

1 Sulfur-fueled chemolithoautotrophs 2 replenish organic carbon inventory in 3 groundwater

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22 **Abstract**

23 Metagenome-assembled genomes (MAGs) have revealed the existence of novel bacterial and
24 archaeal groups and provided insight into their genetic potential. However, metagenomics and even
25 metatranscriptomics cannot resolve how the genetic potential translates into metabolic functions
26 and physiological activity.

27 Here, we present a novel approach for the quantitative and organism-specific assessment of the
28 carbon flux through microbial communities with stable isotope probing-metaproteomics and
29 integration of temporal dynamics in ^{13}C incorporation by Stable Isotope Cluster Analysis (SIsCA). We
30 used groundwater microcosms labeled with $^{13}\text{CO}_2$ and D_2O as model systems and stimulated them
31 with reduced sulfur compounds to determine the ecosystem role of chemolithoautotrophic primary
32 production. Raman microspectroscopy detected rapid deuterium incorporation in microbial cells
33 from 12 days onwards, indicating activity of the groundwater organisms. SIsCA revealed that
34 groundwater microorganisms fell into five distinct carbon assimilation strategies. Only one of these
35 strategies, comprising less than 3.5% of the community, consisted of obligate autotrophs
36 (*Thiobacillus*), with a ^{13}C incorporation of approximately 95%. Instead, mixotrophic growth was the
37 most successful strategy, and was represented by 12 of the 15 MAGs expressing pathways for
38 autotrophic CO_2 fixation, including *Hydrogenophaga*, *Polaromonas* and *Dechloromonas*, with varying
39 ^{13}C incorporation between 5% and 90%. Within 21 days, 43% of carbon in the community was
40 replaced by ^{13}C , increasing to 80% after 70 days. Of the 31 most abundant MAGs, 16 expressed
41 pathways for sulfur oxidation, including strict heterotrophs. We concluded that
42 chemolithoautotrophy drives the recycling of organic carbon and serves as a fill-up function in the
43 groundwater. Mixotrophs preferred the uptake of organic carbon over the fixation of CO_2 , and
44 heterotrophs oxidize inorganic compounds to preserve organic carbon. Our study showcases how
45 next-generation physiology approach like SIsCA can move beyond metagenomics studies by
46 providing information about expression of metabolic pathways and elucidating the role of MAGs in
47 ecosystem functioning.

48 **Introduction**

49 Genome-resolved metagenomics has vastly expanded our knowledge about the microbial
50 communities in Earth's ecosystems (Wrighton et al., 2012; Albertsen et al., 2013; Nielsen et al., 2014;
51 Vollmers et al., 2017). The sequences obtained allowed unprecedented insights into microbial
52 genetic potential and led to the discovery of novel bacterial and archaeal taxa (Brown et al., 2015;
53 Castelle et al., 2015). Studies employing deep sequencing and genome binning approaches have
54 recovered thousands of metagenome-assembled genomes (MAGs) that revealed a broad microbial
55 diversity and greatly expanded the tree of life (Anantharaman et al., 2016; Hug et al., 2016; Parks et
56 al., 2017). Such MAGs enabled researchers to highlight organisms that might play key roles in
57 biogeochemical cycles (Long et al., 2016; Anantharaman et al., 2018), and provide a basis for the
58 establishment of hypothesis about microbial functioning in the investigated habitats. Metagenomics
59 studies targeting the groundwater microbiome, for example, have revealed a high abundance of
60 organisms with the metabolic potential for CO₂ fixation (Emerson et al., 2016; Probst et al., 2018), as
61 well as for the utilization of inorganic electron donors (Anantharaman et al., 2016; Anantharaman et
62 al., 2018; Wegner et al., 2019). These discoveries have led to the assumption that groundwater
63 ecosystems are dominated by chemolithoautotrophic primary production, while other carbon
64 sources, such as the import of surface-derived products of photosynthesis, or ancient organic matter
65 released from sedimentary rocks, play a minor role (Griebler and Lueders, 2009; Akob and Küsel,
66 2011; Schwab et al., 2019). The actual microbial activities and functions of the key players in the
67 groundwater, however, are still unknown. A quantitative assessment of the flow of autochthonous
68 CO₂-derived carbon through the groundwater microbial food web could provide an experimental
69 validation of the dominance of chemolithoautotrophy. Approaches complementary to metagenomics
70 that allow a determination of the active microbial physiological functions and the elucidation of the
71 complex interactions between organisms, to show how the genetic potential is translated into a
72 phenotype, are required to understand ecosystem community functioning.

73 To trace carbon fluxes through microbial communities and identify the active key players, stable
74 isotope probing (SIP) has shown to be a valuable strategy (Neufeld et al., 2007; von Bergen et al.,
75 2013). By introducing ^{13}C -labeled CO_2 to microbial communities, SIP offers the opportunity to
76 illuminate the entire carbon flux from primary production through the food web. However, to
77 pinpoint specific interactions and discern trophic fluxes between individual community members, a
78 highly accurate and time-resolved determination of ^{13}C incorporation in biomolecules must be
79 achieved. Even contemporary nucleic acid-based SIP approaches are typically not able to resolve such
80 slight variations in ^{13}C incorporation patterns (Hungate et al., 2015; Starr et al., 2020). To overcome
81 this problem, we leveraged the high sensitivity of cutting-edge Orbitrap mass spectrometric analysis
82 in SIP-metaproteomics (Taubert et al., 2012) for a highly accurate quantitation of ^{13}C incorporation.
83 To integrate the molecule-specific temporal dynamics in the acquired isotopologue patterns from a
84 ^{13}C -SIP time-series experiment, we developed a PCA-based Stable Isotope Cluster Analysis (SIsCA)
85 approach.

86 We employed this next-generation physiology approach to unravel the role of chemolithoautotrophy
87 for the groundwater microbial community. Groundwater from the Hainich Critical Zone Exploratory
88 (CZE) (Küsel et al., 2016) was supplemented with thiosulfate as electron donor in a $^{13}\text{CO}_2$ SIP
89 microcosm experiment. Thiosulfates are common forms of reduced sulfur in the environment
90 (Grimm et al., 2008), and occur in groundwater through pyrite oxidation (Schippers et al., 1996;
91 Rimstidt and Vaughan, 2003; Kohlhepp et al., 2017). Microbes able to oxidize thiosulfate often
92 possess the genetic potential for autotrophic as well as heterotrophic growth (Ghosh and Dam, 2009;
93 Anantharaman et al., 2018; Wegner et al., 2019), hence, their preferred lifestyle *in situ* is still
94 unknown. By amending the microcosms with thiosulfate, but no organic carbon, we provided ideal
95 conditions for chemolithoautotrophic growth of sulfur oxidizers. We hypothesized that under these
96 conditions, the chemolithoautotrophic activity would be the main source of organic carbon, and a
97 unidirectional carbon flux from autotrophs to heterotrophs would occur in the groundwater
98 microbial community. By mapping the information derived from SIsCA to MAGs, we were able to

99 characterize carbon utilization and trophic interactions of the active autotrophic and heterotrophic
100 key players in the groundwater microbiome over time. This quantitative resolution of the carbon
101 fluxes through a microbial community and the determination of taxon-specific microbial activities has
102 the power to provide novel insights into the fundamental principles governing microbial life.

103

104 **Materials and Methods**

105 **Groundwater sampling & setup of groundwater microcosms**

106 Groundwater was obtained in June 2018 from the Hainich Critical Zone Exploratory (CZE) well H41
107 (51.1150842N, 10.4479713E), accessing an aquifer assemblage in 48 m depth in a trochite limestone
108 stratum. Groundwater from this well is oxic with average dissolved oxygen concentrations of $5.0 \pm$
109 1.5 mg L^{-1} (mean \pm SD), pH 7.2, $< 0.1 \text{ mg L}^{-1}$ ammonium, $1.9 \pm 1.5 \text{ mg L}^{-1}$ dissolved organic carbon, and
110 $70.8 \pm 12.7 \text{ mg L}^{-1}$ total inorganic carbon (Kohlhepp et al., 2017; Schwab et al., 2017). The recharge
111 area of this aquifer assemblage is located in a beech forest (*Fagus sylvatica*). In total, 120 L of
112 groundwater was sampled using a submersible pump (Grundfos MP1, Grundfos, Bjerringbro,
113 Denmark). To collect biomass from the groundwater, 5 L each were filtered through twenty 0.2- μm
114 Supor filters (Pall Corporation, Port Washington, NY, USA). To replace the natural background of
115 inorganic carbon in the groundwater by defined concentrations of ^{12}C or ^{13}C , two times 3 L of filtered
116 groundwater were acidified to pH 4 in 5-L-bottles to remove the bicarbonate present. Following that,
117 ^{12}C - or ^{13}C -bicarbonate was dissolved in the groundwater to a final concentration of 400 mg L^{-1} ,
118 corresponding to 79 mg C L^{-1} . The pH of the groundwater samples was adjusted to 7.2 by addition of
119 ^{12}C - or ^{13}C - CO_2 to the headspace of the bottles.

