Phosphorus availability and arbuscular mycorrhizal fungi limit soil C cycling and influence plant responses to elevated CO₂ conditions.

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9 Abstract

Enhanced soil organic matter (SOM) decomposition and organic phosphorus (P) cycling may
help sustain plant productivity under elevated CO₂ (eCO₂) and P-limiting conditions. P-

12 acquisition by arbuscular mycorrhizal (AM) fungi and their impacts on SOM decomposition

13 may become even more relevant in these conditions. Yet, experimental evidence of the

14 interactive effect of AM fungi and P availability influencing altered SOM cycling under

15 eCO₂ is scarce and the mechanisms of this control are poorly understood. Here, we performed

16 a pot experiment manipulating P availability, AM fungal presence and atmospheric CO₂

17 levels and assessed their impacts on soil C cycling and plant growth. Plants were grown in

18 chambers with a continuous ¹³C-input that allowed differentiation between plant- and SOM-

19 derived fractions of respired CO₂(R), dissolved organic C (DOC) and microbial biomass

20 (MBC) as relevant C pools in the soil C cycle. We hypothesised that under low P availability,

 $21 \qquad \text{increases in SOM cycling may support sustained plant growth under eCO_2 and that AM fungi$

22 would intensify this effect. We found the impacts of CO₂ enrichment and P availability on

soil C cycling were generally independent of each other with higher root biomass and slight

24 increases in soil C cycling under eCO₂ occurring regardless of the P treatment. Contrary to

25 our hypotheses, soil C cycling was enhanced with P addition suggesting that low P conditions

- 26 were limiting soil C cycling. eCO₂ conditions increased the fraction of SOM-derived DOC
- 27 pointing to increased SOM decomposition with eCO₂. Finally, AM fungi increased microbial
- 28 biomass under eCO₂ conditions and low-P without enhanced soil C cycling, probably due to
- 29 competitive interactions with free-living microorganisms over nutrients. Our findings in this
- 30 plant-soil system suggest that, contrary to what has been reported for N-limited systems, the
- 31 impacts of eCO₂ and P availability on soil C cycling are independent of each other.
- 32 *Keywords:* Elevated CO₂, Soil organic matter, decomposition, isotopic natural abundance,
- 33 phosphorus limitation, fertilisation effect, mycorrhizal fungi, carbon pools.

34 **1. Introduction**

35 The effect of increased atmospheric CO₂ concentrations on plant productivity and its 36 potential impacts on soil carbon (C) are regulated by nutrient availability (Ainsworth and 37 Long, 2004; Finzi et al., 2011; Reich et al., 2014; Terrer et al., 2019). Higher plant 38 productivity and increases in soil C with elevated CO_2 (eCO₂) have been observed when 39 combined with nutrient addition (De Graaff et al., 2006; Dieleman et al., 2010; Hungate et al., 2009; Maroco et al., 2002; Piñeiro et al., 2017; van Groenigen et al., 2006). In the 40 41 absence of external nutrient supplies, higher soil organic matter (SOM) decomposition under 42 eCO₂ contributes to meet plant nutrient demands with negative impacts on soil C (Carrillo et 43 al., 2014; Drake et al., 2011; Finzi et al., 2006; Hoosbeek et al., 2004; Phillips et al., 2012; 44 Reich et al., 2006). However, a pre-existent nutrient limitation may constrain productivity 45 responses to eCO₂ (Ellsworth et al., 2017; Reed et al., 2015; Reich et al., 2006), potentially 46 preventing any changes in soil C. The current understanding about nutrient availability 47 regulating eCO₂ impacts on C cycling has been obtained mainly from northern hemisphere 48 systems, where nitrogen (N) often constrains ecosystem productivity. But the impacts of 49 eCO₂ on soil C may be dependent on what the most limiting nutrient is (Dijkstra et al., 2013). 50 Phosphorus-limited tropical and subtropical ecosystems are relevant C sinks that cover a vast 51 area (Keenan et al., 2015; Pan et al., 2011; Soepadmo, 1993) and although the role of P 52 availability regulating the impact of eCO₂ on soil C storage is broadly recognised (Goll et al., 53 2012; Reed et al., 2015; Sun et al., 2017; Vitousek et al., 2010), the mechanisms of this 54 control are poorly understood (Ellsworth et al., 2017; Norby et al., 2015; Terrer et al., 2019). Opposite to N, soil P cycling is less coupled to soil C dynamics because P mobilisation is 55 56 not necessarily linked with SOM decomposition. P cycling can be divided into inorganic and 57 organic sub-cycles (Liu and Chen, 2008; Mullen, 2005). In the inorganic sub-cycle, 58 mobilisation of inorganic P occurs via biochemical mechanisms without SOM mineralisation

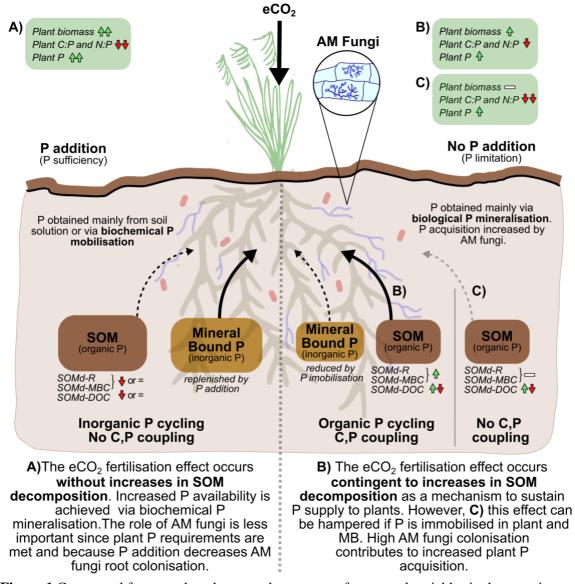
59 nor CO₂ production, thus C and P cycles are not coupled. In the organic P sub-cycle, C and P 60 cycles are more coupled and P mineralisation occurs via biological mechanisms that lead to 61 SOM decomposition and CO₂ production (McGill and Cole, 1981; Sharma et al., 2013; Tate, 62 1984). Most P is made available via inorganic P mobilisation, particularly with P sufficiency. 63 But in low-P systems, mineral P sources are depleted and P derived from rock weathering is 64 minimal (Goll et al., 2012; Reed et al., 2011). Therefore, the majority of P acquisition occurs via internal recycling of organic P sources (Fox et al., 2011; Johnson et al., 2003; Reed et al., 65 66 2011) involving SOM decomposition and CO₂ production. Most of the research on the effects 67 of eCO₂ on soil P availability are focused on its impacts on inorganic P cycling via 68 biochemical P mobilisation mechanisms. For example, higher microbial P enzymatic activity 69 is reported with eCO₂ (Hasegawa et al., 2016; Kelley et al., 2011; Moorhead and Linkins, 70 1997; Souza et al., 2017) as well as higher production of organic acids and siderophores 71 involved in P mobilisation (Fransson and Johansson, 2010; Högy et al., 2010; Tarnawski and Aragno, 2006; Watt and Evans, 1999). Moreover, increases in fine root production and 72 73 changes in pH and soil moisture with eCO₂ can indirectly enhance inorganic P cycling in 74 soils (Dijkstra et al., 2012; Hasegawa et al., 2016; Jin et al., 2015). It is likely that organic P 75 cycling is also affected by eCO_2 due to the influence of eCO_2 on ecosystem stoichiometry 76 (Gifford et al., 2000; Loladze, 2014, 2002) but experimental evidence for this is scarce. A 77 better understanding about how eCO₂ and P availability influence soil C cycling is important 78 to better predict changes in global soil C due to eCO₂ (Finzi et al., 2011; Goll et al., 2012; 79 Yang et al., 2016). 80 In low-P conditions, organic P cycling may help sustain plant productivity under eCO₂ at

