

1                   **Rapid shift in microbial community structure in a neutral hydrothermal**  
2                   **hot spring from Costa Rica**

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16   **Keywords:** Hot Spring, Costa Rica, Microbial mat, Thermophiles, Aquificae, *Sulfurihydrogenibium*

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30  
31   **Abstract**

32

33 In this work, we characterize the geochemistry and microbial community of Bajo las Peñas, a neutral (pH  
34 6.5-7.4), hot spring (T = 62.0-68.0°C) located near Turrialba Volcano, Costa Rica. The microbiota at its  
35 two sources belongs mainly to the family Aquificae, comprising OTUs closely related to the genera  
36 *Sulfurihydrogenibium*, *Thermosulfidibacter*, *Thermodesulfovibrio* and *Thermocrinis* which is consistent  
37 with the presence of moderate levels of sulfate (243-284 mg/L) along the stream. We determined a  
38 dramatic shift in the microbial community just a few meters downstream of the sources of the hot spring  
39 (15-20 meters), with a change from sulfur related chemoautotrophic (e.g. *Sulfurihydrogenibium* and an  
40 OTU closely related to *Thermodesulfovibrio*) to chemoheterotrophic prokaryotes (e.g. *Meiothermus*,  
41 *Nitrososphaera*, *Thermoflexus*, *Thermus*). Thus, in this neutral hot spring, the first level of the trophic  
42 chain is associated with photosynthesis as well other anaerobic CO<sub>2</sub> fixing bacteria. Then, thermotolerant  
43 chemoheterotrophic bacteria colonize the environment to degrade organic matter and use fermentative  
44 products from the first level of the trophic chain. Our data demonstrate how quickly the microbial  
45 community of an ecosystem can change in response to environmental variables and sheds light on the  
46 microbial ecology of less common circumneutral pH hot springs.

47

48

## 49 **Introduction**

50

51 Characterization of microbial communities inhabiting hot springs is of great interest for understanding the  
52 ecology in extreme environments. Due to their high temperatures, macroscopic organisms are limited in  
53 these habitats; as a result, hot spring communities are usually made up exclusively by thermotolerant or  
54 thermophilic microorganisms (Akanbi et al. 2019). These conditions define hot springs as “ecologic  
55 islands”, where even when foreign microorganisms are constantly inoculated into the water, they are not  
56 able to survive the high temperatures. Additionally, due to their geological and chemical properties, a  
57 great spectrum of niches exists in this kind of environments, which in consequence allows the  
58 coexistence of multiple species (Mitri 2019). The great diversity present in the multiple niches (microbial  
59 mat, sediments, surface, and underground water) has been proposed to be a consequence of metabolic  
60 heterogeneity that evokes during the community evolution (Mitri 2019). Such a variation of taxa across  
61 different niches of the same hot spring has been observed when analyzing the microbial mat and the  
62 anoxic region of Mushroom Spring in Yellowstone National Park (Klatt et al. 2011; Thiel et al. 2016; Tank  
63 et al 2017). Furthermore, in this hot habitat, it has been shown that specific taxa are associated to  
64 particular metabolic functions within the community and cross-feeding may be occurring explaining the  
65 high diversity in these habitats (Thiel et al. 2017). The understanding of the rules that govern the ecology  
66 of hot springs is relevant because these extreme habitats have also been proposed to be a good model

67 for understanding the evolution of life, particularly because hot springs resemble hydrothermal vents  
68 where life might have emerged (Damer and Deamer 2020).

69

70 Hot springs can be found at various locations around the globe and their physicochemical parameters  
71 vary widely but they are normally acidic. Previous studies have demonstrated that their acidity levels are  
72 highly correlated to the chemical species that are present in the environment, particularly iron and sulfur  
73 species (Xu et al. 1998, 2000; Pierson et al. 1999, 2000; Sharma et al. 2001). A particularity shared  
74 between the less commonly described neutral pH hot springs with the acidic ones is the significant  
75 proportion of ferrous iron produced from mineral dissolution, which can form precipitates (Konhauser  
76 1997; Pierson et al. 1999; Yee et al. 2003). The iron minerals in neutral hot springs are usually formed by  
77 spontaneous iron oxidation, and the oxides produced consequently nucleate, sometimes in association  
78 with silica, and precipitate as ferric hydroxides that include minerals like ferrihydrite, lepidocrocite or  
79 goethite (Peng et al. 2013; Liu et al. 2014; Cooper et al. 2017). Due to their mild pH and moderate toxic  
80 metals concentrations, the main extreme conditioning factor in neutral hot springs is temperature, limiting  
81 the proliferation of multiple microorganisms and directly shaping the community (Ward et al. 1998).

