et al. 2010). C content per cell was 157 calculated via linear regression of cell counts and elemental content over the dilution series, where 158 the slope of a Model II least squares regression is considered the elemental content per cell (Fig. 159 S1).

160

161 **Combined DIC fixation and DOC release measurements**

162 DIC fixation was measured via the incorporation of [¹⁴C]-bicarbonate as previously described 163 (Herndl et al. 2005) with modifications. [¹⁴C]-bicarbonate (specific activity 56 mCi mmol⁻¹/2.072 x 164 10⁹ Bg mmol⁻¹. Perkin Elmer) was added to 5 mL of culture (between 10-60 µCi were added 165 depending on the activity of the culture). Different incubation times were tested (see Results 166 section) and all consecutive experiments were performed over the entire length of the growth 167 curve. For every culture condition, at least three replicate live samples and one formaldehyde-168 fixed blank (3% v/v) were incubated in temperature-controlled incubators in the dark. Parallel 169 incubations without [¹⁴C]-tracer additions were used to determine cell abundance and nitrite 170 concentration (see above).

Incubations were terminated by adding formaldehyde (3% v/v) to 5 mL of sample. After
30-60 min, every sample was individually filtered onto 25 mm, 0.2 µm pore size polycarbonate
filters (Millipore) and rinsed with 0.5 mL of artificial seawater using a glass filtration set (Millipore).

174 The individual filtrates (5.5 mL per sample) were collected and transferred to scintillation vials to determine the fraction of [¹⁴C]-dissolved organic carbon ([¹⁴C]-DOC). Excess [¹⁴C]-bicarbonate 175 176 from the filters was removed by exposing them to fumes of concentrated HCI (37%) for 24 h. The 177 filters were transferred to scintillation vials and 10 mL of scintillation cocktail (Ultima Gold, Perkin 178 Elmer) was added. The filtrates were acidified to pH ~2 with HCl (25 %) as previously described 179 (Marañón et al. 2004), and filtrates were kept for 24 h in open scintillation vials placed on an orbital 180 shaker before 10 mL scintillation cocktail was added to each vial. Samples were shaken for ca. 181 30 sec and incubated in the dark for at least 24 hours prior to counting the disintegrations per 182 minute (DPM) in a scintillation counter (Beckman Coulter LS6500) for 15 min.

Total radioactivity measurements were performed to verify added [¹⁴C]-bicarbonate concentrations by pipetting 100 μ l of sample into scintillation vials containing 400 μ l betaphenylethylamine (to prevent outgassing of ¹⁴CO₂). Scintillation cocktail was added, vials were shaken for ca. 30 sec and immediately measured in the scintillation counter.

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188 The resulting mean DPM of the samples were corrected for the DPM of the blank, converted into

189 organic carbon fixed over time and corrected for the DIC concentration in the culture media.

190

191 DIC fixation rates were calculated using the following formula:

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193 (DPM_s - DPM_b) × DIC_w/(DPM_{tr} × inc. time)

194

where DPM are the disintegrations per minute measured in the scintillation counter, for the sample (s) and the blank (b). DIC_w denotes the dissolved inorganic carbon concentration in culture

197 medium and DPM tracer (tr) is the DPM for the $[^{14}C]$ -bicarbonate added to the incubations.

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199 DIC concentration measurements

Total alkalinity (TA) of unfixed natural and artificial seawater medium was measured via an opencell endpoint titration using a Mettler-Toledo T5 autotitrator, and pH was measured spectrophotometrically using a Shimadzu UV-1280 UV-VIS spectrophotometer as described previously (Dickson et al. 2007; Hoshijima and Hofmann 2019). Dissolved inorganic carbon (DIC) concentrations were calculated from TA and pH using the CO2SYS software (Pelletier 2007). To calculate DIC concentrations of HEPES-buffered media, TA values were taken from unbuffered artificial seawater medium and the pH was re-measured after adding HEPES.

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208 Calculations of Gibbs free energy (ΔG)

209 The effective Gibbs free energy (ΔG) for ammonia and nitrite oxidation was calculated for the 210 culture conditions in this study using the following formula:

211 $\Delta G = \Delta G^{\circ} + RT \ln Q$

212

where *R* is the ideal gas constant (8.314 J mol⁻¹ K), Q is the reaction quotient, and *T* is the temperature in Kelvin. ΔG° values were obtained from (Amend and Shock 2001).

215

216 Q was calculated based on the following measurements and estimates: NO_2^- concentrations were 217 measured directly (see above); $[NO_3^-]$ and $[NH_4^+]$ were estimated from the decrease or increase 218 in $[NO_2^-]$, respectively; NH_3 concentrations were calculated based on $[NH_4^+]$, pH of the culture 219 medium, and the acid association constant (pKa = 9.4); and O₂ concentrations were estimated to 220 be 235 µM under completely oxic conditions during our incubations. A correction for ionic strength 221 was applied according to (Amend and LaRowe 2019). Calculations can be found in the Supporting 222 Information (Table S1).

223

224 Statistical analyses

225 Pairwise comparisons were performed with a two-sided Mann-Whitnev U Test 226 (pairwise.wilcox.test) using the R software environment (R Core Team 2013). P values were 227 comparisons adjusted for multiple using the Benjamini-Hochberg correction 228 (p.adjust.method="fdr") (Benjamini and Hochberg 1995).

229

230 **Results and Discussion**

231 Elemental composition of cultured nitrifiers

232 We determined the cellular carbon (C) content of cultured isolates of ammonia-oxidizing archaea 233 (AOA), ammonia-oxidizing bacteria (AOB) and nitrite-oxidizing bacteria (NOB) belonging to six 234 different genera. The cellular C contents of AOA were ~11-17 fg C cell⁻¹ (Table 1), which is slightly 235 higher than values reported for natural populations in the deep Atlantic Ocean (~8.39 fg cell⁻¹, 236 (Herndl et al. 2005)) and an AOA enrichment culture from the Baltic Sea (9 fg cell⁻¹, (Berg et al. 237 2014)), but much lower than values reported for AOA from hypoxic shelf waters of the Gulf of 238 Mexico (50 ± 16 fg cell⁻¹, (Kitzinger et al. 2020)). All investigated marine NOB had higher cellular 239 C guotas compared to AOA (Table 1), with *Nitrospina* exhibiting the lowest (~28-55 fg C cell⁻¹) 240 and *Nitrococcus* the highest (~272-1207 fg C cell⁻¹) values (Table 1). The C content of AOA cells 241 remained fairly constant during different growth phases, while C contents of all investigated NOB 242 strains drastically decreased (~40-70%) from early exponential growth to stationary phase, which 243 was supported by the observation of smaller cells in stationary compared to exponentially growing 244 cultures (data not shown). Cell sizes of natural populations of Nitrospinae bacteria have been 245 reported to be 4-fold (Kitzinger et al. 2020) and 50-fold (Pachiadaki et al. 2017) larger than AOA 246 cells, potentially reflecting these variations in cell size and C content during different growth 247 phases.