the expense of higher SOM decomposition. A study by Jing et al, (2017) demonstrated that
inputs of labile C led to higher SOM decomposition in low-P soils mediated by altered
microbial community composition and increases in microbial biomass. Thus, as C allocation

84 belowground increases with eCO₂ (Canadell et al., 1995; Cotrufo and Gorissen, 1997; Pausch 85 and Kuzyakov, 2018), the higher labile C availability, increased microbial biomass and 86 activity and higher SOM decomposition may allow for a sustained nutrient release for plant 87 uptake and growth in low-P soils. However, evidence from field experiments exposed to 88 eCO₂ suggests that under P-limiting conditions, changes in soil C cycling are minor or 89 undetected (Castañeda-Gómez et al., 2020; Dijkstra et al., 2013; Drake et al., 2016) and that 90 the CO₂ fertilization effect is reduced or not observed (Conroy et al., 1992; Deng et al., 2015; 91 Ellsworth et al., 2017; Goll et al., 2012; Peñuelas et al., 2012; Zhang et al., 2014). Thus, if P 92 availability becomes limiting, ecosystem productivity and microbial-mediated SOM 93 decomposition may be hampered, preventing further changes in soil C cycling. Soil microbial 94 communities play a key role in SOM decomposition and both biological and biochemical P 95 mineralisation and mobilisation. Moreover, the responses of SOM decomposition and C 96 stabilisation to P availability have been found to be largely mediated by changes in 97 saprotrophic microbial community composition and activity (Fang et al., 2019; Feng and 98 Zhu, 2019; Liu et al., 2018; Soong et al., 2018; Yuan et al., 2021) and by symbiotic fungi 99 such as arbuscular mycorrhizal (AM) fungi (Xu et al., 2018). Yet the role of saprotrophic and 100 mycorrhizal fungi altering SOM decomposition under eCO₂ conditions and low-P availability 101 is poorly understood.

AM fungi are known to greatly contribute to P mobilisation under eCO₂ conditions, especially in ecosystems with low soil P availability (Delucia et al., 1997; Matamala and Schlesinger, 2000; Terrer et al., 2016) where AM fungi tend to be more abundant (Treseder and Cross, 2006). AM fungi also alleviate P limitation conditions for plants (Soka and Ritchie, 2014; Willis et al., 2013) possibly via excretion of P hydrolytic enzymes (Joner et al., 2000; Sato et al., 2015) and certainly via uptake of inorganic nutrients from the soil solution (Javot et al., 2007; Smith et al., 2011). They have also been linked with increased 109 SOM decomposition and changes in soil C cycling with climate change (Carrillo et al., 2015; 110 Cheng et al., 2012; Talbot et al., 2008; Wei et al., 2019). Higher C transfer to AM fungi 111 under eCO₂ enhances AM fungi growth and activity (Mohan et al., 2014) and so their 112 contribution to plant P acquisition in P-limited ecosystems may be enhanced with eCO₂ 113 (Cavagnaro et al., 2011). On the other hand, AM fungi can also suppress eCO₂-mediated 114 increases in plant growth due to increased P immobilisation in AM fungal biomass and 115 increased competition over P with saprotrophic communities (Jin et al., 2015) that can 116 simultaneously decrease AM fungi-mediated SOM decomposition (Talbot et al., 2008). If P 117 availability is low, the role of AM fungi mobilising P and promoting plant growth may be 118 constrained due to P-limitation (Treseder and Allen, 2002), which might further regulate 119 impacts of eCO₂ on ecosystem productivity in P-limited ecosystems. Due to the significant 120 role of AM fungi on P and SOM dynamics, predictions of altered soil C stocks under eCO₂ 121 require an understanding how eCO₂ can change AM fungal activity, their impact on 122 saprotrophic communities and plant responses to eCO₂, particularly under P limitation. 123 In this study, we explored the role of P availability and AM fungi mediating responses of 124 SOM decomposition and plant productivity to eCO₂. We did this by manipulating P 125 availability, AM fungal presence and atmospheric CO₂ levels in a pot experiment. An 126 Australian native grass species (*Microlaena stipoides*) with the ability to grow in a wide 127 range of P availability conditions and forming arbuscular mycorrhizal associations (Clark et 128 al., 2014; Hill et al., 2010) was grown in chambers that allowed the continuous isotopic 129 labelling of plant tissues so we could directly assess changes in SOM-derived C in different C 130 pools: respired CO_2 (R; the product of decomposition), dissolved organic C (DOC; 131 considered an active pool of C) and microbial biomass C (MBC; the agent of SOM 132 decomposition). This approach allowed for a holistic assessment of the impact of eCO_2 , P 133 availability and AM fungal presence on soil C cycling and plant growth. Changes in

134	saprotrophic communities were also measured (using phospholipid fatty acids) in relation to
135	eCO2 conditions, P treatment and AM fungal presence. These responses, along with nutrient
136	contents in plant tissues and soil extracts (dissolved nutrients), were collected to enhance
137	result interpretations.
138	We hypothesised that for plants grown under eCO2 and enhanced P availability –
139	achieved by P addition – a fertilisation effect would be observed as increased plant biomass
140	without changes in SOM decomposition (Figure 1A) since P addition will allow for plant
141	requirements to be met under eCO ₂ . Contrarily, if P is low, an eCO ₂ fertilisation effect would
142	only occur with increases in SOM decomposition, evidenced as higher SOM-derived respired
143	CO ₂ , MB and MB (Figure 1B). Alternatively, we hypothesised that no fertilisation effect
144	would be observed if soil C and P cycling are decoupled in this experimental system (Figure
145	1C). The presence of AM fungi would increase SOM decomposition as a mechanism to
146	obtain P from SOM via changes in saprotrophic communities. This effect will be observed in
147	the different C pools, particularly with low P conditions and eCO2 since both low P
148	availability and eCO ₂ promote AM fungal colonisation rates and activity.