82

83 Photoautotrophic microorganisms play a key role in hot water bodies (Ward and Castenholz 2000; Klatt et  
84 al. 2011). Previous reports have found that members of Cyanobacteria and Chloroflexi are common  
85 phototrophic groups found in this type of habitats, particularly in microbial mats (Pierson and Castenholz  
86 1974; Miller and Castenholz 2000; Alcamán et al. 2015; Prieto-Barajas et al. 2018; Ward et al. 2018). On  
87 the other hand, some deep sulfurous hot springs have a low abundance of cyanobacteria but a high  
88 concentration of sulfur chemoautotrophic bacteria, particularly from the family Desulfobacteriaceae. In  
89 these environments, these microorganisms are the main CO<sub>2</sub>-fixers (Nunoura 2014). In addition to the  
90 importance of autotrophic microorganisms, multiple taxa classified as heterotroph, such as Deinococcus-  
91 Thermus, Nitrospirae, Proteobacteria, Firmicutes and Actinobacteria are also an important part of hot  
92 spring communities (Valverde et al. 2012; Wang et al. 2013; Daims 2014; Sahoo et al. 2015; Ghilamical  
93 et al. 2017; Rozanov et al. 2017). Complete metagenome studies have also revealed that these microbial  
94 communities can have an extremely diverse and complex metabolism, where nutrients can be recycled  
95 between phylotypes and multiple cross-feeding processes may be occurring (Mangrola et al. 2015;  
96 Saxena et al. 2017; Thiel et al. 2017; Pedron et al. 2019).

97

98 Bajo las Peñas is a natural occurring hot spring located in Cartago province, close to Turrialba and Irazu  
99 volcanoes (Fig. 1). This site is a small lotic environment where water from the hot spring source flows

100 slowly through an extensive bacterial mat for about 20 meters and then joins a cold, mountain stream that  
101 eventually becomes a tributary of the Toro Amarillo River. It is characterized for having a reddish  
102 gelatinous microbial mat covering the site (Figs 1C and 1D), a neutral pH (6.5-7.4) and a temperature of  
103 around 65°C (Uribe-Lorio et al. 2019). Study of circumneutral hot springs like Bajo las Peñas is relevant  
104 since they are uncommon in the world (most are acidic) and can potentially become a source of  
105 microorganisms and thermostable enzymes of biotechnological interest. After we carried out our sampling  
106 in 2016 for this study, a survey of the microbiology of several Costa Rican hot springs reported data of the  
107 microbial community inhabiting the microbial mat of Bajo Las Peñas based on samples taken in 2012.  
108 (Uribe-Lorio et al. 2019). This last study reported that the microbial community in the mat is formed  
109 almost exclusively by Cyanobacteria. Here, we show a more extensive study of the geochemistry and  
110 microbial communities that inhabit the water and the sediments in Bajo las Peñas, including samples  
111 taken right at the source in order to better understand the microbial diversity and the possible metabolism  
112 associated with this close-to-neutral pH hot spring.

113

## 114 **Materials and Methods**

115

### 116 *Study Site*

117

118 Bajo las Peñas is located in a private farm northeast of Irazu Volcano (Cartago, Costa Rica; Fig. 1A). At  
119 the site, we actually identified and sampled two springs (sources) where hot water emerges which are  
120 separated by about 15 m (BP1 and BP3; see Fig. 1). The samples were taken along the whole site, which  
121 is approximately 20 meters long and included two more sampling sites away from the sources. Along the  
122 effluent, a very thick (ca. 30-50 cm) reddish microbial mat covers the site and slows down the flow of  
123 water. Interestingly, despite the high temperature in the surroundings of this thermal spring, grass and  
124 bushes were found. The BP1 point (Fig. 1; 9.993492 N, 83.801597 W) corresponds to the origin (first  
125 spring) of Bajo las Peñas. The second sampling point (BP2; 9.993558 N 83.801639 W) is located 10 m  
126 downstream. Five meters further downstream of BP2 (15 m from the first spring) emerges the second  
127 spring, which corresponds to our third sampling point (BP3; 9.993603 N, 83.801677 W). Finally, 5 m  
128 downstream from BP3 (20 meters downstream from BP1), our fourth sampling point (BP4; 9.993658 N  
129 83.801716 W) was selected, just a few meters before Bajo las Peñas mixes with a cold stream. Permits  
130 for sampling were obtained from the Institutional Commission of Biodiversity of the University of Costa  
131 Rica (resolution N° 066) and the owners of the property.