The molar C : N ratios of all investigated nitrifiers were in the range of 3.4-4.6 : 1 (Table 1), with the exception of previously published values of *Nitrosopumilus maritimus* NAOA6 (Meador et al. 2020) and two AOB strains (Glover 1985). The values observed are lower than values of

heterotrophic marine bacteria (~5 : 1) including *Pelagibacter ubique* (~4.6 : 1) (White et al. 2019), and references therein), with *Nitrospina* cells exhibiting the lowest average C : N ratio (~3.4) of all cultured nitrifiers in our study (Table 1). These low cellular C : N ratios are surprising considering the observation of glycogen storage deposits in cells of *Nitrospina gracilis, Nitrococcus mobilis,* and *Nitrospira marina* (Watson and Waterbury 1971; Watson et al. 1986), as well as polyhydroxbutyrate storage in *Nitrococcus mobilis* (Watson and Waterbury 1971).

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Table 1. Elemental stoichiometry of phylogenetically diverse cultured marine nitrifiers during different growth phases (early exponential, late exponential, stationary) including previously published values. C : N ratios were obtained during exponential growth phase. Cellular C content values are derived from DIC incorporation measurements if not stated otherwise.

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Organism	C : N (mol mol ⁻¹)	Cellular	Ref.		
		Early exponential	Late exponential	Stationary	
Ca. Nitrosopelagicus brevis U25	n.d.	n.d.	10.8	n.d.	this study
Nitrosopumilus sp. CCS1	4.03 ± 0.32	11.8 ± 0.2	12.0 ± 2.0/ 12.5 ^{&}	12.9 ± 2.0/ 16.3 ± 0.2 [§]	this study
Nitrosopumilus adriaticus NF5	3.91	n.d.	16.7 ± 7.5	17.3 ± 2.3	this study, ²
Nitrosopumilus piranensis D3C	3.98	n.d.	16.3	17.2 ± 1.9	this study, ²
Nitrosopumilus maritimus NAOA6	5.8/5.9+	n.d.	n.d.	34 ± 14/ 17 ± 6 ⁺	3
Nitrosomonas sp. C-15	4.31 ± 0.11	n.d.	145.7 ± 11.1	115.2 ± 3.8	this study
Nitrosomonas marina C-25	4.38 ± 0.14	n.d.	302.4 ± 10.0	159.7 ± 13.4	this study
Nitrosomonas marina	5.59-6.11*	241	139	133	1
Nitrosococcus oceani	3.58-4.95*	1115	961	919	1
Nitrospina gracilis Nb-3	3.41 ± 0.05	50.8 ± 3.9	40.1 ± 2.5	28.4 ± 4.6	this study
Nitrospina gracilis Nb-211	3.43 ± 0.18	54.9 ± 4.9 [#]	n.d.	30.4 ± 3.4	this study
Nitrospira marina Nb-295	4.22 ± 0.03	153.5 ± 18.1/ 155.2 ± 6.5 [#]	69.5 ± 7.5	57.8 ± 6.2	this study
Nitrococcus mobilis Nb-231	4.60 ± 0.13	994.6 ± 315.4/ 1206.6 ± 156.1 [#]	442.9 ± 38.0	272.1 ± 60.7	this study
Nitrococcus mobilis	3.07-4.75*	1226	671	384	1

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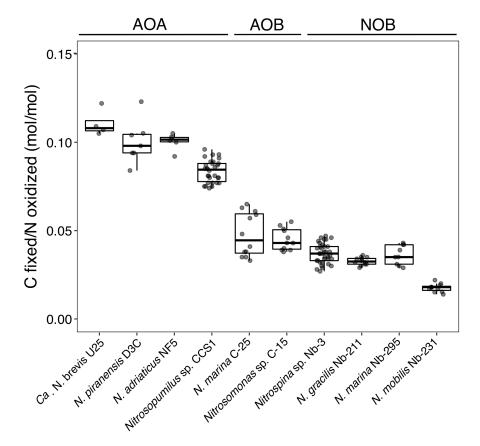
¹ Glover (1985); ² Bayer et al. (2019c); ³ Meador et al. (2020)

- 266 *Range of values obtained during different growth conditions
- 267 [&] Value obtained from TOC dilution series (see Materials and Methods section and Fig. S1)
- 268 [#]Values obtained from CHN elemental analyzer measurements (see Materials and Methods section)
- 269 § Grown in HEPES-buffered medium
- 270 * Values obtained under phosphate-replete and phosphate-deplete conditions (P replete/ P deplete)
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- 272
- 273

DIC fixation yields of marine nitrifiers

274 We conducted combined measurements of DIC fixation, DOC release and ammonia/nitrite 275 oxidation rates of ten nitrifier cultures. Here, we use the term 'DIC fixation yield' to describe the 276 number of moles of inorganic carbon (CO₂ or HCO₃) that are fixed for every mole of N (NH₃ or 277 NO_2^{-}) oxidized, including the proportion that is released/lost as DOC.