150 Figure 1 Conceptual framework and expected responses of measured variables in the experiment 151 (green arrow: increase; red arrow: decrease; -: no effect). Increases in plant biomass under eCO₂ 152 conditions are dependent on nutrient availability. In low phosphorus (P) soils, A) P addition (left side) 153 will support an increase in plant biomass with eCO_2 (fertilisation effect) without major impacts in soil 154 organic matter-derived (SOMd) C pools: R, respired CO₂; MB, microbial biomass; DOC, dissolved 155 organic C. Without P addition (right side), B) P acquisition may be more reliant on organic P cycling 156 (coupled C and P cycles) with enhancements in plant biomass with eCO₂ only observed if SOM 157 decomposition and biological P mineralisation increase. Immobilisation of P by plants, AM fungi and 158 microbial biomass becomes relevant in this scenario further increasing SOM decomposition C) If P 159 availability is too low, becoming limiting, increased P immobilisation and higher P competition will 160 decrease the impact on SOM decomposition and plant biomass.

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162 **2. Materials and methods**

163 We set up a factorial experiment with three factors: **phosphorus** (**P**) **treatment**: P

164 addition (+P)/control; AM fungi: Active AM inoculum (AM)/Non-mycorrhizal (Sterile AM

165	inoculum, NM) and CO ₂ : ambient (400 ppm - aCO ₂) and elevated (640 ppm - eCO ₂) for a
166	total of 8 treatment combinations with 4 replicates each (+P/AM/aCO ₂ ; +P/AM/eCO ₂ ;
167	+P/NM/aCO ₂ ; +P/NM/eCO ₂ ; control/AM/aCO ₂ ; control /AM/eCO ₂ ; control /NM/aCO ₂ ;
168	control /NM/eCO2; n=32, Figure SI 1A). Plants of Microlaena stipoides (Labill.) R.Br., a
169	native Australian C3 grass with a broad P range, were grown from seed in 125mm pots
170	containing around 500 g of dry soil. Plants were grown in growth chambers for twelve weeks
171	and soil water content was kept between 15 and 20 % gravimetric content by addition of
172	MilliQ water as needed every 2-3 days.

173 **2.1. Soil characteristics and treatment preparation**

174 Soil was collected from a Cumberland plain natural woodland in Western Sydney 175 (33°37'01" S, 150°44'26" E, 20 m.a.s.l) near the Eucalyptus Free-Air CO₂ experiment 176 (EucFACE). Soil at this site is an aeric podsol, slightly acidic (pH 5.38 ± 0.02 at 0-10 cm) (Ross et al., 2020), with a N content of 677 mg Kg⁻¹, low total C (1.8%, 0-15 cm), and P 177 content (76.28 mg Kg⁻¹) (Hasegawa et al., 2016). Previous studies have demonstrated that 178 179 both the vegetation (Crous et al., 2015) and soil fauna (Nielsen et al., 2015) are limited by P 180 within this ecosystem. Soil was sieved (to 2 mm) and sterilised (gamma irradiated, 50 kGy) 181 to remove viable AM fungal propagules. To reintroduce a homogenous microbial community 182 to all pots, excluding AM fungi, we prepared a microbial inoculum. For this, approx. 3 Kg of 183 freshly collected soil from a nearby grassland was mixed with water in a 1:3 proportion by 184 volume and the suspension passed through a 20 µm mesh sieve to remove AM fungal spores 185 (Brundett and Australian Centre for International Agricultural Research, 1995). Prior to potting, the filtrate was added to the sterile soil at a rate of 50 mL filtrate per Kg soil. The soil 186 187 was incubated for a week at room temperature, mixing it daily. 188 The soil was then divided into four subsets based on whether they were to receive

additional P and AM fungal inoculum. For the former, triple super phosphate (Richgro, super

190 phosphate fertiliser supplement 9.1% P w/w) was added in a rate of 0.4 g/kg dry soil. The 191 triple super phosphate powder was weighed, diluted and sprayed into the soil while low P 192 treatments received MilliO water only. The soil was then thoroughly mixed and left to settle 193 for a day. Next, AM fungi treatments were created by applying AM fungal inoculum in a 194 1:10 AM fungi inoculum:soil with non-mycorrhizal controls produced by autoclaving the live 195 AM inoculum (121 °C, two hours) and applying it to the soil in the same manner (See 196 supplementary information for AM inoculum production). Once the four types of soils 197 were prepared, soil was added to pots and five surface sterilised *M. stipoides* seeds (30 % 198 H₂O₂ for 10 minutes followed washing) were sown per pot. Pots were randomly split and 199 placed in aCO₂ and eCO₂ chambers. After two weeks of growth, pots were thinned to one 200 plant per pot. Unplanted pots (n=16) with the different P and CO₂ treatments (+P/NM/aCO₂; 201 +P/NM/eCO₂; control/NM/aCO₂; control/NM/eCO₂ n=4 each), were included to estimate 202 effects on SOM decomposition in the absence of plants and AM fungi (Figure SI 1A). 203 Unplanted pots were all prepared using soil with sterile AM inoculum and were kept under 204 the same conditions as the planted pots for the duration of the experiment.

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2.2. Growth Chambers set up

206 The growth chambers (six in total, three per CO_2 treatment) were modified using the 207 approach of Cheng and Dijkstra (2007) to achieve a continuous ¹³C- labelling of plant tissues 208 in both aCO₂ and eCO₂ treatments. The chambers were adapted to take an influx of naturally 209 ¹³C-depleted CO₂ (δ^{13} C= -31.7 °/₀₀ ±1.2) delivered during the photoperiod, combined with a 210 scrubbing system made of a soda lime-filled PVC tube (72 L) that allowed continuous supply 211 of CO₂-free air (Figure SI 1B). Chambers were adjusted to a 16h/8h photoperiod with 25 212 °C/18 °C, 60 % relative humidity, and light intensity of 900 µmol/m²s¹. The chamber 213 atmosphere was sampled frequently to confirm depletion in ¹³C. Air samples from the 214 chambes were extracted via a pump system into a tedlar bag (Tedlar® SCV Gas Sampling

Bag) and analysed for δ^{13} C in a PICARRO G2201i isotopic CO ₂ /CH ₄ analyser (Picarro I

- 216 Santa Clara, CA, USA). To avoid plant-uptake of ¹³C from outside the chambers, watering
- and all other manipulations were performed during the night period aided by green light.
- 218 **2.3. Harvest and sample processing**
- 219 Gas sampling of the plant-soil system

220 After twelve weeks of growth, we quantified rates of total respired CO₂ (R) and its C 221 isotopic composition as described by Carrillo et al., (2015, 2014). Briefly, pots were placed 222 on an elevated platform inside a water filled tray and covered with a PVC chamber (45 cm H 223 x 15 cm D) adapted with an air-tight rubber stopper for air sampling. Free-CO₂ air was 224 circulated for 2 hours using an aquarium pump connected to a CO₂ scrubber (50 cm H x 4 cm 225 D PVC tubing filled with soda lime). After 2 hours of scrubbing, an air sample was taken to 226 determine baseline CO₂ concentrations using a 7890A gas chromatograph with a G1888 227 network headspace sampler (Agilent Technologies, USA). Later, the CO₂ scrubbers were 228 removed and the pump reconnected to the PVC chamber to allow for air circulation while 229 pots were incubated. After two to three hours of incubation, a gas sample was collected in an 230 airtight gas collection bag (Tedlar® PVDF, 1L) using an aquarium pump system and 231 analysed for its C isotopic composition.