132

133 *Sampling and field measurements*

134

135 On 18 August 2016, samples of water (three samples of 1 Liter each) were collected at four points (BP1  
136 to BP4; see Fig. 1) along this thermal water (i.e. twelve water samples). For this purpose, sterile flasks  
137 were attached to a long rod in order to reach down into the water and avoid contamination. At each  
138 sampling point it was possible to take one or two sediment samples, for a total of seven sediment  
139 samples across the entire spring. Specifically, two sediment samples were taken from sampling site 1  
140 (BP1S1 and BP1S2), two from sites 2 (BP2S1 and BP2S2) and 4 (BP4S1 and BP4S2), whereas a single  
141 sediment (BP3S1) was retrieved from point 3. For sediment sampling, masses between 10 and 50 g were  
142 placed into sterile 50 mL centrifuge tubes. The sediments were taken at depths not exceeding 15 cm and  
143 due to the presence of the microbial mat, the sediment samples contain part of it. In the field, temperature  
144 was measured with a partial immersion thermometer. To measure dissolved oxygen (DO), bottles (1 L)  
145 were filled with water from each sampling point, hermetically sealed and cooled to room temperature (~22  
146 °C). At that temperature the dissolved oxygen was measured with a dissolved oxygen meter Model 550A  
147 (Yellow Springs Instrument Company Inc, Ohio, USA). pH was measured with a pH meter (Scholar 425  
148 pH meter, Corning, Inc., Corning, NY) in the laboratory. Water samples for chemical analysis were  
149 collected in clean glass bottles, chilled on ice, and stored at 4 °C until analysis. Water samples for  
150 microbial community analysis were collected in clean and sterile bottles, stored at 4 °C and processed  
151 within 24 h.

152

153 *Chemical Analysis*

154

155 For chemical analyses, water samples were filtered through polycarbonate membrane filters (0.45 µm;  
156 Sartorius 23006-47N) before analysis. Major anionic components (Cl<sup>-</sup>, F<sup>-</sup>, NO<sub>3</sub><sup>-</sup>, SO<sub>4</sub><sup>2-</sup>) in the water  
157 samples were analyzed by ion-exchange chromatography (IC, MIC-II, Metrohm Co., Switzerland) using  
158 an anionic exchange resin (Metrosep A Supp 5 - 100/4.0). Operating conditions were a mobile phase at  
159 33 °C, Na<sub>2</sub>CO<sub>3</sub> (3.2 mM) / NaHCO<sub>3</sub> (1.0 mM) and flow rate 0.7 mL/min. The anions were identified and  
160 quantified relative to certified commercial standards (Certipur ®Anion multi-element standard I, II, Merck,  
161 Germany). On the other hand, elemental components (Al, As, Cd, Ca, Cu, Cr, Fe, Mg, Mn, Ni, Pb, K, Na)  
162 were analyzed by inductively-coupled-plasma mass spectrometry (ICP-MS, Agilent 7500 instrument,  
163 Agilent Technologies, Tokyo, Japan) using a certified multi-element stock solution (Perkin-Elmer Pure

164 Plus standard, product number 9300233). All determinations were made with three independent samples  
165 and the reported results correspond to their average. The uncertainty ( $\pm$ ) was calculated using the  
166 standard deviation (s) of the three repetitions (n = 3) with a coverage factor K = 2.