120 For the ^{13}C -SIP experiment, eighteen microcosms were set up by transferring one filter each into a
121 500-ml-bottle with 300 ml of treated groundwater as described, 9 control microcosms with water
122 containing ^{12}C -bicarbonate, 9 microcosms with water containing ^{13}C -bicarbonate. Additionally, two
123 microcosms were set up by transferring one filter each into a 1-L-bottle with 350 ml untreated
124 groundwater. One of these bottles was supplemented with 150 ml D_2O , the second bottle was
125 supplemented with 150 ml H_2O . To all microcosms, sodium thiosulfate was added to a final
126 concentration of 2.5 mM, and ammonium chloride was added to a final concentration of 15 μM . All
127 microcosms were incubated in the dark at 15 °C with shaking at 100 rpm.

128 **Hydrochemical analyses**

129 During the incubation, in the 18 microcosms supplemented with ¹²C- or ¹³C-bicarbonate,
130 concentrations of oxygen, thiosulfate and sulfate were monitored at regular intervals. Oxygen
131 concentrations were determined using the contactless fiber-optic oxygen sensor system Fibox 4 trace
132 with SP-PSt3-SA23-D5-YOP-US dots (PreSens Precision Sensing GmbH, Regensburg, Germany).
133 Measurements were taken in three ¹²C microcosms and three ¹³C microcosms every two days for the
134 first three weeks, and every week thereafter. Thiosulfate concentration was determined in a
135 colorimetric titration assay using iodine. Samples from all microcosms were measured every 4 to 7
136 days. For each measurement, 1 ml of sample was mixed with 2 mg potassium iodide, then 10 µl of
137 zinc iodide-starch-solution (4 g L⁻¹ starch, 20 g L⁻¹ zinc chloride and 2 g L⁻¹ zinc iodide) and 10 µl of
138 17% (v:v) phosphoric acid were added. Titration was performed using 0.005 N iodine solution in steps
139 of 5 µl until formation of a faint blue color. The concentration of thiosulfate ($C_{thiosulfate}$ in mg L⁻¹) was
140 calculated according to equation (I), where V_{iodine} is the volume of iodine solution added and V_{sample} is
141 the sample volume:

$$142 \quad C_{thiosulfate} = \frac{V_{iodine} \times 561}{V_{sample}} \quad (I)$$

143 The concentration of sulfate was determined using a turbidimetric assay (Tabatabai, 1974) in
144 samples from all microcosms every 4 to 7 days. For each measurement, 1 ml of sample, standard (50
145 µM to 1000 µM potassium sulfate) or blank (dH₂O) was mixed with 0.4 ml 0.5 M HCl and 0.2 ml BaCl₂-
146 gelatin reagent (0.5 g gelatin and 8 g BaCl₂ in 200 ml dH₂O). After incubation for 1 h in the dark, the
147 absorbance at 420 nm was measured in a DR3900 spectrophotometer (HACH, Düsseldorf, Germany).

148 **Detection of cellular activity by Raman microspectroscopy**

149 The two microcosms supplemented with D₂O or H₂O were sampled regularly during the first 7 weeks
150 of incubation for single cell Raman analysis, to quantify the incorporation of deuterium in the
151 biomolecules of active cells as carbon-deuterium-(C-D)-bonds. For Raman spectroscopic analysis, 1
152 ml sample was filtered through a 0.5-µm-filter and subsequently washed three times by centrifuging

153 at 10,000g for 2 min (Hettich Rotina 380R), discarding the supernatant and re-suspending the pellet
154 in 1 ml of sterile ddH₂O. Afterwards, the pellet was re-suspended in 50 µl ddH₂O, and 10 µl of the
155 final bacterial suspension were placed on a nickel foil (Raman substrate) and air-dried at room
156 temperature. Microbial cells were located by dark field microscopy. The measurements were
157 performed with a Raman microscope (BioParticleExplorer 0.5, rap.ID Particle Systems GmbH) with an
158 excitation wavelength of 532 nm from a solid-state frequency-doubled Nd:YAG module (Cobolt
159 Samba 25 mW) and a laser power of 13 mW at the sample. The laser light was focused with an x100
160 objective (Olympus MPLFLN 100xBD) and with a spot size <1 µm laterally. The 180°-backscattered
161 light was diffracted by a single-stage monochromator (Horiba Jobin Yvon HE 532) with a 920 line mm⁻¹
162 grating and then registered with a thermoelectrically cooled CCD camera (Andor DV401-BV)
163 resulting in a spectral resolution of about 8 cm⁻¹. An integration time of 5 s was used per Raman
164 spectrum (-57 to 3203 cm⁻¹).

165 **Pre-processing and analysis of Raman data**

166 The data pre-processing and statistical analysis were carried out in the software GNU R (R Core
167 Team, 2019). First, cosmic spikes were detected and removed from the spectra (Ryabchykov et al.,
168 2016). Subsequently, a wavenumber calibration was applied using 4-acetamidophenol standard
169 spectra (Dörfer et al., 2011), while an intensity calibration was performed using the SRM2242
170 standard material (Bocklitz et al., 2015; Guo et al., 2017). The fluorescence contribution was
171 removed from the spectra using the asymmetric least-squares (ALS) baseline correction method
172 (Liland et al., 2020). Finally, the spectra were vector normalized and subjected to dimensionality
173 reduction using principal component analysis (PCA). Five principal components were used to build a
174 linear discriminant analysis (LDA) classification model to differentiate between deuterium labeled
175 and non-labeled bacterial cells. The deuterium uptake was expressed as the C-D ratio $A(\text{C-D}) / [A(\text{C-D})$
176 $+ A(\text{C-H})]$, which was calculated by integrating the areas of the C-H (2800 - 3100 cm⁻¹) and C-D (2040 -
177 2300 cm⁻¹) stretching vibration bands. The detection and quantification of deuterium incorporation

178 from D₂O in microbial cells was used to assess metabolic activity and decide on time points for
179 sampling of microcosms.

180 **Sampling, DNA and protein extraction**

181 After 21, 43 and 70 days of incubation, microbial biomass from the microcosms was recovered by
182 filtration of the water phase through 0.2- μ m Supor filters (Pall Corporation). The filters used for
183 biomass enrichment before incubation were retrieved from the microcosms as well and combined
184 with the filters used for the water phase. Combined DNA and protein extraction was performed using
185 a phenol/chloroform-based protocol as previously described (Taubert et al., 2018).

186 **Amplicon sequencing**

187 For taxonomic characterization of the bacterial community in the microcosms, amplicon sequencing
188 of the bacterial 16S rRNA gene, region V3 to V5, was done. Polymerase chain reaction was performed
189 using primer pair Bact_341F/Bact_805R (Herlemann et al., 2011) and HotStarTaq Mastermix (Qiagen,
190 Hilden, Germany) as described previously (Kumar et al., 2018). Amplicons were purified using
191 NucleoSpin Gel & PCR Clean-Up Kit (Macherey-Nagel, Düren, Germany). The NEBNext Ultra DNA
192 Library Prep Kit for Illumina (New England Biolabs, Frankfurt, Germany) was used to prepare libraries
193 for amplicon sequencing, following the manufacturer's instructions. Amplicons were purified using
194 AMPure XP Beads (Beckman Coulter, Krefeld, Germany), and amplicon sequencing was then carried
195 out in-house on a MiSeq Illumina platform (Illumina, Eindhoven, The Netherlands) with v3 chemistry.
196 Raw sequence data was analyzed using mothur (v.1.39) (Schloss et al., 2009), according to the
197 mothur standard operating procedures (Kozich et al., 2013) as previously described (Taubert et al.,
198 2018). OTU binning with a 3% identity cutoff was performed, followed by OUT classification using the
199 SILVA reference database release SSU 132 (Quast et al., 2013). Raw Illumina MiSeq sequencing data
200 have been deposited in the Sequence Read Archive (SRA) of NCBI under accession numbers
201 SRR11805453 to SRR11805472.