278 Marine AOA, including three axenic *Nitrosopumilus* strains and one *Ca*. Nitrosopelagicus 279 enrichment culture, exhibited the highest DIC fixation yields (mean \pm sd= 0.091 \pm 0.012, n=47) in 280 our study, which were on average \sim 2-times higher than those of marine AOB (mean±sd= 0.047) 281 \pm 0.010. *n*=23) (Fig. 1). AOA encode the 3-hydroxypropionate/4-hydroxybutyrate (3-HP/4-HB) 282 cycle for DIC fixation (Walker et al. 2010; Santoro et al. 2015; Bayer et al. 2016), which is 283 suggested to be the most energy-efficient aerobic autotrophic DIC fixation pathway, requiring four 284 moles of ATP to fix two moles of C (Könneke et al. 2014). In contrast, AOB use the Calvin-Benson-285 Bassham (CBB) cycle (Utåker et al. 2002; Stein et al. 2007), which has a higher ATP requirement 286 and an estimated 20% loss of fixed DIC due to the oxygenase side-reaction of ribulose-1,5-287 bisphosphate carboxylase/oxygenase (Berg 2011). DIC fixation yields of two Nitrosopumilus 288 strains were recently reported to be up to ten times higher (0.18-1.2, (Meador et al. 2020)) 289 compared to values in our study and previously published values of Nitrosopumilus adriaticus 290 NF5 (0.1, (Bayer et al. 2019c)) and a *Nitrosarchaeum* enrichment culture (0.1, (Berg et al. 2014)). 291 However, such high values would require unrealistically high ATP yields (up to 2.4 moles ATP per 292 mole NH₃ oxidized) compared to reported estimates of 0.15-0.28 ATP/NH₃ (mol/ mol) (Li et al. 293 2018).



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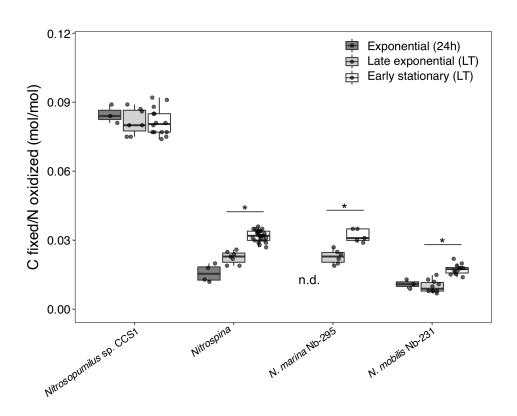
Fig.1 Comparison of DIC fixation yields of ten different phylogenetically diverse marine nitrifiers. Plotted values include both, the fraction of C incorporated into biomass and the fraction of C released as DOC. For NOB, only measurements conducted over the entire length of the growth curve (until stationary phase) are shown (see Fig. 2). Values obtained from cultures grown under different conditions (see Fig. 3 and Fig. S3) are included in this plot.

300

301 DIC fixation yields of marine NOB (*Nitrospina/Nitrospira*: mean \pm sd=0.036 \pm 0.005, *n*=47; 302 *Nitrococcus*: mean \pm sd=0.018 \pm 0.002, *n*=11) were lower compared to those of ammonia oxidizers 303 (Fig. 1). Nitrococcus mobilis, which uses the CBB cycle for DIC fixation (Füssel et al. 2017) had 304 ~2-times lower DIC fixation yields compared to Nitrospina and Nitrospira which use a O_2 -tolerant 305 version of the reverse TCA cycle (Lücker et al. 2010, 2013). Zhang et al. (2020) measured ~1.7-306 times lower DIC fixation yields of Nitrospina gracilis 3/211 and a terrestrial Nitrospira isolate 307 compared to values in our study. We observed that radiotracer incubations conducted over the 308 entire length of the growth curve (until early stationary phase, see Fig. S2) resulted in ~1.4 to 1.7-

times higher DIC fixation yields of NOB compared to incubations conducted until late exponential growth (when NO₂⁻ was completely oxidized) (Fig. 2), suggesting that, in contrast to AOA where ammonia oxidation and DIC fixation were tightly coupled, nitrite oxidation might be partly decoupled from DIC fixation in NOB. While short incubation times (24 h) are typically favored over longer times for environmental measurements to avoid cross-feeding of reaction products, our results indicate that DIC fixation yields of NOB might be underestimated using these established protocols (Fig. 2).

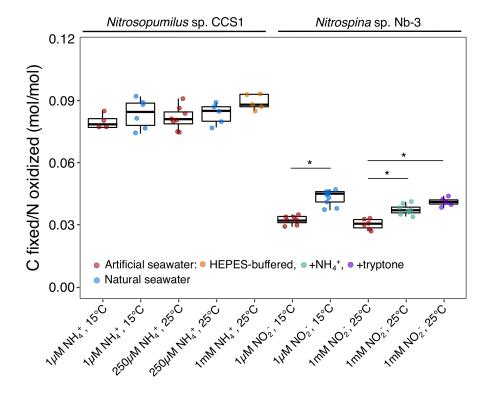




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Fig. 2 Comparison of DIC fixation yields obtained from short-term (24h) radiotracer incubations during exponential growth, and long-term (LT) radiotracer incubations carried out until either late exponential growth or early stationary phase. Measurements of both *Nitrospina* strains (Nb-3 and Nb-211) were combined in this plot. Statistical significance (adj. *p*-value <0.01) of within-condition comparisons are indicated by an asterisk (*). Statistical results of all pairwise comparisons are reported in Table S2. Representative growth curves can be found in the Supporting Information (Fig. S2).

326 We further explored the effect of multiple culture conditions, including environmentally 327 relevant conditions of low substrate concentrations (1 µM) and low temperature (15°C), on DIC 328 fixation yields of Nitrosopumilus sp. CCS1 and Nitrospina sp. Nb-3. We observed that Nitrospina 329 sp. Nb-3 was \sim 1.4-times more efficient in converting energy to growth when grown in natural 330 seawater compared to artificial seawater medium, which was not observed for *Nitrosopumilus* sp. 331 CCS1 (Fig. 3). We hypothesize that reduced N compounds present in natural seawater 332 (ammonium and/or organic N compounds) might be responsible for the observed differences due 333 to the high metabolic costs (6 reducing equivalents) associated with assimilatory NO_2^- reduction 334 (Einsle et al. 2002). Additions of ammonium or tryptone to artificial seawater medium likewise 335 resulted in significantly higher DIC fixation yields (Fig. 3, Fig. S3), corroborating this hypothesis. 336 Environmental populations of *Nitrospinae* have previously been shown to favor ammonium and 337 the organic N sources urea and cyanate over nitrite (Kitzinger et al. 2020). Our data suggest that 338 in addition to urea and cvanate, marine NOB can assimilate more complex organic N sources 339 such as peptides and/or amino acids thereby saving energy that can instead be invested in C 340 assimilation. The two most recent estimates for global ocean DIC fixation by NOB differ by one 341 order of magnitude (Pachiadaki et al. 2017; Zhang et al. 2020) (Table S3), potentially also 342 reflecting some of these uncertainties. Furthermore, we observed slightly higher DIC fixation 343 vields of *Nitrosopumilus* sp. CCS1 in HEPES-buffered artificial seawater compared to unbuffered 344 culture medium (Fig. 3), which coincided with higher cellular C quota (Table 1). While we cannot 345 explain these observations, the differences in DIC fixation yield did not seem to be caused by 346 variations in pH, which remained constant in unbuffered culture medium containing low substrate 347 concentrations (1 μ M NH₄⁺).