232 Plant and soil harvest

One day after gas sampling, pots were destructively harvested. Aboveground biomass was cut and roots separated from the soil and washed. Plant aboveground biomass, roots and a subsample of fresh soil were oven dried at 60 °C for measurements of dry plant biomass, soil gravimetric water content and total nutrients. A fresh soil subsample was stored at -20 °C for microbial community analyses and the remaining fresh soil was used for assessments of dissolved nutrients and microbial biomass.

239 Plant, soil nutrients and microbial biomass C and N

240	Two sub-samples of soil were weighed. Dissolved nutrients in soil were extracted
241	from the first subsample with a 0.05M K ₂ SO ₄ solution in a 4:1 solution to soil ratio, shaking
242	at 180 rpm for an hour. Samples were filtered through a Whatman # 42 filter paper and
243	frozen (-20°C) until analyses. The second subsample was fumigated for 5 days with
244	chloroform and then nutrients were extracted as for unfumigated samples. The fumigated and
245	unfumigated extracts were analysed for total dissolved organic C and N (Shimadzu® TN,
246	TOC-L, Japan) and microbial biomass C and N calculated by subtracting fumigated and
247	unfumigated samples (Vance et al., 1987). The volume left of these K2SO4 extractions was
248	used to obtain the isotopic composition of DOC and microbial biomass (see section below).
249	Phosphates were extracted following the Bray 1-P method for acidic soils, in a 1:7
250	solution:soil ratio using a 0.03M NH ₄ F solution in 0.025 M hydrochloride (HCl) adjusted pH
251	to 2.6 ± 0.05 with HCl (Rayment et al., 2010). Samples were manually shaken for 60s and
252	immediately poured over a Whatman # 42 filter paper. Collected extracts were analysed by
253	colorimetry (AQ2 Discrete Analyser, SEAL Analytical, Mequon, WI, USA).
254	Total P and N were determined from ground oven-dried soil and root samples. Total P
255	concentration was analysed by an X-ray fluorescence spectrometer (PANalytical, spsilon 3.
256	10Kv, 0.9mA. Lelyweg, Almelo, Netherlands) while total C and N from soil, roots and soil
257	extracts were analysed along their isotopic composition.
258	2.4. Respired, dissolved organic C (DOC) and microbial biomass isotopic
259	composition and partitioning
260	Gas samples from plant-soil system incubations were analysed in a PICARRO analyser
261	(G2201i; Picarro, Santa Clara, CA, USA. Precision values below 0.16‰) for the δ^{13} C and
262	CO_2 concentration of the total respired CO_2 (R). For the isotopic composition of the DOC and
263	microbial biomass C (MBC), fumigated (f) and unfumigated (uf) soil K ₂ SO ₄ extracts were
264	oven dried at 60°C. The dried extracts were scraped and weighed for analysis on a Thermo

GC-C-IRMS system (Trace GC Ultra gas chromatograph, Thermo Electron Corp., Milan,

266 Italy; coupled to a Delta V Advantage isotope ratio mass spectrometer through a GC/C-III); University of California, Davis Campus, USA). The δ^{13} C of unfumigated soil extracts was 267 268 used as the isotopic composition of the DOC while $\delta^{13}C$ of MBC ($\delta^{13}C_{MBC}$) were calculated from δ^{13} C values of both fumigated (*f*) and unfumigated (*uf*) extractions (See supplementary 269 270 information for calculations of the isotopic composition of the microbial biomass). 271 Samples of dried soil were also analysed on the Thermo GC-C-IRMS system for C, N and δ^{13} C. 272 273 To calculate the fractions of SOM-derived C in the total respired CO₂ (R), dissolved 274 organic C (DOC) and microbial biomass C (MBC) we used isotopic partitioning as: SOM. $C_{R,DOC,MBC} = (\delta^{13}C_{R,DOC,MBC} - \delta^{13}C_p) / (\delta^{13}C_{control} - \delta^{13}C_p)$. Where $\delta^{13}C_{R,DOC,MBC}$ were the 275 δ^{13} C of either the CO₂ measured in the total respired CO₂ (R), dissolved organic C (DOC) or 276 microbial biomass C (MBC)) from each pot; $\delta^{13}C_p$ is the isotopic ratio of the plant source 277 averaged across P treatments per CO₂ condition (root biomass at aCO₂ δ^{13} C= -40.01±0.08 278 279 and eCO₂ δ^{13} C = -43.20±0.04) and δ^{13} C_{control} is the average δ^{13} C of the native SOM source obtained from the dried bulk soil across CO₂ and P conditions ($\delta^{13}C = -25.31 \pm 0.02$). The 280 281 fractions of plant-derived C were obtained by subtracting SOM-derived fractions from the 282 unit. Measured total respired CO₂ rates. DOC and MBC were partitioned into SOM-derived 283 and plant-derived, using these fractions to obtain the mass of the SOM derived R, MBC and 284 DOC pools. In this study, we focus only on the SOM-derived C pools and fractions.

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2.5. Microbial community analysis: Phospholipid-derived fatty acids (PLFA) and Neutral lipid-derived fatty acids (NLFA)

Soil PLFAs were extracted to assess the overall microbial communities while the
16:105c NLFA was used as an indicator of arbuscular mycorrhizal (AM) fungi presence and

abundance (Olsson, 1999). Freeze-dried soil from planted pots (n=32) were extracted
following the protocol by Buyer & Sasser (2012) with modifications by Castañeda-Gómez et
al. (2020) (See supplementary information for PLFA and NLFA analyses). Functional
groups were defined as shown in Table SI 1. Fungal to bacterial ratio (F:B) was calculated
by dividing fungal PLFAs (not including AM fungi) by the sum of bacterial PLFAs. The sum
of individual lipids was used as an indicator of the size of the community (µg PLFA g⁻¹ dry
soil).