202 **Metagenomic analysis**

203 To obtain genomes of the key organisms in the microcosms, metagenomic sequencing was
204 performed on DNA samples from 4 selected ^{12}C microcosms, one replicate after 21 days of
205 incubation and 43 days of incubation, and two replicates after 70 days of incubation. These samples
206 were selected to cover the majority of taxonomic diversity, based on principal component analysis of
207 the 16S rRNA gene amplicon sequencing data. DNA sizing, quantitation, integrity, and purity were
208 determined using the Agilent 2100 Bioanalyzer system (Santa Clara, CA, USA). Library preparation
209 was performed with a NEBNext Ultra II DNA Lib Prep Kit (New England Biolabs, Ipswich, MA, USA) as
210 recommended by the manufacturer, followed by multiplexed sequencing on one flow cell of an
211 Illumina NextSeq 500 system (300 cycles) to generate 150-bp paired-end reads.

212 Raw sequencing data was quality filtered using BBDuk (Bushnell, 2014) following assembly using
213 metaSPAdes v3.13.0 (Nurk et al., 2017). Using contigs with more than 1000 bp, three different
214 binning algorithms with default parameters were used to obtain genomic bins: MaxBin 2.0 v2.2.7
215 (Wu et al., 2016), MetaBAT 2 v2.12.1 (Kang et al., 2019), and BinSanity v0.2.7 (Graham et al., 2017).
216 Bin refinement was performed using the MetaWRAP pipeline v1.1.3 (Uritskiy et al., 2018). Only bins
217 with more than 50% completeness and less than 10% contamination were considered. Bins were
218 classified using GTDB-Tk v0.3.2 (Chaumeil et al., 2019) and completeness parameters were assessed
219 using CheckM v1.0.12 (Parks et al., 2015). Bins from different samples were dereplicated using
220 FastANI v1.0 (Jain et al., 2018). The Prokka pipeline v1.13.3 (Seemann, 2014) was used for calling and
221 functional annotation of gene sequences, and for translation into amino acid sequences for
222 metaproteomics analysis. Selected bins of key organisms covered by the metaproteomics analysis
223 were manually refined using Anvi'o v6.1 (Eren et al., 2015) to obtain the final metagenome-
224 assembled genomes (MAGs). Normalized coverage values for all MAGs were calculated by dividing
225 raw coverage values by the relative abundance of *rpoB* genes of each metagenome. The *rpoB* gene
226 abundances were determined using ROcker with the precomputed model (Orellana et al., 2017).
227 Raw Illumina NextSeq metagenome sequencing data have been deposited in the Sequence Read

228 Archive (SRA) of NCBI under accession numbers SRR12964611 to SRR12964614. Curated MAGs have
229 been deposited in the Whole Genome Shotgun (WGS) database of NCBI under accession numbers
230 JADMJD000000000 to JADMKH000000000 as well as under BioSample accessions SAMN16635721 to
231 SAMN16635751.

232 **Metaproteomics analysis**

233 Proteins extracted from the microcosms were first subjected to SDS polyacrylamide gel
234 electrophoresis, followed by in-gel tryptic cleavage as previously described (Taubert et al., 2018).
235 After reconstitution in 0.1% formic acid (v:v), LC-MS/MS analysis was performed on a Q Exactive HF
236 instrument (Thermo Fisher Scientific, Waltham, MA, USA) equipped with a TriVersa NanoMate
237 source (Advion Ltd., Ithaca, NY, USA) in LC chip coupling mode. Raw data files were analyzed with
238 Proteome Discoverer (v1.4.1.14, Thermo Fisher Scientific, Waltham, MA, USA) using the Sequest HT
239 search algorithm. Amino acid sequences derived from translation of the genes from all contigs of the
240 metagenomes were used as reference database for protein identification, to allow the assignment of
241 identified proteins to MAGs. Within Proteome Discoverer, the enzyme specificity was set to trypsin,
242 with two missed cleavages allowed. A 5 ppm peptide ion tolerance and a 0.05 Da MS/MS tolerance
243 were used. Oxidation (methionine) and carbamidomethylation (cysteine) were selected as
244 modifications. Peptides were considered identified when they scored a q-value < 1% based on a
245 decoy database and when their peptide rank was 1. Functional classification of peptides was done
246 according to the gene annotation by Prokka, and taxonomic classification based on the dereplicated
247 and refined MAGs described above.

248 **Determination of ¹³C incorporation in peptides**

249 To quantify the incorporation of ¹³C, peptide identifications from samples of unlabeled microcosms
250 were mapped to mass spectra of the corresponding ¹³C-labeled samples, by comparison of expected
251 peptide mass, chromatographic retention time and MS/MS fragmentation pattern. The molecular
252 masses of peptides were calculated based on their amino acid sequence. Isotopologue patterns of

253 labeled peptides were extracted manually from mass spectral data using the Xcalibur Qual Browser
254 (v3.0.63, Thermo Fisher Scientific, Waltham, MA, USA), and the ^{13}C incorporation was calculated as
255 previously described (Taubert et al., 2011).

256 To determine and compare carbon utilization of the organisms represented by the MAGs, Stable
257 Isotope Cluster Analysis (SIsCA) was performed using R (R Core Team, 2019). Measured isotopologue
258 patterns were compared with 21 predicted isotopologue patterns with varying ^{13}C relative isotope
259 abundance (RIA) in 5% intervals, from 0% to 100% ^{13}C RIA. To determine the most probable RIA, the
260 coefficient of determination (R^2) for each comparison was calculated. We expected that microbes
261 fixing CO_2 would exhibit a high (> 90%) ^{13}C RIA, while microbes utilizing organic carbon should feature
262 a lower ^{13}C RIA with an increasing trend over time. To differentiate lifestyles of different microbes, R^2
263 values were averaged between samples obtained from replicate microcosms and peptides assigned
264 to the same MAG. The resulting datasets of 21 R^2 values per time point for each MAG were
265 compared using principal component analysis in the software package vegan (Dixon, 2003).

266 Generation times of individual taxa were calculated by comparing the relative intensity of unlabeled
267 and labeled peptide signals in mass spectrometric data as previously described (Taubert et al., 2012).
268 First, the number of doubling was calculated according to equation (II), where n is the number of
269 doublings, $I_{12\text{C}}$ is the signal intensity of the unlabeled peptide and $I_{13\text{C}}$ is the signal intensity of the
270 labeled peptide.

$$271 \quad n = \log_2 \frac{I_{12\text{C}} + I_{13\text{C}}}{I_{13\text{C}}} \quad (\text{II})$$

272 If the mass spectrometric signals of labeled and unlabeled peptides were overlapping, the
273 monoisotopic peak was used to determine the total abundance of unlabeled peptide based on the
274 natural distribution of heavy isotopes, as previously described (Taubert et al., 2011). The generation
275 time t_d was subsequently calculated following equation (III), with Δt being the incubation time.

$$276 \quad t_d = \frac{\Delta t}{n} \quad (\text{III})$$

277 **Assessment of DNA density shifts**

278 To evaluate the ^{13}C incorporation on nucleic acid level, DNA samples extracted from the triplicate
279 groundwater microcosms supplemented with ^{13}C and ^{12}C bicarbonate after 21 and 70 days of
280 incubation were subjected to SIP ultracentrifugation in CsCl gradients as previously described
281 (Neufeld et al., 2007). Ultracentrifugation was carried out in an NVT 90 rotor (Beckman Coulter, Brea,
282 CA, USA) in a Sorvall Discovery 90SE ultracentrifuge (Thermo Fisher Scientific, Waltham, MA, USA) at
283 40,900 rpm and 20 °C for 60 h. The gradients were separated into 12 to 14 fractions, covering a
284 buoyant density range from 1.77 g ml $^{-1}$ to 1.68 g ml $^{-1}$. The buoyant density of each fraction was
285 determined using a Reichert AR200 digital refractometer (Reichert Analytical Instruments, Depew,
286 NY, USA). To account for tube-to-tube and spin-to-spin variations of the density gradient (Sieradzki et
287 al., 2020), gradients were normalized by centering the density of ^{12}C DNA onto 1.70 g ml $^{-1}$. The DNA
288 in the density fractions was purified by precipitation with NaCl-PEG as previously described (Taubert
289 et al., 2019) and quantified fluorometrically using the Qubit dsDNA broad-range assays (Thermo
290 Fisher Scientific). Subsequently, amplicon sequencing of the bacterial 16S rRNA gene was performed
291 as described above. OTU-wise DNA buoyant density profiles over the gradients were obtained as
292 previously described (Taubert et al., 2017). Only OTUs that were represented by at least 10 reads in
293 one fraction of each ^{12}C and each ^{13}C replicate were included in the analysis. This was done
294 separately for both time points. From the DNA buoyant density profiles of each OTU, the DNA
295 density of that OTU ($\bar{\rho}_{OTU}$) was determined by performing least-squares regression of the OTU
296 abundance A_{OTU} per fraction f to the fraction density ρ_f using a nonlinear model (IV), where α^2
297 represents the variance of the OTU DNA density.

$$298 \quad A_{OTU,f} = e^{-\frac{(\rho_f - \bar{\rho}_{OTU})^2}{2\alpha^2}} \quad (IV)$$

299 The average ^{12}C and ^{13}C OTU DNA density was determined as the mean of the respective triplicates.
300 The density shift between ^{12}C and ^{13}C samples was determined separately for each time point. The
301 significance of the shift was assessed using Student's t -test based on the respective triplicates.