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Fig. 3 DIC fixation yields of *Nitrosopumilus* sp. CCS1 and *Nitrospina* sp. Nb-3 under different culture conditions (substrate concentrations: 1 μ M, 250 μ M, 1mM; temperature: 15°C, 25°C) and culture media (natural seawater, artificial seawater, HEPES-buffered artificial seawater). Plotted values include both, the fraction of C incorporated into biomass and the fraction of C released as DOC. Ammonium (50 μ M) or tryptone (150 mg L⁻¹) served as additional, reduced nitrogen source for *Nitrospina* sp. Nb-3. Statistical significance (adj. *p*-value <0.01) of within-condition comparisons are indicated by an asterisk (*). Statistical results of pairwise comparisons are reported in Table S2.

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357 The theoretical Gibbs free energy release (ΔG) for conditions in our study was 3.6-times 358 higher for ammonia compared to nitrite oxidation (Table 2), yet DIC fixation yields of 359 Nitrosopumilus sp. CCS1 and Ca. Nitrosopelagicus brevis U25 (Table 2) were only 2 to 2.6-times 360 higher compared to *Nitrospina* sp. Nb-3. Similar observations were made by Kitzinger et al (2020) 361 who reported that Nitrospinae bacteria in low O₂ waters of the Gulf of Mexico are more efficient 362 in translating energy gained from nitrite to C assimilation than AOA are in translating energy 363 gained from ammonia oxidation. In addition to thermodynamics and the efficiency of the DIC 364 fixation pathway itself, additional factors can contribute to realized energy yields, including the 365 requirement of four out of six generated electrons by ammonia monooxygenase to reduce

366 molecular oxygen in ammonia oxidizers (Stahl and de la Torre 2012; Caranto and Lancaster 367 2018). When considering that a maximum of 53.8% of the energy released from catabolism are 368 available to ammonia oxidizers for growth (González-Cabaleiro et al. 2019), AOA appear to have 369 slightly higher DIC fixation efficiencies compared to NOB encoding the rTCA cycle (Table 2). 370 While oxygen protection likely increases the energy demands of the rTCA cycle (Berg 2011), our 371 results indicate that the cycle is also highly efficient under oxic conditions that are found in most 372 regions of the global ocean.

373

374**Table 2.** Thermodynamic considerations and comparison of DIC fixation efficiencies and biomass yields of375marine AOA and NOB grown under environmentally relevant conditions (substrate concentration: 1 μ M;376temperature: 15°C) in artificial and natural seawater medium. Gibbs free energy calculations for NH₃377oxidation and NO₂⁻ oxidation can be found in Table S1.

378

	Ca. Nitrosopelagicus U25§	§ Nitrosopumilus sp. CCS1		Nitrospina sp. Nb-3	
Culture medium	Natural seawater	Artificial seawater	Natural seawater	Artificial seawater	Natural seawater
Gibbs free energy (kJ mol ⁻¹)	280 / 151*	276 / 149*	276 / 149*	77	77
DIC fixation yield (mol mol ⁻¹)	0.111 ± 0.008	0.080 ± 0.004	0.085 ± 0.008	0.032 ± 0.002	0.043 ± 0.004
DIC fixation efficiency (μmol C kJ ⁻¹)	396 ± 29 / 735 ± 53*	290 ± 15 / 537 ± 27*	308 ± 29 / 570 ± 54*	416 ± 26	558 ± 52
Biomass yield ^{&} (gBio gN ⁻¹)	0.187 ± 0.019	0.135 ± 0.010	0.143 ± 0.019	0.056 ± 0.005	0.076 ± 0.010

- 379
- 380

381 [§]Ca. Nitrosopelagicus U25 was grown at 22°C with initial substrate concentrations of 50μM.

382 [&]The average chemical formula of bacterial biomass (CH_{1.7}O_{0.4}N_{0.2}, (Popovic 2019)) was adjusted using the C:N ratios

383 from Table 1 (AOA: CH_{1.7}O_{0.4}N_{0.25}; *Nitrospina*: CH_{1.7}O_{0.4}N_{0.29}).

*When considering 53.8% of the energy released is available for growth according to González-Cabaleiro et al. 2019.
 385

Multiple studies have used estimates of DIC fixation yields to infer DIC fixation rates associated with nitrification in diverse marine and estuarine environments (Dore and Karl 1996; Lam et al. 2004; Wuchter et al. 2006; Middelburg 2011; Lee et al. 2015), and a value of 0.1 for

389 archaeal ammonia oxidation has widely been used in the literature (Wuchter et al. 2006: 390 Reinthaler et al. 2010; Middelburg 2011) without direct experimental evidence. Previous 391 measurements of DIC fixation yields were mainly derived from cultures of ammonia and nitrite 392 oxidizers that are not representative for the majority of nitrifiers found in marine environments and 393 were highly variable (AOB: 0.033-0.130; NOB: 0.013-0.031; (Prosser 1990) and references 394 therein). The variations in DIC fixation yields we observe for marine nitrifiers across different 395 species and culture conditions are comparably low within AOA (mean \pm sd=0.091 \pm 0.012; *n*=47) 396 and Nitrospina/Nitrospira (mean \pm sd=0.036 \pm 0.005; n=56), suggesting that these values are more 397 constrained than previous estimates and particularly useful for modelling approaches in marine 398 systems.

399

400 **DOC** release by chemolithoautotophs

401 We measured DOC release rates of ten nitrifier cultures and tested how different culture 402 conditions affected the amount of DOC released in proportion to the amount of fixed DIC. All 403 investigated strains released DOC during exponential growth, and DOC release ceased when 404 cultures reached stationary phase (as determined by comparing the total amount of released DOC 405 until late exponential vs stationary phase, see Fig. S4), suggesting that DOC release is a feature 406 of metabolically active nitrifiers. This is in agreement with earlier observations of amino acid 407 release by exponentially growing Nitrosopumilus cells (Bayer et al. 2019a). The amount of 408 chemoautotrophically fixed DIC that was released as DOC by nitrifiers made up on average ~5-409 15% (Fig. 4a). This is within the range observed for phytoplankton, which released 2-10% and 4-410 42% of their photosynthetically fixed DIC in culture and environmental studies, respectively 411 (Carlson 2002), and references therein).