296 **2.6. Statistical analyses**

297 The effect of CO₂ condition, AM fungi treatment and P addition and their interactions on 298 the response variables was analysed with a linear model fitted with the function "lm" from 299 the stats package in R version 3.3.2 (R Core Team 2016). This approach was used instead of 300 a mixed effects modelling approach since pots were moved among chambers with the same 301 CO₂ treatment so it was not possible to estimate a random effect associated with each 302 chamber. The normality of the residuals of each model was inspected to check the 303 appropriateness of the fit and transformations were performed when needed. Statistical 304 significance was determined performing an ANOVA (Analyses of variance) with the 305 "Anova" function ("car" package, Fox et al. 2017). Multiple mean comparisons were 306 performed using the Tukey test using the "glht" function ("multcomp" package, Hothorn et 307 al. 2017). As soil moisture can affect soil respiration measurements, we tested for 308 correlations between soil moisture and the total respired CO₂, SOMd- respired CO₂ fraction 309 and total mass of SOMd- respired CO₂. For response variables with a significant correlation 310 with moisture, a 3-way ANCOVA was performed with soil moisture as covariate, to account 311 for the variability brought by the slightly different moisture contents at the time of CO₂ 312 sampling. The homogeneity of the regression slopes, normality of residuals and homogeneity of variances was tested when performing the ANCOVA. Significance levels were: $\leq 0.1(.), \leq 0.05(*), \leq 0.01$ (**), $\leq 0.001(***)$.

315 Microbial communities were analysed as the PLFA-based total microbial biomass (sum of all individual lipids, in μ g PLFA g⁻¹ dry soil), as absolute microbial abundance (reflecting 316 the size or biomass of the community, in μ g PLFA g⁻¹ dry soil) and as relative abundance (in 317 318 percentage, reflecting microbial community composition) of the different microbial groups 319 and as individual lipids. To visualise significant three-way interactions of experimental 320 factors in the abundance of microbial groups from the ANOVA, the "emmip" function from 321 the "emmeans" package (Lenth et al., 2020) was used to show the estimated marginal means 322 from the fitted linear model.

323 **3. Results**

324 **3.1.** Influence of elevated CO₂, P availability and AM fungi on SOM decomposition 325 and C pools

Higher SOM-C losses under eCO₂ were expected for control P conditions (low P) and 326 327 particularly when AM fungi were inoculated (Figure 1). However, we found that in general, 328 there was not a significant interactive impact of the three experimental factors on the respired 329 CO₂, microbial biomass C (MBC) and dissolved organic C (DOC) as initially expected. 330 Instead, we found that the interaction of AM and P treatments was relevant determining the 331 fate of these soil C pools, regardless of eCO₂. AM and P treatments significantly affected the 332 total respired CO₂ (Figure 2), which decreased with P addition but only for the NM treatment 333 (P<0.05, Tukey's multiple comparison). However, we did not find any significant effects of CO₂ conditions, AM or P treatments on the SOM-derived respired CO₂ (Figure 2). On the 334 335 other hand, most of the microbial biomass C (MBC) was derived from SOM (above 80%, 336 Table SI 2) and while neither the total microbial biomass C nor the SOM-derived MBC (Figure 2) were affected by the experimental factors, the C sources used by the microbial 337

- biomass were affected by the interaction between AM and P treatment (Table SI 2). The
- fraction of MBC derived from SOM marginally increased by 6% with P addition in NM
- 340 treatments (P=0.1, Tukey's multiple comparison. **Table SI 2**). Similar to the MBC pool, most
- of the DOC was derived from SOM (above 80%, **Table SI 2**) and both the total DOC and
- 342 SOM-derived DOC were affected by the interaction between AM and P treatments (Figure
- 343 2) with P addition significantly increasing the SOM-derived DOC in NM treatment (P=0.001,
- 344 Tukey's multiple comparison). Finally, although not evidenced in the total DOC or SOM-
- 345 derived DOC, the fraction of DOM-derived DOC significantly increased with eCO₂ (Table
- 346 **SI 2**).

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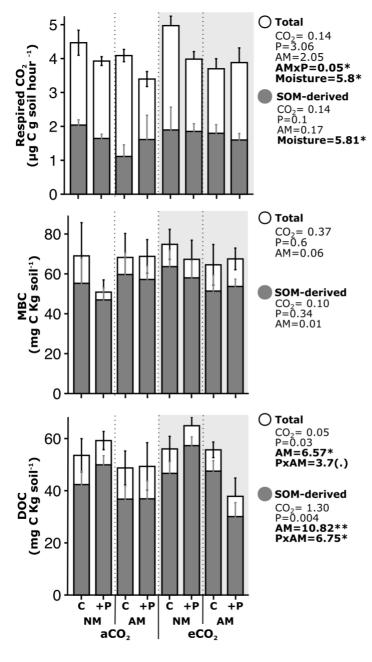


Figure 2. Mean values (± standard error, n=4) of the total(white) and SOM-derived (grey) respired CO₂, microbial biomass C (MBC) and dissolved organic C (DOC) per CO₂ condition, AM fungi and P addition treatment (C: control and +P: P addition). CO₂ conditions are shown on light grey (elevated) and white (ambient) backgrounds. $F_{(1,24)}$ values of ANCOVAs (for respired CO₂) and ANOVAs from linear model for response variables displayed on the right. Significance levels: $\leq 0.1(.), \leq$ $0.05(*), \leq 0.01 (**), \leq 0.001(***).$ Only significant interactions of the treatments displayed.

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3.2. Soil nutrient responses to elevated CO₂, P availability and AM fungi

383 We expected eCO₂ to increase C and nutrient availability, P addition to increase soil P,

and AM fungi to decrease soil P and N due to higher nutrient acquisition and immobilisation

- 385 by the fungal biomass. These expectations were partially supported with our observations.
- 386 eCO2 increased total soil N, but only for control P pots while dissolved N marginally
- decreased with eCO₂ for AM pots (**Table 1**). Total soil P was unaffected by eCO₂ conditions
- 388 while soil dissolved P significantly decreased with eCO₂ conditions, particularly for control P

and AM pots (Table 1). P addition increased total soil P while AM decreased total P and

390 dissolved N (Table 1). Finally, we detected lower dissolved C:N with P addition for AM pots

391 but higher dissolved C:N for AM pots with eCO₂ (Figure 3). Dissolved C:P decreased with P

addition for pots under eCO₂ conditions while it increased with eCO₂ for AM pots (**Figure 3**).

Table 1 Soil nutrients (total and dissolved (ppm)) per CO₂ condition, AM fungi and P addition treatment. Dissolved nutrients as: DOC – total dissolved organic Carbon, TN – total nitrogen and phosphates (PO₄). Mean values (± standard error, n=4) followed by the same letter are not significantly different (p≤ 0.05, Tukey multiple comparison test), no letters indicate non-significant effect of the treatments. AM: Arbuscular mycorrhizal and NM: non-mycorrhizal. Below, ANOVA results from linear model (lm) for response variables (× natural log transformed). F_(1,24) values displayed with significance levels: $\leq 0.1(.), \leq 0.05(*), \leq 0.01$ (**), $\leq 0.001(***)$.