167

#### 168 *X-Ray Diffraction (XRD) analysis*

169

170 Powder X-ray diffraction patterns were recorded with a D8 Advance diffractometer (Bruker AXS, Madison,  
171 USA, LynxEye detector) with Cu radiation ( $\lambda = 0.154$  nm) from  $2\theta$   $15^\circ$  to  $80^\circ$  and step size  $0.018^\circ$ .  
172 Identification was made with ICDD PDF-2 2007 Database [ICDD 2007] and comparison with published  
173 ferrihydrite diffraction patterns.

174

#### 175 *Scanning Electron Microscope and Electron Dispersive Spectrometer (SEM-EDS).*

176

177 The chemical elements that constitute the reddish mat and sediments were determined with SEM-EDS.  
178 For that purpose, samples were analyzed with a scanning electron microscope (Hitachi S-570) with  
179 energy-dispersive X-ray spectra (SEM-EDS).

180

#### 181 *Total DNA isolation, construction of 16S rRNA gene libraries and Illumina sequencing*

182

183 The three water samples from each sampling point (3 x 1L each) were pooled and filtered through a vac-  
184 uum system under sterile conditions using a membrane filter (pore size  $0.22 \mu\text{m}$ ; Millipore, GV CAT No  
185 GVWP04700). To prevent rupture, another filter membrane (pore size  $0.45 \mu\text{m}$ ; Phenex, Nylon Part No  
186 AF0-0504) was placed below. The upper filter was collected and stored at  $-80^\circ\text{C}$  until processing. The  
187 DNA was extracted from aseptically cut pieces of the filter with a DNA isolation kit (PowerSoil®, MoBio,  
188 Carlsbad, CA, USA) as described by the manufacturer. Cell lysis was accomplished by two steps of bead  
189 beating (FastPrep-24, MP Biomedicals, Santa Ana, CA, USA) for 30 s at  $5.5 \text{ m s}^{-1}$ . To process the sedi-  
190 ments, a homogeneous sample of 500 mg was collected and DNA was extracted using the same proto-  
191 col. For the construction of microbial 16S rRNA amplicon libraries, the V5-V6 hypervariable regions were  
192 PCR-amplified with universal primers 807F (5'-GGATTAGATACCCBRGTAGTC-3') and 1050R (5'-  
193 AGYTGDCGACRRCRTGCA-3') (Bohorquez et al. 2012). The barcoding of the DNA amplicons and the  
194 addition of Illumina adaptors were conducted by PCR from approximately 1-10 ng of total DNA as de-  
195 scribed previously (Burbach et al. 2016). The PCR-amplified amplicons were verified by 1% agarose gel

196 electrophoresis, purified using the QIAquick Gel Extraction Kit (Qiagen, Hilden, Germany) and quantified  
197 with the Quant-iT PicoGreen dsDNA reagent kit (Invitrogen, Darmstadt, Germany). The individual  
198 amplicons were pooled to equimolar ratios (200 ng of each sample) and purified with the QIAquick PCR  
199 Purification Kit (Qiagen, Hilden, Germany). Finally, PCR-generated amplicon libraries were subjected to  
200 250 nt paired-end sequencing on a MiSeq platform (Illumina, San Diego, CA, USA).

201

### 202 *Bioinformatic and phylogenetic analysis of 16S rDNA amplicon data*

203

204 The sequence data were deposited in the sequence-read archive (SRA) of GenBank under the BioProject  
205 PRJNA678689. Bioinformatic processing was performed as previously described (Schulz et al. 2018).  
206 Raw reads were merged with the Ribosomal Database Project (RDP) assembler (Cole et al. 2014).  
207 Sequences were aligned within MOTHUR (gtoth algorithm using the SILVA reference database; (Schloss  
208 et al. 2009) and subjected to preclustering (diffs=2) yielding the so-called operational taxonomic units  
209 (OTUs) that were filtered for an average abundance of  $\geq 0.001\%$  and a sequence length  $\geq 250$  bp before  
210 analysis. OTUs were taxonomically classified based on the SILVA v138 database with the RDP taxonomy  
211 (Yilmaz et al. 2014) as reported by the SINA classification tool (Pruesse et al. 2012). OTUs were  
212 assigned to a taxonomic rank only if their best hit in the SILVA database (Pruesse et al. 2007) had an  
213 identity higher than the threshold established by Yarza et al. (2014) for that rank (94.5% for genus, 86.5%  
214 for family, 82.0% for order, 78.5% for class and 75.0% for phylum). Moreover, the sequences of some  
215 highly abundant OTUs were also manually examined by means of BLASTN (Altschul et al. 1997) against  
216 the non-redundant and against the bacterial and archaeal 16S rRNA databases. Sequences that weren't  
217 assigned to at least a phylum were excluded for further analysis. The statistical analyses and their  
218 visualization were performed with the R statistical program (R-Core-Team 2019) and Rstudio interface.  
219 Package Vegan v2.5-6 (Oksanen et al. 2019) was used to calculate alpha diversity estimators (Observed  
220 richness, Shannon index, Simpson index). Data tables with the OTUs abundances were normalized into  
221 relative abundances and then converted into a Bray–Curtis similarity matrix. The effects of factors on the  
222 bacterial community composition were evaluated with non-parametric permutational analysis of variance  
223 (PERMANOVA) (adonis2 function with 999 permutations) and using multivariate analysis such as  
224 Principal Coordinate Analysis (PcoA). Significant differences between a priori predefined groups of  
225 samples were evaluated using Permutational Multivariate Analysis of Variance (PERMANOVA), allowing  
226 for type III (partial) sums of squares, fixed effects sum to zero for mixed terms. Monte Carlo p-values