302 Results

303 Sulfur oxidation by active groundwater microbes

304 The addition of thiosulfate as an electron donor to the groundwater microcosms was expected to
305 result in an activation of chemolithoautotrophic microbes. To determine the onset of metabolic
306 activity, single-cell Raman microspectroscopic analysis was performed in real time on the microcosm
307 labeled with D₂O. The majority of single cell Raman spectra acquired showed a distinct C-D-band at a
308 wavelength position between 2,100 and 2,200 cm⁻¹ within 12 days of incubation (Figure 1). Hence,
309 most cells in the microcosms had become metabolically active, synthesizing new biomolecules and
310 incorporating deuterium from D₂O into new carbon-deuterium bonds. The relative C-D-band
311 intensity increased further over time, from 18.3% after 12 days to 25.7% after 47 days of incubation
312 (median values), indicating continued microbial growth and cross-feeding of deuterium-labeled
313 organic carbon.

314 Thiosulfate consumption rates in the microcosms with ¹²C- and ¹³C-bicarbonate started at 1.7 ± 1.9
315 μmol d⁻¹ during the first three weeks, and increased significantly to 7.2 ± 2.0 μmol d⁻¹ in the final four
316 weeks (*p* = 0.0006, Student's *t*-test) (Figure S1). Oxygen consumption increased significantly from 5.5
317 ± 2.0 μmol d⁻¹ to 12.8 ± 3.2 μmol d⁻¹ in this period (*p* = 0.0013). Sulfate was produced with rates of
318 8.1 to 9.6 μmol d⁻¹ during the duration of the experiment. The approximated stoichiometry observed
319 over the entire time course of incubation was 2.8:1:2.6 (oxygen:thiosulfate:sulfate), close to the
320 theoretical values of 2:1:2 for oxygen-dependent thiosulfate oxidation. The hydrochemical analyses
321 thus demonstrated sulfur oxidation in the microcosms, providing evidence for
322 chemolithoautotrophic primary production. No differences in consumption or formation rates were
323 observed between the ¹²C- and ¹³C-bicarbonate supplemented microcosms, indicating that the
324 isotope treatment had no effect on microbial activity.

325 **Organism-specific ^{13}C incorporation reveals distinct lifestyles**

326 To identify the microbial key players responsible for chemolithoautotrophic primary production, we
327 performed metagenomic sequencing of DNA from selected microcosms after 21, 43 and 70 days to
328 obtain MAGs of the groundwater organisms. We then conducted genome-resolved SIP-
329 metaproteomics based on these MAGs using all ^{12}C - and ^{13}C -bicarbonate-supplemented microcosms.
330 Our newly developed SISCA (Stable Isotope Cluster Analysis) approach revealed five clusters of MAGs
331 with distinct carbon utilization patterns among the 31 most abundant MAGs (Figure 2, Figure S2).
332 The organisms represented by the MAGs in cluster I exhibited a ^{13}C RIA of 95% during the entire 70
333 days of the experiment (Figure 2). This stable and high ^{13}C incorporation can have only resulted from
334 exclusive assimilation of CO_2 , and hence these organisms, related to *Thiobacillus*
335 (*Betaproteobacteriales*), displayed a strictly autotrophic lifestyle. The generation times of these strict
336 autotrophs were shorter than our two-day detection limit within the first 21 days of incubation
337 (Figure 3). Hence, the organisms of cluster I were rapidly producing new, ^{13}C -labeled biomass from
338 $^{13}\text{CO}_2$.

339 Organisms in cluster II displayed a ^{13}C RIA of 65% on average after 21 days of incubation (Figure 2).
340 After 43 and 70 days, their ^{13}C RIA had increased substantially to 91%. The intermediate ^{13}C
341 incorporation at the first time point suggested that these organisms were using organic carbon
342 produced by the autotrophs from cluster I, in addition to organic carbon from the groundwater.
343 Potentially due to limitation of this carbon, they switched to chemolithoautotrophic growth later on.
344 The members of cluster II displaying such a mixotrophic lifestyle were related to *Methyloversatilis*,
345 *Polaromonas* and *Dechloromonas* (all *Betaproteobacteriales*), and showed generation times from 2
346 to 4 days (Figure 3).

347 In cluster III, average ^{13}C RIAs increased from 65 to 76% and in cluster IV from 18 to 53% from 21 to
348 70 days of incubation (Figure 2). The intermediate levels of ^{13}C incorporation indicated the
349 assimilation of organic carbon from chemolithoautotrophic primary production of clusters I and II.

350 The observed increase in ^{13}C RIA reflects the increasing labeling of the organic carbon in the
351 microcosms due to $^{13}\text{CO}_2$ fixation. Variations in the ^{13}C RIAs between species suggested different
352 extents of cross-feeding on the chemolithoautotrophically produced organic carbon, potentially due
353 to preferences of different organic carbon compounds. Most of the organisms in these clusters,
354 related to various *Betaproteobacteriales* and other bacterial groups, exhibited generation times of 3
355 to 4 days (Figure 3). Members of cluster III affiliated with *Hydrogenophaga*, *Vitreoscilla* and
356 *Rubrivivax*, however, showed faster growth, comparable with cluster I. In cluster V, average ^{13}C RIAs
357 reached 6% after 21 days of incubation and did not change at the later time points, suggesting that
358 the organisms were active only in this first period, where they displayed a heterotrophic lifestyle.
359 Generation times in this period were slightly longer and displayed more variability than for clusters III
360 and IV, ranging from 3.5 days for one *Acidivorax* species to eight days for *Aquabacterium* (Figure 3).
361 Approximately 43% of carbon in the microbial community was replaced by ^{13}C after 21 days of
362 incubation, based on the corresponding peptide RIAs of all analyzed MAGs. This portion increased
363 further to 68% after 43 days and 80% after 70 days of incubation. This increase of ^{13}C incorporation
364 across the entire microbiome demonstrated the extensive cross-feeding and recycling of organic
365 carbon. SisCA revealed the carbon flux from autotrophic cluster I to mixotrophic cluster II. On both of
366 these clusters, further cross-feeding by the heterotrophic cluster III to V occurred. Recycling of
367 organic carbon likely played a role within cluster II as well as among the heterotrophic clusters. In
368 agreement, the number of ^{13}C -labeled taxa increased from 21 OTUs after 21 days of incubation to 65
369 OTUs after 70 days of incubation, observable by a significant shift of DNA buoyant density (Figure
370 S3). While after 21 days, these OTUs were mainly affiliated with *Betaproteobacteriales* such as
371 *Thiobacillus*, *Hydrogenophaga* and *Polaromonas*, after 70 days, various other *Alpha*- and
372 *Gammaproteobacteria* were included. The average buoyant density shift of these OTUs likewise
373 increased significantly from $0.021 \pm 0.010 \text{ g ml}^{-1}$ to $0.028 \pm 0.013 \text{ g ml}^{-1}$ ($p = 0.017$, Student's *t*-test) in
374 this period. This highlights the increasing role of $^{13}\text{CO}_2$ -derived carbon introduced by

375 chemolithoautotrophic activity into the microbial carbon pool in the groundwater incubations, and
376 the flux of ^{13}C through the microbial food web.