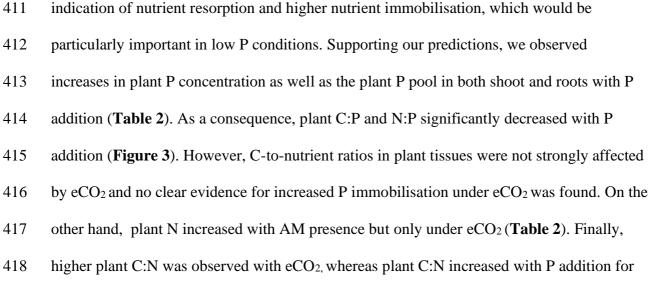
			Tota	ıl	Dissolved		
			Soil N	Soil P	TN≭	PO₄×	
		control	347.2(9.5)a	166(20)a	4.2(0.4)	2.49(0.8)ab	
	NM	+ P	558.2(39.7)ab	268(10)b	4.2(0.5)	1.25(0.1)ab	
		control	472.9(23.5)a	167(12)a	3.6(0.6)	3.86(0.3)b	
aCO ₂	AM	+ P	526.1(60.7)ab	207(32)a	4.8(1.0)	2.75(0.6)ab	
		control	665.0(116.4)b	155(26)a	4.4(0.1)	1.08(0.1)a	
	NM	+ P	425.1(63.0)ab	238(21)b	4.5(0.5)	3.40(0.8)ab	
		control	633.7(103.9)b	132(10)a	3.7(0.2)	1.09(0.1)a	
eCO ₂	AM	+ P	643.0(70.3)ab	195(11)a	3.0(0.8)	1.00(0.1)a	
		CO ₂	5.46*	2.52	1.74	8.84**	
		Р	0.03	28.01***	0.16	0.0001	
		AM	2.00	5.41*	6.58*	0.08	
		CO ₂ xP	6.25*	0.003	0.60	7.64*	
		CO ₂ xAM	0.22	0.02	3.64(.)	11.45**	
		AMxP	0.21	2.30	0.15	3.46	
		PxAMxCO ₂	4.23(.)	0.56	1.22	2.90	

401

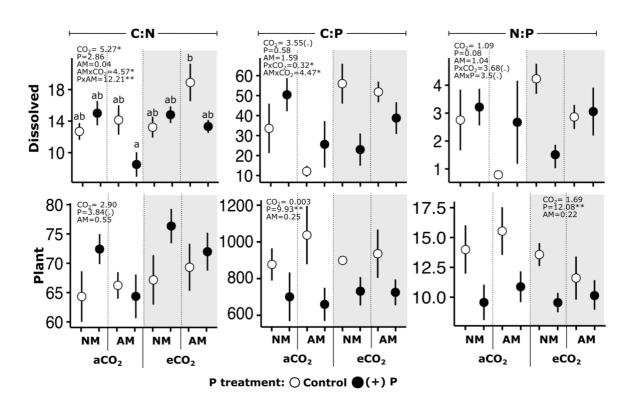
400

402 3.3. Plant responses to elevated CO₂, P availability and AM fungi

403 We expected that plant biomass would increase with eCO₂ when P availability was high, 404 and be unresponsive to eCO₂ when P availability was low unless AM fungi were present, 405 supporting further plant growth by increasing SOM decomposition and P uptake (Figure 1). We did not observe a significant impact of eCO₂ on aboveground plant biomass regardless of 406 407 the P and AM treatments. However, eCO₂ effects increased root biomass when AM fungi were present, regardless of the P treatment (Table 2). Increases in plant biomass with P 408 409 addition were not observed either, but we expected to see increases in plant P concentrations 410 in response to P addition and decreases in plant C-to-nutrient ratios with eCO₂ as an



- 419 NM pots (**Figure 3**).
- 420





422Figure 3. Mean values (\pm standard error, n=4) of mass nutrient ratios (columns) in plant and soil423solution (dissolved) per treatment. eCO2 conditions shown as grey areas and P treatment as424white(control) or black (+P) circles. NM (non-mycorrhizal). $F_{(l,24)}$ values of ANOVAs from linear425model for response variables displayed. Significance levels: $\leq 0.1(.), \leq 0.05(*), \leq 0.01$ (**), \leq 4260.001(***); only significant interactions of the treatments displayed. Treatments with different letters427are significantly different (p ≤ 0.05 , Tukey multiple comparison test).428

- 429
- 430

431 **Table 2** Plant biomass (shoot and roots), nutrient contents and plant P and N pools (plant P and N

432 content per unit of shoot and root biomass) per CO₂ condition, AM fungi and P addition treatment.

433 NM (non-mycorrhizal). Mean values (± standard error, n=4) followed by the same letter are not

434 significantly different ($p \le 0.05$, Tukey multiple comparison test), no letters indicate non-significant

435 effect of the treatments. Below, results of the ANOVA from linear model (lm) for response variables

436 (\times log transformed). $F_{(1,24)}$ values displayed with significance levels: $\leq 0.1(.), \leq 0.05(*), \leq 0.01$ (**), \leq

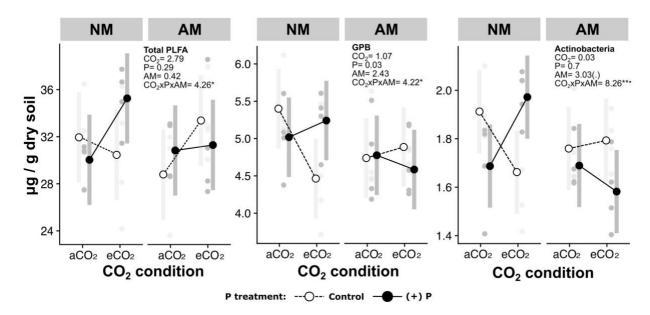
437 0.001(***).