DOC release varied between closely related species (Fig. 4a). *Nitrosopumilus piranensis* released more DOC compared to the two other investigated *Nitrosopumilus* species, which is in agreement with (Bayer et al. 2019b) who reported higher amino acid release rates of *N. piranensis*

415 compared to *N. adriaticus*. Differences in the amount of released DOC have also been recently 416 reported between the closely related aquarium strain *Nitrosopumilus maritimus* SCM1 (9-19% of 417 fixed DIC) and the environmental strain *Nitrosopumilus maritimus* NAOA6 (5% of fixed DIC) 418 (Meador et al. 2020). Within NOB, *Nitrospina* sp. Nb-3 consistently released less DOC compared 419 to *N. gracilis* Nb-211 and the two phylogenetically more distantly related species *N. marina* and 420 *N. mobilis*.

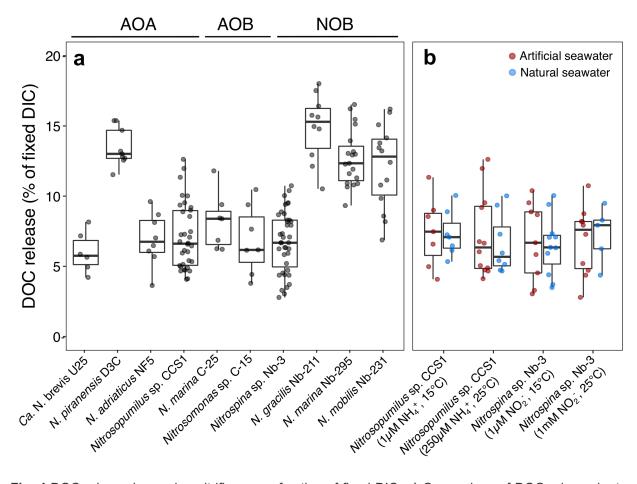


Fig. 4 DOC release by marine nitrifiers as a fraction of fixed DIC. a) Comparison of DOC release by ten different phylogenetically diverse marine nitrifiers. Values obtained from cultures grown under different conditions (see panel b) are included in this plot. DOC release by *Ca*. N. brevis might be underestimated due to the presence of heterotrophic bacteria that could take up some of the released DOC. b) Comparison of DOC release by *Nitrosopumilus* sp. CCS1 and *Nitrospina* sp. Nb-3 grown under different culture conditions (substrate concentrations: 1µM, 250µM, 1mM; temperature: 15°C, 25°C) in artificial or natural seawater medium. Statistical results of all pairwise comparisons are reported in Table S2.

430 The fraction of released DOC remained constant across different culture conditions 431 including environmentally relevant conditions of low substrate concentration (1 µM) and at low 432 temperature (15°C) in natural seawater (Fig. 4b). This suggests that DOC release is not an artifact 433 of unrealistic culture conditions but likely a feature exhibited by nitrifier populations in the 434 environment. While the composition of DOM released by bacterial nitrifiers is currently unknown, 435 a fraction of the DOM released by AOA has been shown to consist of labile compounds, such as 436 amino acids, thymidine and B vitamins, that can limit microbial heterotrophic activity in open ocean 437 waters (Bayer et al. 2019a).

438

439 **Conclusions**

Our results suggest that DIC fixation yields of marine NOB might be underestimated by conventional short-term tracer incubations, due to a partial decoupling between NO_2^- oxidation and C assimilation. Additionally, DIC fixation yields of *Nitrospina* were positively affected by the presence of ammonium or complex organic N compounds, which might influence metabolic interactions with ammonia oxidizers and/or heterotrophic prokaryotes in the environment, suggesting a potentially underappreciated role for competition in the N cycle (Santoro 2016).

446 DIC fixation yields of marine nitrifiers obtained in our study will help to further constrain 447 the relationship between C and N fluxes in the nitrification process and inform theoretical models 448 about how to connect observations at microscale to regional and global scales. Using a mean 449 global value of organic C export from the euphotic zone of ~6 Pg C yr^{-1} (Siegel et al. 2014) and 450 a mean C:N ratio of sinking marine particles (at the surface) of ~7.1 (Schneider et al. 2003), we estimate that the resulting global ocean organic N export of ~0.85 Pg yr⁻¹ could fuel up to 0.13 Pg 451 452 C y⁻¹ of chemoautotrophic DIC fixation (0.094 Pg C y⁻¹ by AOA and 0.037 Pg C y⁻¹ by NOB) in the 453 dark ocean, which is lower than previous estimates (0.15-1.4 Pg C y⁻¹, see Table S3 and 454 references therein). Furthermore, we show that nitrifiers release significant amounts of DOC

- 455 under environmentally relevant conditions, equating to fluxes of 0.006-0.02 Pg C y⁻¹ of fixed DIC
- 456 released as DOC. Elucidating the lability and fate of the DOM released by nitrifiers will be crucial
- 457 to understand its implications for the marine carbon cycle.
- 458

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- 473
- 474 The authors declare no conflict of interest.
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477 **References**

- 478 Amend, J. P., and D. E. LaRowe. 2019. Minireview: demystifying microbial reaction energetics.
- 479 Environmental Microbiology **21**: 3539–3547. doi:10.1111/1462-2920.14778