377 **Functional characterization of MAGs reveals putative mixotrophs**

378 All putative autotrophs detected employed the Calvin-Benson-Bassham (CBB) cycle for CO_2 fixation
379 (Figure 4). Subunits of the key enzyme ribulose-1,5-bisphosphate carboxylase/oxygenase (RuBisCO)
380 were detected for 15 of the 31 MAGs. In 14 of these, further enzymes of the CBB cycle were present.
381 Interestingly, proteins of the CBB cycle were not only present in strict or facultative autotrophs
382 related to *Thiobacillus* from cluster I or *Methyloversatilis*, *Polaromonas* and *Dechloromonas* from
383 cluster II, but also in heterotrophic organisms affiliated with genera *Hydrogenophaga*, *Rhodofera*,
384 *Paucibacter* and *Rubrivivax*, from cluster III and IV. Hence, the ability for mixotrophic growth was
385 present in a large number of organisms in the groundwater microcosms also beyond cluster II. No
386 enzymes allowing a complete reconstruction of other CO_2 fixation pathways, like the Wood-
387 Ljungdahl pathway (no carbon monoxide dehydrogenase/acetyl-CoA synthase (CODH/ACS) or no
388 accessory enzymes), the reductive tricarboxylic acid cycle, or the 3-hydroxypropanoate cycle, were
389 detected.

390 **Functional composition of the whole microbial community**

391 Mapping the functional information obtained to the corresponding taxa in 16S rRNA gene profiles
392 allowed us to classify the lifestyle of up to 50% of the microbial community. Only $3.2 \pm 3.1\%$ (mean \pm
393 sd) of the community were composed of strict autotrophs, primarily affiliated with *Thiobacillus*.
394 Mixotrophs comprised $17.6 \pm 4.3\%$ of the total community and were dominated by *Rhodofera* and
395 *Hydrogenophaga*, but the largest fraction of the total community, with $20.1 \pm 7.8\%$, consisted of
396 heterotrophs, primarily affiliated to *Sediminibacterium* (*Bacteroidetes*), *Pseudomonas*
397 (*Gamma*proteobacteria), *Sericytochromatia* (*Cyanobacteria*) and *Microbacterium* (*Actinobacteria*)
398 (Figure S4, Figure S5).

399 Based on the metagenomics dataset, cluster I containing the strict autotrophs had the lowest
400 normalized coverage, making up 11% of all five clusters. The normalized coverage of the mixotrophic
401 cluster II and the heterotrophic/mixotrophic cluster III were more than twice as high, with 26% and
402 28% of all clusters, respectively. The clusters IV and V also had higher normalized coverages than
403 cluster I, representing 20% and 15% of all clusters, respectively. This demonstrates that the strict
404 autotrophs represented a minor part of the microbial community, suggesting that mixotrophs were
405 of higher importance for the fixation of CO₂ in the microcosms.

406 **MAGs express pathways for the utilization of reduced sulfur compounds**

407 Sixteen of the MAGs obtained were expressing proteins for sulfur oxidation, belonging to the Sox or
408 the Dsr enzyme system (Figure 4). Organisms from clusters II to IV, affiliated with the genera
409 *Methyloversatilis*, *Dechloromonas*, *Hydrogenophaga*, *Rhodofera* and other *Betaproteobacteriales*,
410 were exclusively using the Sox system, which allows the oxidation of both the sulfane group and the
411 sulfone group of thiosulfate to sulfate, without any free intermediates (Ghosh and Dam, 2009). The
412 corresponding MAGs exhibited gene clusters sharing a conserved *soxCDYZAXB* gene order (Figure
413 S6), featuring the core components of the Kelly-Friedrich pathway (Kelly et al., 1997; Friedrich et al.,
414 2001). Further genes, including *soxVW*, *soxEF*, *soxTRS*, and *soxH*, were typically found in different
415 regions of the MAGs, showing variable arrangements. Hence, the clusters differ from the canonical
416 operon structure *soxVWXYZABCDEFGH* described for the *Alphaproteobacterium Paracoccus*
417 *pantotrophus* (Friedrich et al., 2001; Ghosh and Dam, 2009). Organisms from cluster I, affiliated with
418 *Thiobacillus*, were found to produce enzymes from both the Sox and the Dsr system. The
419 corresponding MAGs contained a truncated *soxXYZAB* gene cluster, lacking the genes *soxCD* required
420 for oxidation of the sulfane group of thiosulfate, and hence likely utilized the branched thiosulfate
421 oxidation pathway. In this pathway, the reverse activity of the Dsr enzyme system is used to oxidize
422 the sulfane-derived sulfur atom to sulfite, via elemental sulfur as an intermediate (Ghosh and Dam,
423 2009). The MAGs in cluster I contained the conserved operon structure *dsrABEFHCMKLJOPNR*, thus
424 including the genes *dsrEFH* and *dsrL* described to be typical for sulfur-oxidizers, but lacking the gene

425 *dsrD* indicative of sulfate-reducers (Grimm et al., 2008). In addition, these organisms also expressed
426 the *aprAB* and *sat* genes encoding Adenosine-5'-phosphosulfate reductase and ATP sulfurylase,
427 which can likewise act in reverse to oxidize sulfite to sulfate (Beller et al., 2006). Hence, facultative
428 chemolithoautotrophs in the groundwater microcosms used the Sox system to oxidize thiosulfate to
429 sulfate, whereas obligate chemolithoautotrophs seemed to have utilized an incomplete version of
430 the Sox system to oxidize the sulfone group and the Dsr/Apr/Sat system to oxidize the sulfane group
431 of thiosulfate.

432 **Sulfur oxidizers can use different electron acceptors and donors**

433 For 15 of the MAGs of sulfur oxidizers, cytochrome c oxidase and other enzymes of the respiratory
434 chain were detected, and for 12 of them, enzymes for nitrate reduction (nitrate reductase, nitrite
435 reductase, and nitric oxide reductase) were found (Figure 4). Multiple organisms such as
436 *Dechloromonas* and *Rhodoferrax* had both pathways expressed simultaneously. Proteins for ammonia
437 oxidation, ammonia monooxygenase and hydroxylamine oxidoreductase, were produced by various
438 organisms in clusters I to IV, such as *Thiobacillus* and *Methyloversatilis*. Furthermore, MAG_77
439 (*Thiobacillus*), MAG_55 (*Dechloromonas*) and MAG_7 (*Hydrogenophaga*) expressed [NiFe]-
440 hydrogenase genes. This indicates that the sulfur oxidizers were able to perform both aerobic
441 respiration and denitrification, and that also nitrogen compounds and hydrogen could be used as
442 electron donor in the microcosms.

443 **MAGs salvage a variety of organic carbon sources**

444 For organic carbon utilization, various pathways, targeting simple sugars (Glycolysis, pentose
445 phosphate pathway), amino acids (TCA cycle), fatty acids (beta-oxidation), C₁ compounds and
446 aromatic compounds, were found throughout clusters II to V (Figure 4). The obligate autotrophs in
447 cluster I only expressed pathways for degradation of simple sugars. The TCA cycle was present as one
448 of the most abundant metabolic modules in the organisms of clusters II to V. Pathways for toluene
449 and ethylbenzene utilization were expressed by organisms affiliated with *Dechloromonas* and

450 *Rhizobacter* (*Betaproteobacteriales*), respectively. Enzymes for naphthalene and catechol utilization
451 were detected, e.g., for organisms related to *Hydrogenophaga* and *Pseudomonas*. The utilization of
452 C₁ compounds was primarily detected for organisms related to *Methyloversatilis*, which possessed
453 methanol dehydrogenase, formate dehydrogenase and various enzymes for
454 tetrahydromethanopterin-dependent C₁-cycling. Enzymes for degradation of complex carbohydrates
455 such as starch and chitin were produced by organisms related to *Microbacterium* and
456 *Sediminibacterium*. Reflecting this high versatility for organic carbon utilization, a wide range of
457 proteins with import functions was observed (Figure 4). In mixotrophs and heterotrophs from
458 clusters II to V, peptides of import systems for amino acids and carboxylic acids (e.g. alpha-keto acids,
459 C₄-dicarboxylates, lactate) were highly abundant. The microorganisms from clusters III to V, which
460 were exclusively growing heterotrophically in our incubations, possessed further transport systems
461 for carbohydrates and nucleotides, and showed the broadest variety of import functions. For the
462 obligate autotrophs in cluster I, only transporters targeting cations (mostly iron transporters) and
463 phosphate were found. This suggests an increasing ability to utilize organic carbon from cluster I to
464 clusters III - V.