227 (p<sub>MC</sub>) were generated using unrestricted permutation of raw data (Anderson 2001) IN PRIMER (v.7.0.11,  
228 PRIMER-E, Plymouth Marine Laboratory, UK). We used the Kruskal Wallis test to determine significant  
229 differences in alpha diversity estimators (Simpson and Shannon-index).

230

## 231 **Results and Discussion**

232

### 233 *Chemical Analysis*

234

235 The major physicochemical properties of Bajo las Peñas are shown in Table 1. This habitat corresponds  
236 to a hot spring with moderately high temperature, ranging from 62.0-68.0°C, and pH nearly neutral,  
237 ranging from 6.5 to 7.4. Despite the small differences in pH and temperature throughout the site, we  
238 observed small but consistent differences between BP1 and BP3 (the two sources) and the stream sites  
239 BP2 and BP4 being the sources slightly more acidic and hotter. The oxygen levels (2.01-2.53 mg O<sub>2</sub>/L  
240 measured at 22°C) in all the samples indicate that Bajo las Peñas corresponds to a dysoxic environment.  
241 (Spietz et al. 2015). Physicochemical analysis of water samples showed the presence of moderate levels  
242 of sulfate (243-284 mg/L), magnesium, manganese, sodium, aluminum, calcium, iron and potassium. Our  
243 data (samples taken in 2016) agree with the chemical composition reported by Uribe-Lorio et al (2019) of  
244 samples taken in 2012, suggesting that the inorganic chemistry of Bajo Las Peñas has remained  
245 relatively stable. Examination of the dried sediment samples revealed considerable heterogeneity and  
246 organic matter content. Even a few ca. 0.2 mm sized magnetite and sand grains were present and were  
247 removed prior to the X-Ray measurements. This heterogeneity is expected as the sediments do not  
248 consist of a pure mineral. EDS data showed sediments to be iron-rich though extremely variable (50-70%  
249 Fe) and we assign them to ferric hydroxide precipitates. The samples were very amorphous, required  
250 several hours of Powder X-Ray data collection and only one did diffract strongly enough to identify a  
251 mineral phase. The XRD pattern (Fig. S1) of the reddish sediment sample from site BP1 was assigned to  
252 6-line ferrihydrite rather than the less crystalline 2-line ferrihydrite. The spacings for the diffraction peaks  
253 at 2.53, 2.22, 1.96, 1.72 1.50 and 1.46 Å compare very well to the literature (Fig. S1; Drits et al. 1993).  
254 Only an unassigned peak at 2.44 Å remained. Interestingly, the conversion from 2-line ferrihydrite into 6-  
255 line ferrihydrite has been reported to occur under biotic conditions (Kukkadapu et al. 2003). Ferrihydrite is  
256 normally very amorphous and the presence of organic matter made collecting Powder X-Ray data  
257 difficult. We could not obtain a diffraction pattern from other sediment samples, and we include as an  
258 example the essentially absent diffraction of the sample from the BP2 site as reference (Fig. S1).  
259 Additionally, previous reports of similar environments indicate that the mixture of the sediments and the



260 microbial mat with the water flow produces the characteristic red color of these habitats (Casanova et al.  
261 1999; Bishop and Murad 2002; Naren et al. 2012), as is the case in Bajo las Peñas (see Fig. 2).