- Amend, J. P., and E. L. Shock. 2001. Energetics of overall metabolic reactions of thermophilic
 and hyperthermophilic Archaea and Bacteria. FEMS Microbiology Reviews 25: 175–243.
 doi:10.1016/S0168-6445(00)00062-0
- 483 Bayer, B., R. L. Hansman, M. J. Bittner, B. E. Noriega-Ortega, J. Niggemann, T. Dittmar, and G.
- 484 J. Herndl. 2019a. Ammonia-oxidizing archaea release a suite of organic compounds 485 potentially fueling prokaryotic heterotrophy in the ocean. Environmental Microbiology **21**:
- 486 4062–4075. doi:10.1111/1462-2920.14755
- 487 Bayer, B., C. Pelikan, M. J. Bittner, T. Reinthaler, M. Könneke, G. J. Herndl, and P. Offre. 2019b.
- 488 Proteomic Response of Three Marine Ammonia-Oxidizing Archaea to Hydrogen Peroxide
- 489 and Their Metabolic Interactions with a Heterotrophic Alphaproteobacterium. mSystems **4**:
- 490 e00181-19. doi:10.1128/mSystems.00181-19
- 491 Bayer, B., M. A. Saito, M. R. McIlvin, S. Lücker, D. M. Moran, T. S. Lankiewicz, C. L. Dupont, and
- 492 A. E. Santoro. 2021. Metabolic versatility of the nitrite-oxidizing bacterium Nitrospira marina 493 and its proteomic response to oxygen-limited conditions. ISME Journal **15**: 1025–1039.
- 494 doi:10.1038/s41396-020-00828-3
- 495 Bayer, B., J. Vojvoda, P. Offre, and others. 2016. Physiological and genomic characterization of
- 496 two novel marine thaumarchaeal strains indicates niche differentiation. ISME Journal **10**:
- 497 1051–1063. doi:10.1038/ismej.2015.200
- Bayer, B., J. Vojvoda, T. Reinthaler, C. Reyes, M. Pinto, and G. J. Herndl. 2019c. Nitrosopumilus
 adriaticus sp. nov. and Nitrosopumilus piranensis sp. nov., two ammonia-oxidizing archaea
 from the Adriatic Sea and members of the class Nitrosophaeria. International Journal of
- 501 Systematic and Evolutionary Microbiology **7**: 1892–1902. doi:10.1099/ijsem.0.003360
- 502 Benjamini, Y., and Y. Hochberg. 1995. Controlling the False Discovery Rate: A Practical and
- 503 Powerful Approach to Multiple Testing. Journal of the Royal Statistical Society: Series B
- 504 (Methodological) **57**: 289–300. doi:https://doi.org/10.1111/j.2517-6161.1995.tb02031.x

505 Berg, C., L. Listmann, V. Vandieken, A. Vogts, and K. Jürgens. 2014. Chemoautotrophic growth

506 of ammonia-oxidizing Thaumarchaeota enriched from a pelagic redox gradient in the Baltic

507 Sea. Frontiers in Microbiology **5**: 786. doi:10.3389/fmicb.2014.00786

- 508 Berg, I. A. 2011. Ecological aspects of the distribution of different autotrophic CO2 fixation
- 509 pathways. Applied and Environmental Microbiology **77**: 1925–1936. 510 doi:10.1128/AEM.02473-10
- 511 Bock, E., and M. Wagner. 2013. Oxidation of Inorganic Nitrogen Compounds as an Energy 512 Source, p. 83–118. *In* E. Rosenberg, E.F. DeLong, S. Lory, E. Stackebrandt, and F.
- 513 Thompson [eds.], The Prokaryotes: Prokaryotic Physiology and Biochemistry. Springer 514 Berlin Heidelberg.
- Buchwald, C., A. E. Santoro, R. H. R. Stanley, and K. L. Casciotti. 2015. Nitrogen cycling in the
 secondary nitrite maximum of the eastern tropical North Pacific off Costa Rica. Global
 Biogeochemical Cycles 29: 2061–2081. doi:10.1002/2015GB005187
- Caranto, J. D., and K. M. Lancaster. 2018. Nitric oxide is an obligate bacterial nitrification
 intermediate produced by hydroxylamine oxidoreductase. Proceedings of the National
 Academy of Sciences 115: E8325–E8325. doi:10.1073/pnas.1812827115
- 521 Carlson, C. A. 2002. Chapter 4 Production and Removal Processes, p. 91–151. *In* D.A. Hansell
- and C.A. Carlson [eds.], Biogeochemistry of Marine Dissolved Organic Matter. AcademicPress.
- Carlson, C. A., D. A. Hansell, N. B. Nelson, D. A. Siegel, W. M. Smethie, S. Khatiwala, M. M.
 Meyers, and E. Halewood. 2010. Dissolved organic carbon export and subsequent
 remineralization in the mesopelagic and bathypelagic realms of the North Atlantic basin.
 Deep-Sea Research Part II: Topical Studies in Oceanography 57: 1433–1445.
 doi:10.1016/j.dsr2.2010.02.013
- 529 Choy, C. A., B. N. Popp, C. C. S. Hannides, and J. C. Drazen. 2015. Trophic structure and food 530 resources of epipelagic and mesopelagic fishes in the north pacific subtropical Gyre

- 531 ecosystem inferred from nitrogen isotopic compositions. Limnology and Oceanography **60**:
- 532 1156–1171. doi:10.1002/lno.10085
- 533 Dickson, A. G., C. L. Sabine, and J. R. Christian, eds. 2007. Guide to best practices for ocean
- 534 CO₂ measurements, PICES Special Publication 3.
- 535 Dore, J. E., and D. M. Karl. 1996. Nitrification in the euphotic zone as a source for nitrite, nitrate,
- and nitrous oxide at Station ALOHA. Limnology and Oceanography **41**: 1619–1628.
- 537 doi:10.4319/lo.1996.41.8.1619
- 538 Einsle, O., A. Messerschmidt, R. Huber, P. M. H. Kroneck, and F. Neese. 2002. Mechanism of
- 539 the Six-Electron Reduction of Nitrite to Ammonia by Cytochrome c Nitrite Reductase. Journal
- 540 of the American Chemical Society **124**: 11737–11745. doi:10.1021/ja0206487
- 541 Füssel, J., S. Lücker, P. Yilmaz, and others. 2017. Adaptability as the key to success for the 542 ubiquitous marine nitrite oxidizer Nitrococcus. Science Advances **3**: 2–10. 543 doi:10.1126/sciadv.1700807
- Giering, S. L. C., R. Sanders, R. S. Lampitt, and others. 2014. Reconciliation of the carbon budget
 in the ocean's twilight zone. Nature **507**: 480–483. doi:10.1038/nature13123
- 546 Glover, H. E. 1985. The relationship between inorganic nitrogen oxidation and organic carbon
- 547 production in batch and chemostat cultures of marine nitrifying bacteria. Archives of
 548 Microbiology 142: 45–50. doi:10.1007/BF00409235
- 549 González-Cabaleiro, R., T. P. Curtis, and I. D. Ofiţeru. 2019. Bioenergetics analysis of ammonia-
- 550 oxidizing bacteria and the estimation of their maximum growth yield. Water Research **154**:
- 551 238–245. doi:10.1016/j.watres.2019.01.054
- Hannides, C. C. S., B. N. Popp, C. Anela Choy, and J. C. Drazen. 2013. Midwater zooplankton
- 553 and suspended particle dynamics in the North Pacific Subtropical Gyre: A stable isotope
- 554 perspective. Limnology and Oceanography **58**: 1931–1946. doi:10.4319/lo.2013.58.6.1931

- 555 Hansman, R. L., S. Griffin, J. T. Watson, E. R. M. Druffel, A. E. Ingalls, A. Pearson, and L. I.
- 556 Aluwihare. 2009. The radiocarbon signature of microorganisms in the mesopelagic ocean.