465 Discussion

466 An understanding of the function of microbial communities in the environment cannot be achieved
467 from genomic data alone, as information about their activity is required. The combination of ¹³C- and
468 D₂O-SIP with genome-resolved metaproteomics and Raman microspectroscopy allowed us to
469 determine the activity of organisms on the level of specific taxa as well as on the level of single cells.
470 SIsCA revealed the carbon flux between the active taxa qualitatively and quantitatively. By comparing
471 the ¹³C incorporation patterns after different periods of incubation, we were able to follow the
472 metabolic processes of the organisms over time (Figure 5). We expected that sulfur-oxidizing
473 chemolithoautotrophs would be activated first. Later on, organic carbon provided through
474 chemolithoautotrophic primary production by this dominant population would be released, for

475 example through dying cells releasing necromass into the water and fuel a heterotrophic satellite
476 community. The genomic makeup of the organisms identified supported this classical separation of
477 autotrophic and heterotrophic lifestyles on first glance. Raman analysis, however, showed that most
478 microbial cells were active in a short period, questioning our concept of a strict temporal and
479 functional separation. We observed that most organisms were combining an autotrophic and a
480 heterotrophic lifestyle to establish individual, optimized strategies for survival, adapted to the
481 conditions they experienced at the time. This gave rise to a high functional diversity in the
482 groundwater microcosms. We could show exactly how the diverse organisms leveraged their
483 genomic potential to come up with strategies to succeed in the groundwater microcosms, moving
484 beyond a separation of different lifestyles based on the genomic potential of organisms.

485 Mixotrophs, and not obligate chemolithoautotrophs, were the most abundant active organisms in
486 our groundwater microcosm. These mixotrophs strongly preferred heterotrophic growth. Their ^{13}C
487 incorporation patterns, raising from 65 up to 91% RIA, clearly demonstrated the uptake of organic
488 carbon. This is astonishing as thiosulfate and oxygen were available in sufficient amounts throughout
489 the whole experiment, providing conditions that should have favored chemolithoautotrophic growth.
490 A heterotrophic lifestyle was likely favored due to the higher metabolic cost of fixing CO_2 via the CBB
491 cycle during autotrophic growth (Berg, 2011). Consequently, CO_2 fixation only occurred when
492 insufficient organic carbon was available. The organisms of cluster II switched from heterotrophic
493 growth to CO_2 fixation at later time points, likely due to such limitations of suitable organic carbon
494 compounds. The only strict autotrophs were the *Thiobacillus*-related organisms from cluster I.
495 *Thiobacillus* species are known to be obligate autotrophs, featuring an incomplete TCA cycle that
496 prevents heterotrophic growth (Boden et al., 2017). Performing thiosulfate and hydrogen-driven
497 denitrification, *Thiobacillus* constituted up to 50% of a previous enrichment culture obtained from
498 the groundwater investigated here (Kumar et al., 2018), but only made up a minor fraction of the
499 total community in our experiment. Nevertheless, cluster I represented the first entry point of CO_2 -
500 derived ^{13}C . Their low abundance despite the short generation times of less than 2 days suggested a

501 high turnover of biomass, allowing rapid cross-feeding by other organisms. Within 21 days, 43% of
502 the microbial biomass was ¹³C-labeled. The further increase of ¹³C-labeling in the microbial biomass,
503 up to 80% after 70 days, was driven through a cascade of recycling reactions by organisms with a
504 preference for heterotrophic growth. Similar continuous recycling of microbial biomass is also
505 characterizing the transport of organic matter from soils (Stevens, 1997; Roth et al., 2019). When no
506 suitable organic carbon was available, these organisms turned to autotrophic growth instead, and
507 further increased the ¹³C-labeling in the community. This opportunistic strategy to use autotrophy to
508 fill up their carbon demand allowed them to dominate the microbial community, with generation
509 times almost as short as the strict autotrophs, with 4 days to less than 2 days. Considering the overall
510 low concentrations but high diversity of organic carbon compounds in the groundwater (Benk et al.,
511 2019; Schwab et al., 2019), such a versatile and adaptive strategy seemed to be superior to focusing
512 on one particular lifestyle. In contrast to organisms restricted to a purely heterotrophic lifestyle in
513 cluster IV and V with generation times of up to 8 days, the ability to use CO₂ fixation as a fill-up
514 function for organic carbon proved highly advantageous in the oligotrophic groundwater.

515 Sulfur oxidation was not restricted to strict chemolithoautotrophs in our groundwater microcosms.
516 Mixotrophs and heterotrophs expressed sulfur oxidation pathways as well. The use of such inorganic
517 electron donors as energy sources is an alternative to the oxidation of organic compounds. Especially
518 when the concentrations of organic carbon present are low, as in pristine groundwater,
519 microorganisms might rather use it as carbon source than to gain energy, as the re-fixation of CO₂ is
520 associated with higher energy costs (Berg, 2011). Under such conditions, the ability to use inorganic
521 electron donors would be an advantageous part of the survival strategy. For various facultative
522 autotrophic taxa from clusters II to IV, including *Polaromonas*, *Dechloromonas*, *Hydrogenophaga*, and
523 *Rhodofera*, the ability for sulfur oxidation has been suggested based on genomic evidence (Kämpfer
524 et al., 2005; Mattes et al., 2008; Salinero et al., 2009; Jin et al., 2020). However,
525 chemolithoautotrophic growth on reduced sulfur compounds has not been observed for these
526 genera in pure culture so far. Members of these taxa frequently occur in ecosystems where the

527 oxidation of thiosulfate plays an important role, e.g., linked to denitrification (Kumar et al., 2017;
528 Kumar et al., 2018; Sun et al., 2020). Our study demonstrates that these organisms can use reduced
529 sulfur compounds as energy source, and *Polaromonas*, *Dechloromonas* and potentially
530 *Hydrogenophaga* used it to fuel autotrophic growth. Interestingly, for most of the observed taxa the
531 growth on H₂ and CO₂ was experimentally demonstrated (Willems et al., 1989; Shrouf et al., 2005;
532 Sizova and Panikov, 2007), and other H₂ oxidizing organisms are known to grow
533 chemolithoautotrophically also on reduced sulfur compounds (Sano et al., 2010). This suggests a
534 close link between the hydrogen and the sulfur cycle, potentially via the production of H₂ from
535 reduced sulfur compounds, in chemolithoautotrophic sulfur oxidizers.

536 The ability to use organic carbon compounds showed an inverse, gradual relationship to the CO₂
537 fixation activity of the organisms in our groundwater microcosms. Organisms with a strictly
538 autotrophic strategy had no transport functions for organic carbon. Their metabolic capabilities were
539 restricted to simple sugars, likely to allow the utilization of organic carbon assimilated via the CBB
540 cycle (Berg, 2011). Organisms in cluster II with a strategy that combined autotrophic and
541 heterotrophic lifestyle possessed importers for amino acids, and highly expressed the TCA cycle to
542 facilitate their assimilation. This suggests that amino acids were central compounds in the carbon
543 transfer between microbes. As amino acids are major components of microbial necromass, it is
544 plausible that lysis of autotrophic cells made these compounds available in the microcosms. This
545 release of amino acids by the strict autotrophs of cluster I was limited, so cluster II adapted its
546 strategy by switching to fixation of CO₂ later on. The success of this strategic mix of
547 chemolithoautotrophy with utilization of easily available organic carbon compounds allowed them to
548 become 2.5-fold as abundant as cluster I. The organisms of cluster III were able to utilize a wider
549 range of organic compounds, including carbohydrates, methyl compounds, nucleotides and aromatic
550 compounds. The strategy of diversifying their organic carbon diet allowed them to become the most
551 abundant cluster in the microcosms, being more independent of autotrophic CO₂ fixation until the
552 end of the experiment. Whether they grew entirely heterotrophic or obtained a minor part of carbon

553 from CO₂ via the expressed CBB cycle is difficult to determine. Also during purely heterotrophic
554 growth, a small part of CO₂-derived carbon would be incorporated into biomass due to anaplerotic
555 reactions (Alonso-Sáez et al., 2010; Spona-Friedl et al., 2020). The strict heterotrophs in cluster IV
556 expressed most of the organic carbon transport and assimilation pathways, allowing them to succeed
557 without autotrophic CO₂ fixation capabilities by relying entirely on the organic carbon in the
558 groundwater microcosms. The organisms in the groundwater microcosms hence relied on various
559 strategies combining chemolithoautotrophy and heterotrophy to different extents, by utilizing no
560 organic carbon, only the most easily available compounds, a diverse range, or as many compounds as
561 possible. The more diverse the metabolic potential of an organism is, the higher the tendency to use
562 organic carbon and the higher the independence from CO₂ fixation. In its entirety, the community
563 was highly efficient in using the diverse carbon compounds available in the groundwater microcosms,
564 despite facing a diverse mixture of organic carbon compounds.