			Biomass		Nutrient Concentration		
			Shoot	Root	Plant C	Plant N×	Plant P
				g		g/Kg	
		control	2.7(0.2)	2.1(0.08)ab	356.9(30.2)	5.5(0.2)	0.45(0.07) a
	NM	+ P	2.8(0.3)	2.0(0.3)ab	380.6(14.6)	5.3(0.1)	0.59(0.09) b
		control	2.8(0.1)	1.7(0.3)a	395.0(11.5)	6.0(0.4)	0.41(0.07) a
aCO ₂	AM	+ P	4.3(1.2)	1.8(0.3)a	400.8(6.4)	6.2(0.4)	0.64(0.08) b
		control	3.2(0.5)	1.7(0.2)ab	391.6(8.9)	5.9(0.3)	0.44(0.01) a
	NM	+ P	3.2(0.4)	2.1(0.2)ab	390.1(8.5)	5.1(0.2)	0.55(0.06) b
		control	3.4(0.3)	2.4(0.4)b	376.1(16.1)	5.5(0.2)	0.44(0.09) a
eCO ₂	AM	+ P	3.6(0.3)	2.4(0.3)b	382.7(14.8)	5.3(0.2)	0.55(0.09) b
		CO ₂	0.3516	2.1	0.03	2.48	0.19
		Р	1.0574	0.2835	0.65	2.52	8.87**
		AM	2.1422	0.1943	0.61	2.35	0.03
		CO ₂ xP	0.9543	0.4923	0.29	1.13	0.54
		CO2xAM	0.4513	5.1081*	3.49(.)	5.22*	0.03
		AMxP	1.241	0.4535	0.06	2.81	0.08
		PxAMxCO ₂	0.758	1.0528	0.33	0.04	0.08
			Р	pool	N p	ool	
			Shoot P	Root P	Shoot N×	Root N	
					mg		
		control	1.2(0.2) a	0.89(0.14) a	15.14(0.87)	11.64(0.65)	
	NM	+ P	1.6(0.2) b	1.25(0.35) b	14.46(1.48)	10.52(1.34)	
		control	1.2(0.2) a	0.66(0.04) a	16.75(1.13)	10.09(0.99)	
aCO ₂	AM	+ P	3.0(1.1) b	1.06(0.06) b	19.38(2.22)	13.84(2.15)	
		control	1.4(0.2) a	0.72(0.06) a	19.29(3.75)	9.85(1.44)	
	NM	+ P	1.8(0.3) b	1.23(0.22) b	16.59(3.05)	11.25(1.37)	
		control	1.5(0.3) a	0.99(0.10) a	18.97(2.59)	13.27(2.04)	
eCO ₂	AM	+ P	2.0(0.3) b	1.22(0.03) b	18.89(1.23)	12.46(1.68)	
		CO ₂	0.03	0.40	1.65	0.03	
		Р	5.62*	10.98**	0.08	0.55	
		AM	1.50	0.11	2.60	2.19	
		CO ₂ xP	1.08	0.002	0.32	0.22	
		CO2xAM	0.67	2.25	0.35	0.43	
		AMxP	1.24	0.28	1.10	0.38	
		PxAMxCO ₂	0.91	0.49	0.01	2.67	

438

439 **3.4.** *Microbial communities including AM*

440 AM fungi presence and abundance was assessed with the $16:1\omega5c$ neutral lipid, which 441 was high in AM-inoculated pots lacking additional P. However, P addition reduced the 442 concentration of the AM fungal lipid in soil to a level approaching the uninoculated pots 443 (AMxP: $F_{(1,24)} = 6.28$, P < 0.001 **Table SI 2**). We assessed the biomass and composition of the microbial community as we expected that AM fungi would alter soil saprotrophic
communities as a mechanisms to enhance SOM decomposition under low P conditions. We
observed a significant increase in microbial biomass with AM fungi in low P (control pots)
and eCO₂ conditions. The increase in microbial biomass in these conditions was particularly
due to the response of Gram positive bacteria and Actinobacteria (Figure 4). Gram negative
bacteria and fungi significantly increased with eCO₂ while protozoa increased with eCO₂ only
for NM pots (Figure SI 2).



451 Figure 4. Estimated marginal means with original data (grey circles) and confidence intervals (grey bars) of the total microbial biomass, GPB (Gram positive bacteria) and actinobacteria 452 $(\mu g/g dry soil)$ from fitted lineal models per experimental treatments. P treatment as white 453 454 circles and dashed lines (control) or black circles and solid lines (+P) for estimated means, 455 and as light grey (control) and dark grey (+P) for original data. NM: non-mycorrhizal. ANOVA results displayed with significance levels: $\leq 0.1(.), \leq 0.05(*), \leq 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), <0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01(**), < 0.01$ 456 457 0.001(***); only significant interactions of the treatments displayed. Grey circles are the original data input for the models. 458

459

460 **4. Discussion**

461 Our experiment investigated whether, under low P conditions, increased SOM

462 decomposition and organic P cycling would allow for a sustained plant fertilisation effect

463 with eCO₂. We expected to observe increases in plant biomass under eCO₂ conditions and

464 low P availability paired with higher SOM derived C pools (SOM derived respired CO₂,

465 MBC, DOC), with further enhancements of SOM decomposition when AM fungi were 466 inoculated due to the role of this symbiotic fungi on P acquisition. We did not find strong 467 evidence suggesting that increases in soil C cycling with low P availability supported plant 468 growth under eCO₂. Instead, we found that the impacts of CO₂ enrichment and P availability on plant growth and soil C cycling were generally independent of each other with increases in 469 470 root biomass and soil C cycling under eCO₂ occurring regardless of the P treatment. Contrary 471 to our hypotheses (Figure 1), soil C cycling was enhanced with P addition, suggesting that 472 low P conditions were limiting soil C cycling. On the other hand, root biomass increased with 473 eCO₂ conditions and AM presence, while microbial biomass increased with eCO₂ in AM-474 inoculated pots and control P conditions mainly due to the response of Gram positive bacteria 475 and Actinobacteria. Taken together, our findings in this plant-soil system demonstrate that C 476 and P biogeochemical cycles may not become coupled to sustain an eCO₂ fertilisation effect 477 but instead, that low soil P will limit C cycling responses to eCO₂ with AM potentially 478 outcompeting saprotrophic communities for nutrients and hampering SOM decomposition. 479 The CO₂ fertilization effect is modulated by the interaction of mycorrhizal associations 480 and nutrient availability (Terrer et al., 2019, 2016; Treseder, 2004). In this study, although 481 the increases in root biomass under eCO₂ conditions occurred when AM fungi were present, 482 this effect occurred regardless of P availability. The CO₂ fertilization effect is a product of 483 higher plant photosynthetic activity leading to excess C in plant tissues leading to increased 484 growth. In this study, we observed a marginally higher plant C:N ratio with eCO₂ (Figure 3) 485 but no significant increases in plant biomass. The observed additional assimilated C was 486 potentially allocated belowground, driving the detected increases in root biomass and AM 487 fungi growth and activity (Cheng and Johnson, 1998; Mohan et al., 2014; Treseder, 2004). 488 The presence of AM fungi has been previously reported to increase root biomass with eCO₂ 489 (Baslam et al., 2012; Dong et al., 2018; Zhu et al., 2016) and root biomass responses to eCO₂

490 may not be necessarily affected by low P conditions (Jiang et al., 2020). It is hypothesised 491 that the positive feedback of eCO₂ and AM fungi on plant growth is mediated by the nutrient 492 uptake role of these symbiotic fungi (Alberton et al., 2005). While P availability has been 493 found to drive the CO₂ fertilization effect on ectomycorrhizal plants, the response of 494 arbuscular mycorrhizal plants to eCO₂ seems to be more dependent on soil N availability 495 (Terrer et al., 2019). Although we did not find that P availability mediated increases in root 496 biomass with eCO₂ and AM fungal presence, we observed that AM fungi marginally 497 increased plant N contents in the aboveground biomass (Table 2), as well as decreased 498 dissolved N and P in soil with eCO₂ (**Table 1**) and plant C:N with P addition (**Figure 3**), all 499 of this indicating that AM fungi might have promoted N uptake when P was added and 500 facilitated the observed increases in root biomass.