557 Proc. Natl. Acad. Sci. USA **106**: 6513–6518. doi:10.1073/pnas.0810871106

- Herndl, G. J., T. Reinthaler, E. Teira, H. van Aken, C. Veth, A. Pernthaler, and J. Pernthaler. 2005.
- 559 Contribution of Archaea to total prokaryotic production in the deep Atlantic Ocean. Applied 560 and environmental microbiology **71**: 2303–2309. doi:10.1128/AEM.71.5.2303
- Hoshijima, U., and G. E. Hofmann. 2019. Variability of seawater chemistry in a kelp forest
 environment is linked to in situ transgenerational effects in the purple sea urchin,
 Strongylocentrotus purpuratus. Frontiers in Marine Science 6: 62.
 doi:10.3389/fmars.2019.00062
- Kitzinger, K., H. K. Marchant, L. A. Bristow, and others. 2020. Single cell analyses reveal
 contrasting life strategies of the two main nitrifiers in the ocean. Nature Communications 11.
 doi:10.1038/s41467-020-14542-3
- 568 Könneke, M., A. E. Bernhard, J. R. de la Torre, C. B. Walker, J. B. Waterbury, and D. A. Stahl.
- 569 2005. Isolation of an autotrophic ammonia-oxidizing marine archaeon. Nature 437: 543–546.
 570 doi:10.1038/nature03911
- Könneke, M., D. M. Schubert, P. C. Brown, and others. 2014. Ammonia-oxidizing archaea use
 the most energy-efficient aerobic pathway for CO2 fixation. Proceedings of the National
 Academy of Sciences of the United States of America 111: 8239–8244.
 doi:10.1073/pnas.1402028111
- Lam, P., J. P. Cowen, and R. D. Jones. 2004. Autotrophic ammonia oxidation in a deep-sea
 hydrothermal plume. FEMS Microbiology Ecology 47: 191–206. doi:10.1016/S01686496(03)00256-3
- Lee, D. Y., M. S. Owens, M. Doherty, E. M. Eggleston, I. Hewson, B. C. Crump, and J. C. Cornwell.
 2015. The Effects of Oxygen Transition on Community Respiration and Potential

- 580 Chemoautotrophic Production in a Seasonally Stratified Anoxic Estuary. Estuaries and 581 Coasts **38**: 104–117. doi:10.1007/s12237-014-9803-8
- 582 Li, F., W. Xie, Q. Yuan, and others. 2018. Genome-scale metabolic model analysis indicates low
- 583 energy production efficiency in marine ammonia-oxidizing archaea. AMB Express 8: 0–11.
- 584 doi:10.1186/s13568-018-0635-y
- Lomas, M. W., and F. Lipschultz. 2006. Forming the primary nitrite maximum: Nitrifiers or phytoplankton? Limnology and Oceanography **51**: 2453–2467. doi:10.4319/lo.2006.51.5.2453
- 588 Lücker, S., B. Nowka, T. Rattei, E. Spieck, and H. Daims. 2013. The genome of Nitrospina gracilis
- illuminates the metabolism and evolution of the major marine nitrite oxidizer. Frontiers in
 Microbiology 4: 27. doi:10.3389/fmicb.2013.00027
- Lücker, S., M. Wagner, F. Maixner, and others. 2010. A Nitrospira metagenome illuminates the
 physiology and evolution of globally important nitrite-oxidizing bacteria. Proceedings of the
 National Academy of Sciences of the United States of America **107**: 13479–13484.
- 594 doi:10.1073/pnas.1003860107
- Marañón, E., P. Cermeño, E. Fernández, J. Rodríguez, and L. Zabala. 2004. Significance and
 mechanisms of photosynthetic production of dissolved organic carbon in coastal eutrophic
- 597 ecosystems. Limnology and Oceanography **49**: 1652–1666. doi:10.4319/lo.2004.49.5.1652
- 598 Meador, T. B., N. Schoffelen, T. G. Ferdelman, O. Rebello, A. Khachikyan, and M. Könneke.
- 599 2020. Carbon recycling efficiency and phosphate turnover by marine nitrifying archaea.
 600 Science Advances 6. doi:10.1126/sciadv.aba1799
- Middelburg, J. J. 2011. Chemoautotrophy in the ocean. Geophysical Research Letters 38: 94–
 97. doi:10.1029/2011GL049725
- Mueller, A. J., M. Y. Jung, C. R. Strachan, C. W. Herbold, R. H. Kirkegaard, M. Wagner, and H.
- Daims. 2021. Genomic and kinetic analysis of novel Nitrospinae enriched by cell sorting.
- 605 ISME Journal **15**: 732–745. doi:10.1038/s41396-020-00809-6

Pachiadaki, M. G., E. Sintes, K. Bergauer, and others. 2017. Major role of nitrite-oxidizing bacteria
in dark ocean carbon fixation. Science **358**: 1046–1051. doi:10.1126/science.aan8260