565 Metagenomics studies of multiple groundwater sites suggested a diverse metabolic potential, e.g. for
566 using multiple electron donors and acceptors, to be a common trait of chemolithoautotrophic sulfur
567 oxidizers (Anantharaman et al., 2016; Jewell et al., 2016; Wegner et al., 2019). Based on the
568 strategies observed in our microcosms, sulfur oxidizers should have two main advantages in shallow
569 aquifers of mixed siliciclastic-limestone alternations. (1) Reduced sulfur compounds are released
570 autochthonously from weathering of interspersed pyrite minerals (Schippers et al., 1996; Rimstidt
571 and Vaughan, 2003; Kohlhepp et al., 2017) so the energy source for sulfur oxidizers is independent
572 from surface input dynamics. (2) Sulfur oxidizers can tailor their organic carbon requirements
573 dependent on the availability. Within the subsurface, carbon can be sourced from surface based
574 primary production and translocated with groundwater recharge. However, this surface carbon has
575 typically already been recycled through the microbial loop and incorporated into labile microbial
576 carbon (Roth et al., 2019). Second, microbial cells are transported down from the soils, representing
577 another source of microbial carbon (Herrmann et al., 2019). Third, subsurface organic carbon can
578 also be released from sedimentary rocks, which is often characterized by a higher fraction of

579 aromatic and aliphatic functions, similar to organic carbon from marine sediments rather than soil
580 (Benk et al., 2019). The ability of sulfur oxidizers to access all of these carbon sources, and
581 additionally fix CO₂, should allow them to thrive independent of variations in the organic carbon
582 supply.

583 The majority of the active sulfur oxidizers observed in our groundwater microcosms were affiliated
584 with the class *Gammaproteobacteria*. In the groundwater of well H41, *Gammaproteobacteria* are the
585 second most abundant class after *Cand. Parcubacteria* (Yan et al., 2020). Organisms affiliated with
586 *Betaproteobacteriales* were found to have the highest abundance of RuBisCO-encoding transcripts
587 (Herrmann et al., 2015), highlighting their importance for CO₂ fixation in this well. The rapid
588 replacement of organic carbon by ¹³C in our microcosms likewise suggested a high importance of CO₂
589 fixation for carbon cycling in the groundwater. Indeed, genes for pathways of CO₂ fixation, such as
590 the CBB cycle, are widely distributed: In 85% of 48 groundwater samples obtained from various
591 locations in Germany and Austria, RuBisCO genes were detected (Alfreider et al., 2009), and at the
592 site investigated here, RuBisCO transcripts made up 4% of the total metatranscriptomic reads
593 (Wegner et al., 2019). The recent finding that the metabolic potential for chemolithoautotrophy
594 coincides with the complexity of subsurface food webs (Herrmann et al., 2020) demonstrates that
595 this essential function has consequences for the entire ecosystem.

596 **Conclusions**

597 Stable Isotope Cluster Analysis utilizing high-resolution ¹³C isotopologue patterns of peptides
598 obtained by genome-resolved SIP-metaproteomics allowed us to provide MAG-specific information
599 of CO₂ fixation, organic carbon utilization and activity of the microorganisms in our groundwater
600 model system. We observed that mainly mixotrophs were abundant in the groundwater microcosms,
601 and largely preferred the uptake of organic carbon to CO₂ fixation. Chemolithoautotrophy is an
602 essential function for the recycling of organic carbon, and seems to serve as a “fill-up function” in the
603 groundwater, to replenish organic carbon that has been lost or in situations where not enough

604 organic carbon is available for the organisms. A variety of autotrophic, mixotrophic and
605 heterotrophic organisms were able to utilize reduced sulfur compounds as energy sources. These
606 organisms thus do not rely solely on the oxidation of the precious organic carbon to CO₂ for energy
607 gain, and instead can focus on employing it as carbon source. For this, the groundwater microbes
608 show a high metabolic versatility to utilize as much as possible of the scarce organic carbon
609 compounds present in the oligotrophic groundwater. The presented next generation physiology
610 approach provides information about the metabolic phenotype of individual taxa as well as the
611 carbon fluxes between them. This information is essential for the confirmation of hypotheses
612 established in metagenomics studies and to understand the role of microbial communities in
613 ecosystem functioning.

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851 **Figure and Table legends**

852 **Figure 1: Quantification of deuterium incorporation by single cell Raman microspectroscopy.**

853 Boxplots depict the relative intensity of the Raman C-D-band determined as $A(C-D) / [A(C-D) + A(C-$
854 $H)]$ from single cell Raman spectra. Spectra were obtained from samples of groundwater microcosms
855 with 30% D₂O (shaded) or with H₂O only (empty) at different time points of incubation. Boxes show
856 median, first and third quartile. Whiskers indicate 5th and 95th percentile. Outliers are depicted as
857 dots. At least n=147 spectra were obtained at each time point.

858 **Figure 2: Clustering of selected MAGs based on carbon utilization.** Left panel: Stable isotope cluster

859 analysis based on PCA of ¹³C incorporation profiles over incubation time obtained by SIP-
860 metaproteomics of samples from the ¹³C-microcosms. Each point represents an organism associated
861 with one MAG. Clusters of MAGs are indicated with Latin numbers. Ellipses depict 95% confidence
862 intervals. Only MAGs are shown where at least two replicates of ¹³C incorporation patterns per time
863 point could be acquired. Right panel: Representative ¹³C incorporation profiles of MAGs marked with
864 an asterisk are given for each cluster. Heatmaps depict the extent of ¹³C incorporation in peptides of
865 the corresponding MAG after 21 (T1), 43 (T2) and 70 days (T3) of incubation in 5%-intervals, ranging
866 from 0 to 100% ¹³C relative isotope abundance.

867 **Figure 3: Generation times of groundwater organisms in the microcosms.** Given values were

868 determined for the first 3 weeks of incubation, based on the relative abundance of ¹²C and ¹³C
869 peptides. Shown are mean and standard deviation based on n ≥ 4 replicate determinations. Colored
870 horizontal lines indicate average generation time for each cluster. bdl: generation time was below
871 the detection limit of 2 days. na: no quantification of the generation time was possible.

872 **Figure 4: Metabolic functions of selected MAGs.** The sizes of the bubbles correspond to the total

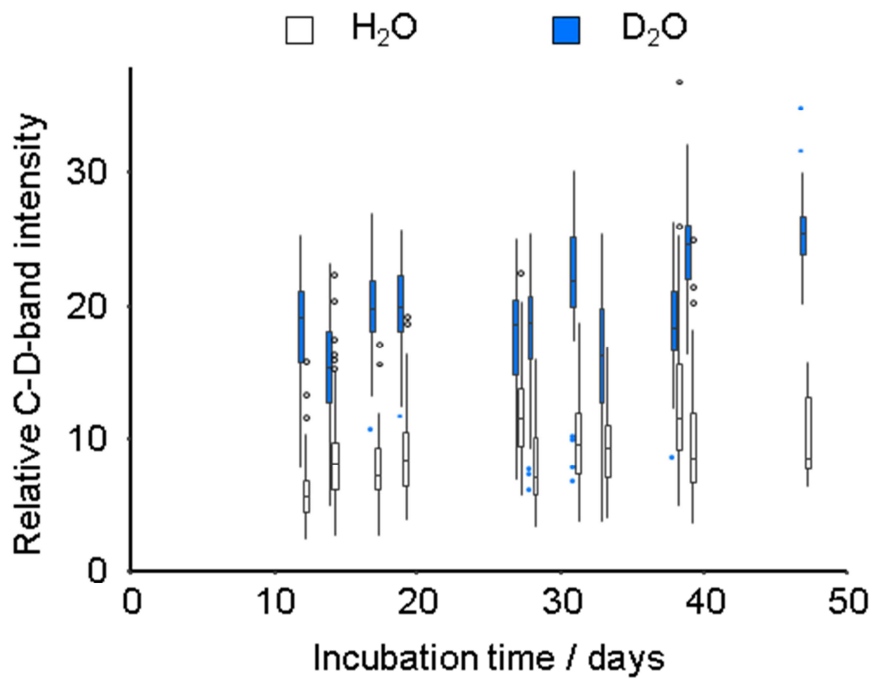
873 number of peptides for each MAG and each functional category identified at any time point.
874 Metabolic functions are grouped into CO₂ fixation (red), sulfur cycling (yellow), nitrogen cycling
875 (green), aerobic respiration and ATP synthesis (blue) and organic carbon utilization (black). The

876 taxonomic categories “others” include peptides that were assigned to multiple MAGs affiliated with
877 the same genus. Only MAGs included in the stable isotope cluster analysis are shown. RuBisCO:
878 ribulose-1,5-bisphosphate carboxylase/oxygenase, CODH/ACS: carbon monoxide
879 dehydrogenase/acetyl-CoA synthase, TCA cycle: tricarboxylic acid cycle.