501 We hypothesised that higher SOM decomposition would occur with low P availability 502 as organic P cycling and biological P mobilisation mechanisms are crucial when mineral P 503 sources are depleted (Reed et al., 2011). Contrary to our expectations, we found that P 504 addition enhanced soil C cycling. Increases in SOM decomposition with P addition in P-505 limited systems have been previously reported (Cleveland et al., 2006) and can be explained 506 by the higher affinity of P to mineral surfaces and the consequent release of labile C that 507 promotes microbial activity (Mori et al., 2018). We expected that AM fungi presence would 508 further enhance SOM decomposition in low P conditions (Figure 1). Labile C funnelled via 509 AM fungi to the saprotrophic community may enhance soil microbial community activity and 510 promote SOM decomposition (Frey, 2019). Although we observed higher microbial biomass 511 under eCO₂ and AM presence in control P conditions, we suggest that the lack of significant 512 increases in SOM decomposition in these conditions may be due to higher competition 513 between AM fungi and saprotrophic communities that might have limited enhanced soil C 514 cycling (Zhou et al., 2019). The changes in plant C:N, soil dissolved C:N, total soil C:N with

AM fungi explained above further support the idea that AM fungi might outcompete saprotrophic microbes in N and P uptake. Finally, we detected a decrease in SOM decomposition with P addition and AM-inoculated pots (**Figure 2**), but caution must be taken when interpreting this result since P addition also significantly reduced AM fungi presence and thus, it cannot be claimed that the reduction in SOM decomposition with P addition in AM pots is due to AM fungi presence.

521 We expected that SOM decomposition and, thus, organic P cycling would increase with 522 eCO₂ particularly under low P conditions as a mechanism to sustain plant nutrient demands 523 (Figure 1). Our results show evidence for increased C cycling with CO₂ enrichment, 524 particularly for the fraction of SOM-derived DOC, but this effect was not dependent on P 525 availability. Increases in DOC with eCO₂ have been previously attributed to increases in C 526 allocation belowground and higher rhizodeposition (Drake et al., 2011; Freeman et al., 2004; 527 Lukac et al., 2003; Phillips et al., 2011) but it can also be due to increased SOM 528 decomposition (Hagedorn et al., 2008, 2002). Our isotopic analyses allowed us to detect that 529 increases in DOC under eCO_2 in this system were due to enhanced SOM decomposition 530 given the observed increase in SOM-derived DOC fraction (Table SI 2). The DOC made 531 available via SOM decomposition can be either incorporated in the microbial biomass or lost 532 via leaching. We did not observed changes in the SOM-derived MBC with eCO₂ and thus the 533 extra DOC was likely not incorporated in the microbial biomass but rather lost via leaching 534 (Kindler et al., 2011). On the other hand, only marginally higher soil respiration with eCO₂ 535 was found in the present study which can be due to the low diffusivity of CO₂ from soils, 536 particularly with high soil moisture (Davidson et al., 2000; Hashimoto and Komatsu, 2006; 537 Maier and Schack-Kirchner, 2014). In our experiment, volumetric soil water content was on 538 average 23% for planted pots, which is slightly above the expected (12-21%) soil water

holding capacity for a soil with sandy to loamy texture as the one used for this experiment(Datta et al., 2017; Easton and Bock, 2016).

541 We observed higher P concentrations in plant tissues but this was not accompanied by increases in biomass under any CO₂ condition. Higher P concentration in plant tissues 542 543 without increases in biomass can occur due to a P luxury consumption of plants, where 544 higher P immobilisation is not necessarily related to increases in growth (Brar and Tolleson, 545 1975). Lack of responses to P addition in biomass (Nie et al., 2009) and abundance 546 (Robinson et al., 1993) for Microlaena stipoides have been observed before and attributed to 547 a nutrient accumulation strategy of this grass species, the time and frequency of fertilizer 548 additions and the low response that Australian native grasses typically have to fertilizer 549 application. In this study, only one addition of P as triple superphosphate was done at the 550 start of this experiment. Australian native grasses are less responsive to fertilizer addition 551 during establishment (Nie et al., 2009), which explains the lack of biomass effects of M. 552 stipoides in this experiment. Also, triple superphosphate is highly soluble in water and 553 becomes rapidly available for plant uptake (Mullins et al., 1995) by the first week of 554 application (Ghosal and Chakraborty, 2012). Hence, increases in P concentrations in plant 555 tissues occurred at the beginning of the experiment when AM fungi had not fully colonised 556 the roots and so, their role in soil C transformations and plant P uptake were less relevant. 557 Current understanding of the impacts of eCO₂ on plant productivity with low nutrient 558 conditions have focused on N-limited ecosystems. Higher SOM decomposition under eCO2 559 when N availability is low occurs as a mechanism to sustain nutrient supply and plant 560 growth, with mycorrhizal fungi aiding to deliver the mined nutrients to the plants. For P-561 limited ecosystems however, low P availability generally constrains ecosystem responses to 562 CO₂ enrichment (Ellsworth et al., 2017; Jiang et al., 2020; Reed et al., 2015; Reich et al., 563 2006) and the role of AM fungi mediating plant and soil C responses to eCO₂ with low P

availability is not fully clear. Thanks to the detailed measurement of soil C cycling as changes

- 565 in total C pools, SOM-derived fractions and total mass of SOM-derived C pools, we show
- that the impacts of CO₂ enrichment and P availability on plant growth and soil C cycling
- 567 were independent of each other and are not likely to become coupled with eCO₂ conditions.
- 568 We also demonstrate that although AM fungi may contribute to increases in microbial
- 569 biomass with eCO₂ and low-P conditions, this effect may not translate into enhanced SOM
- 570 decomposition due to increased nutrient competition that may limit saprotrophic
- 571 communities. Our findings highlight that ecosystem responses to eCO₂ with P limitation are
- 572 different from those reported for N-limited systems and therefore, inferences of the behaviour
- 573 of P-limited ecosystems based on current knowledge about N-limited ecosystems are not
- 574 ideal. Moreover, our results also contribute to the current gap in knowledge regarding the
- 575 impacts of soil C cycling with low P availability exposed to eCO₂ conditions.
- 576 5. Acknowledgements
- 577 We thank Gavin McKenzie and Goran Lopaticki for their help with the chamber set up,
- 578 maintenance and troubleshooting for the duration of this experiment. Thanks to Pushpinder
- 579 Matta and Christopher Mitchell for their help with nutrient analyses and to Johanna Pihlblad
- and Johanna Wong- Bajracharya for their assistance during the harvest of this experiment and
- sample processing.

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