608 Pelletier, G. L. E. W. D. 2007. CO2SYS.xls: A calculator for the CO2 System in Seawater for

- 609 Microsoft Excel/VBA . Washington State Department for Ecology, Olympia, WA; Brookhaven
- 610 National Laboratory, Upton, NY.
- 611 Popovic, M. 2019. Thermodynamic properties of microorganisms: determination and analysis of
- 612 enthalpy, entropy, and Gibbs free energy of biomass, cells and colonies of 32 microorganism
- 613 species. Heliyon **5**. doi:10.1016/j.heliyon.2019.e01950
- 614 Prosser, J. I. 1990. Autotrophic Nitrification in Bacteria, p. 125–181. *In* A.H. Rose and D.W.
- 615 Tempest [eds.]. Academic Press.
- R Core Team. 2013. R: A language and environment for statistical computing. R Foundation for
 Statistical Computing, Vienna, Austria. http://www.R-project.org.
- 618 Reinthaler, T., H. M. van Aken, and G. J. Herndl. 2010. Major contribution of autotrophy to
- 619 microbial carbon cycling in the deep North Atlantic's interior. Deep-Sea Research Part II:
- 620 Topical Studies in Oceanography **57**: 1572–1580. doi:10.1016/j.dsr2.2010.02.023
- 621 Santoro, A. E. 2016. The do-it-all nitrifier. Science **351**: 342–343.
- 622 Santoro, A. E., and K. L. Casciotti. 2011. Enrichment and characterization of ammonia-oxidizing
- archaea from the open ocean: Phylogeny, physiology and stable isotope fractionation. ISME
 Journal 5: 1796–1808. doi:10.1038/ismej.2011.58
- Santoro, A. E., K. L. Casciotti, and C. A. Francis. 2010. Activity, abundance and diversity of
 nitrifying archaea in the central California Current. Environmental Microbiology 12: 1989–
- 627 2006. doi:10.1111/j.1462-2920.2010.02205.x
- 628 Santoro, A. E., C. L. Dupont, R. A. Richter, and others. 2015. Genomic and proteomic 629 characterization of "Candidatus Nitrosopelagicus brevis": An ammonia-oxidizing archaeon 630 from the open ocean. Proc. Natl. Acad. Sci. USA 112: 1173–1178. 631 doi:10.1073/pnas.1416223112

- 632 Santoro, A. E., R. A. Richter, and C. L. Dupont. 2019. Planktonic Marine Archaea. Annual Review
- 633 of Marine Science **11**: 131–158. doi:10.1146/annurev-marine-121916-063141
- 634 Santoro, A. E., C. M. Sakamoto, J. M. Smith, and others. 2013. Measurements of nitrite production
- 635 in and around the primary nitrite maximum in the central California Current. Biogeosciences
- 636 **10**: 7395–7410. doi:10.5194/bg-10-7395-2013
- 637 Schneider, B., R. Schlitzer, G. Fischer, and E. M. Nöthig. 2003. Depth-dependent elemental
- 638 compositions of particulate organic matter (POM) in the ocean. Global Biogeochemical
- 639 Cycles **17**. doi:10.1029/2002gb001871
- 640 Siegel, D. A., K. O. Buesseler, S. C. Doney, S. F. Sailley, M. J. Behrenfeld, and P. W. Boyd. 2014.
- 641 Global assessment of ocean carbon export by combining satellite observations and food-
- web models. Global Biogeochemical Cycles 28: 181–196.
 doi:10.1002/2013GB004743.Received
- 644 Stahl, D. A., and J. R. de la Torre. 2012. Physiology and Diversity of Ammonia-Oxidizing Archaea.
- 645 Annual Review of Microbiology **66**: 83–101. doi:10.1146/annurev-micro-092611-150128
- 546 Stein, L. Y., D. J. Arp, P. M. Berube, and others. 2007. Whole-genome analysis of the ammonia-
- 647 oxidizing bacterium, Nitrosomonas eutropha C91: Implications for niche adaptation.
- 648 Environmental Microbiology **9**: 2993–3007. doi:10.1111/j.1462-2920.2007.01409.x
- Strickland, J. D. H., and T. R. Parsons. 1972. A Partical Handbook of Seawater Analysis. Fish.
 Res. Bd. Can., Bull. No. 167.
- Tolar, B. B., M. J. Ross, N. J. Wallsgrove, Q. Liu, L. I. Aluwihare, B. N. Popp, and J. T. Hollibaugh.
- 652 2016. Contribution of ammonia oxidation to chemoautotrophy in Antarctic coastal waters.
- 653 ISME Journal **10**: 2605–2619. doi:10.1038/ismej.2016.61
- Utåker, J. B., K. Andersen, Å. Aakra, B. Moen, and I. F. Nes. 2002. Phylogeny and functional
 expression of ribulose 1,5-bisphosphate carboxylase/oxygenase from the autotrophic
- ammonia-oxidizing bacterium Nitrosospira sp. isolate 40KI. Journal of Bacteriology **184**:
- 657 468–478. doi:10.1128/JB.184.2.468-478.2002

- Walker, C. B., J. R. de la Torre, M. G. Klotz, and others. 2010. Nitrosopumilus maritimus genome
- 659 reveals unique mechanisms for nitrification and autotrophy in globally distributed marine
- 660 crenarchaea. Proceedings of the National Academy of Sciences **107**: 8818–8823.
- 661 doi:10.1073/pnas.0913533107
- Ward, B. B. 2011. Nitrification in the Ocean, p. 325–346. *In* B.B. Ward, D.J. Arp, and M.G. Klotz
 Ieds.l. Nitrification. ASM Press.
- 664 Watson, S. W., E. Bock, F. W. Valois, J. B. Waterbury, and U. Schlosser. 1986. Nitrospira marina
- gen. nov. sp. nov.: a chemolithotrophic nitrite-oxidizing bacterium. Archives of Microbiology
- 666 **144**: 1–7. doi:10.1007/BF00454947
- 667 Watson, S. W., and J. B. Waterbury. 1971. Characteristics of two marine nitrite oxidizing bacteria,
- Nitrospina gracilis nov. gen. nov. sp. and Nitrococcus mobilis nov. gen. nov. sp. Archiv für
 Mikrobiologie 77: 203–230. doi:10.1007/BF00408114
- 670 White, A. E., S. J. Giovannoni, Y. Zhao, K. Vergin, and C. A. Carlson. 2019. Elemental content
- and stoichiometry of SAR11 chemoheterotrophic marine bacteria. Limnology &
 Oceanography Letters 4: 44–51. doi:10.1002/lol2.10103
- Wuchter, C., B. Abbas, M. J. L. Coolen, and others. 2006. Archaeal nitrification in the ocean. Proc.
- 674 Natl. Acad. Sci. USA **103**: 12317–12322.
- Zhang, Y., W. Qin, L. Hou, and others. 2020. Nitrifier adaptation to low energy flux controls
 inventory of reduced nitrogen in the dark ocean. Proceedings of the National Academy of
 Sciences of the United States of America **117**: 4823–4830. doi:10.1073/pnas.1912367117