880 **Figure 5: Carbon flux between microbial clusters.** Red arrow inlays illustrate the portion of $^{13}\text{CO}_2$ -
881 derived carbon assimilated by each microbial cluster after 21, 43 and 70 days. Overall arrow width
882 scales with total amount of carbon assimilated based on the relative abundance of the respective
883 microbial cluster in the metagenomics analysis. Fading grey arrows indicate uptake of unlabeled
884 organic carbon from the groundwater. Checkboxes highlight the presence and activity of metabolic
885 functions for CO_2 fixation, utilization of organic carbon and sulfur oxidation.

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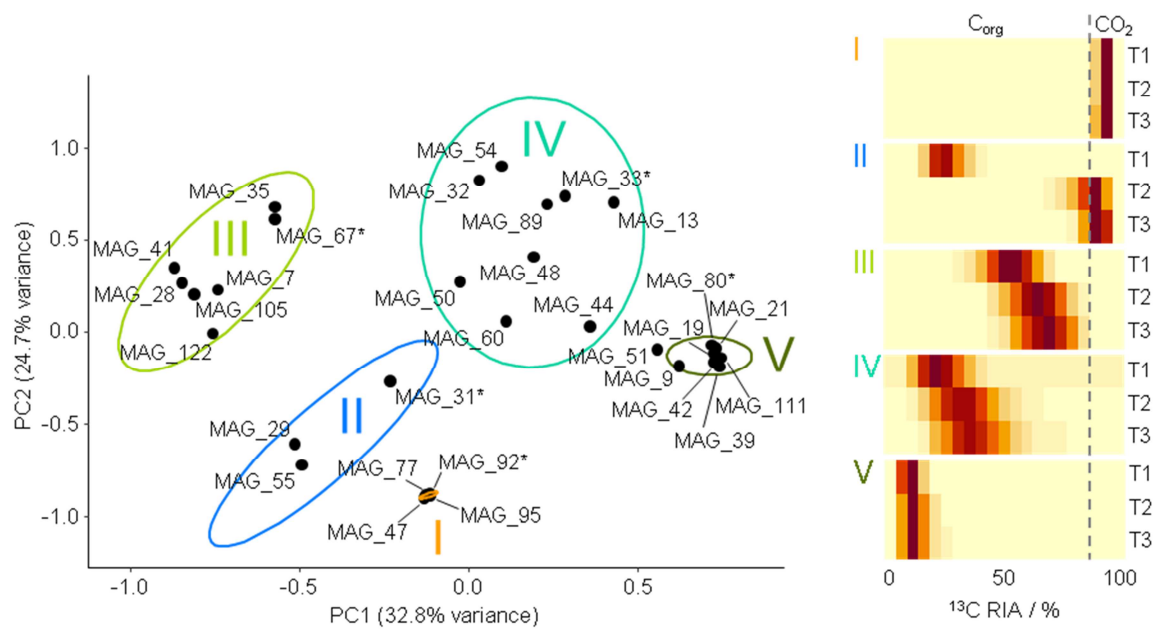
887 **Figure 1**



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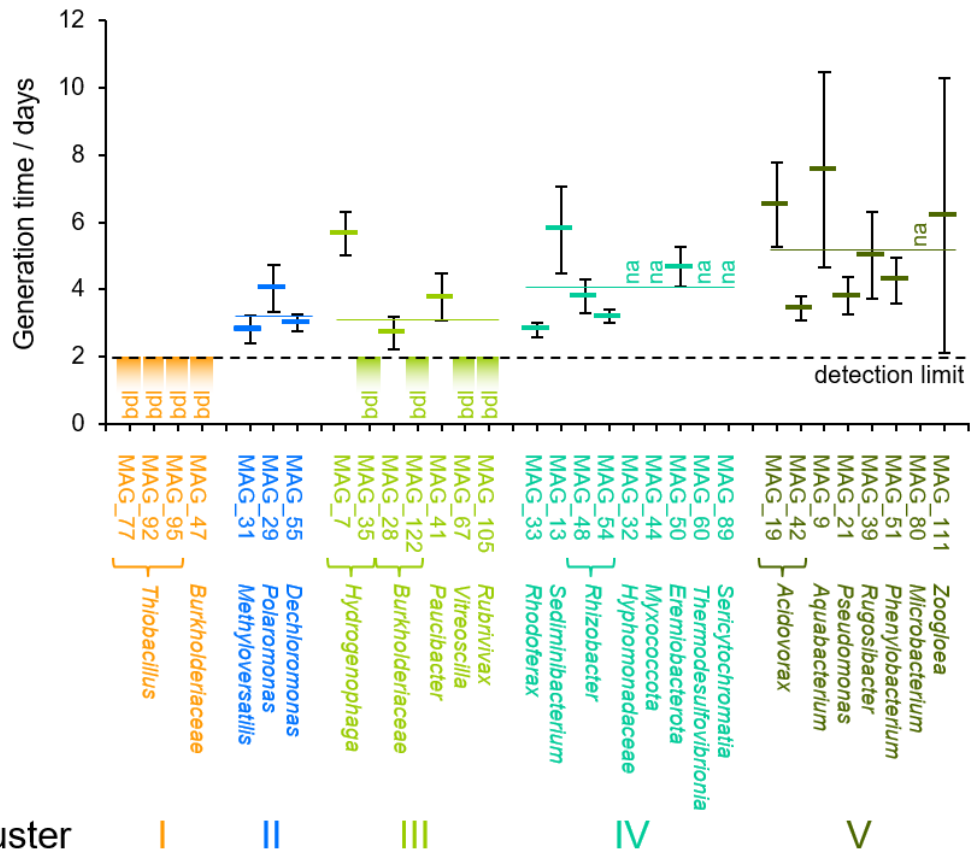
890 **Figure 2**



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893 **Figure 3**



894 cluster

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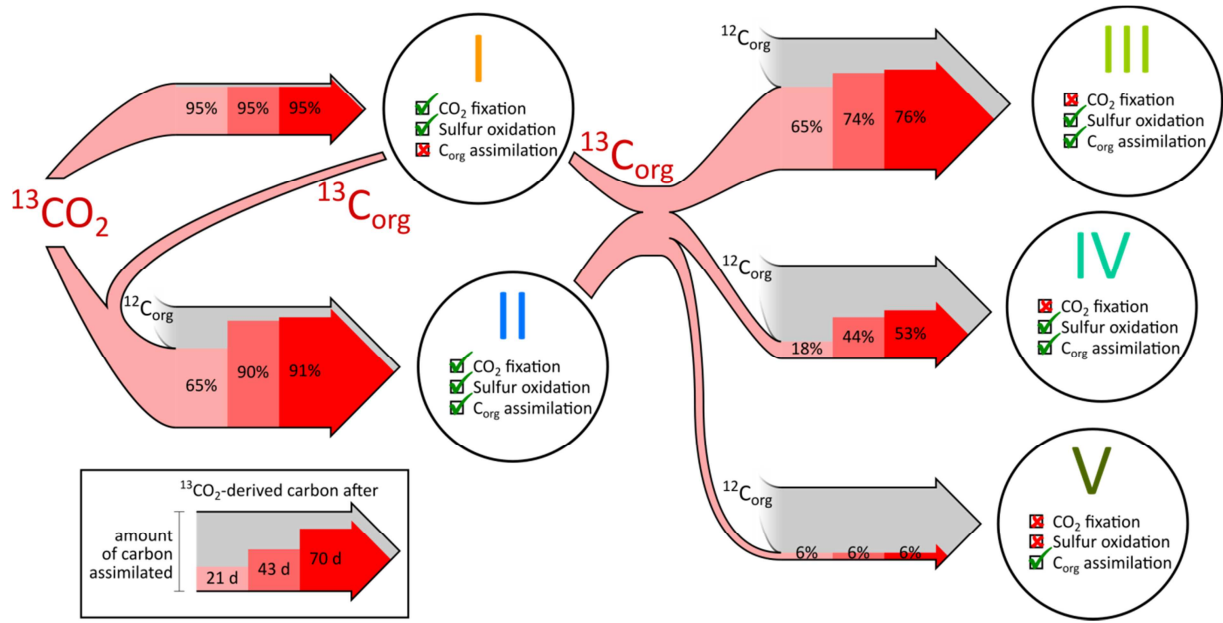
896 **Figure 4**



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899 **Figure 5**



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