262

263 In spite that Bajo las Peñas is located close to other acidic rivers (e.g. Rio Sucio and Toro Amarillo rivers)  
264 likely associated with the area's sulfate and chloride-rich hydrothermal water, Bajo las Peñas  
265 physicochemical properties are particularly different: it is endowed with moderate sulfate concentrations in  
266 the water column, low chloride, neutral pH and precipitation of ferrihydrite rather than ferric-oxide sulfate  
267 minerals. All these observations suggest that the sulfate source in Bajo las Peñas is different from Acid  
268 Rock Drainage (ARD) and Volcanic influenced Acid Rock Drainage (VARD) environments. We have  
269 described VARD ecosystems associated with nearby Irazu Volcano's hydrothermal system (Arce-  
270 Rodríguez et al. 2020) in which the sulfate and the acidity come from the volcano hydrothermal system  
271 rather than from biological oxidation of pyrite as in ARD systems (Amils et al. 2002; Arce-Rodríguez et al.  
272 2017, 2019). However, they do not seem to influence Bajo las Peñas and instead, the likely sources of  
273 sulfate are anhydrite/gypsum dissolution perhaps aided by some acid rain as the area is downwind from  
274 active Turrialba Volcano (Porowski et al. 2019). The hydrothermal system at Bajo las Peñas contributes  
275 therefore mostly heat and ferrous iron to the aquifer rather than acid-forming magmatic gases.

276

### 277 *Analysis of microbial communities*

278

279 We determined the presence of 916 OTUs in all samples, which were taxonomically assigned to 35 phyla  
280 of Bacteria and Archaea (Supplementary Table 1). The most-abundant phyla were Actinobacteria (0.8-  
281 40.0%), Proteobacteria (4.9-39.7%), Chloroflexi (0.8-29.1%), Nitrospirae (2.6-28.6%), Firmicutes (0.4-  
282 25.9%), Thaumarchaeota (0.1-25.8%), Crenarchaeota (0.1-24.0%), Deinococcus-Thermus (0.4-19.6%),  
283 Aquificae (0-16.0%), Ignavibacteriae (0-14.0%), Acetothermia (0-5.6%), Dictyoglomi (0-5.1%),  
284 Acidobacteria (0.8-4.7%), Cyanobacteria (0-4.4%), and Planctomycetes (0-4.0%). Nonetheless, we  
285 observed large variations across samples (Fig. 3).

286

287 Comparisons between the global bacterial assemblages in the different sites showed the presence of  
288 three well-defined clusters: (i) water springs (BP1, BP3), (ii) sediments in the springs (BP1S1, BPS2,  
289 BP1S3), (iii) waters and sediments downstream from the springs (BP2, BP2S1, BP2S2, BP4, BP4S1,  
290 BP4S2) (Fig. 4) (PERMANOVA  $p = 0.001$ ). All groups of samples were significantly different from one  
291 another as indicated by pairwise tests (source water vs downstream samples  $p_{MC} = 0.05$ , source  
292 sediments vs downstream samples,  $p_{MC} = 0.01$  and source water vs source sediment,  $p_{MC} = 0.04$ ). This  
293 clustering pattern indicates a fast community turnover right after communities emerge from underground

294 to the surface, possibly in response to new environmental conditions such as increase in oxygen and  
295 carbon input.

296

297 Alpha diversity estimations (Fig. S2) showed that all samples present a Shannon index > 2.8. Kruskal-  
298 Wallis test showed no significant differences in the diversity indices across the clusters described.  
299 Previous analyses of the microbial mats in Bajo las Peñas (Uribe-Lorío et al. 2019) showed a high  
300 abundance of photoautotrophic prokaryotes, where almost the entire the community was dominated by  
301 few OTUs. However, our analyses, including samples in the water column and sediments showed that  
302 this environment holds a very variable microbial community (See Figs 3 and 4). The high dynamism that  
303 exists in this extreme environment, where the community fluctuates according to minor changes in the  
304 water conditions and external factors, such as the climate (Ferris and Ward 1997; Fraterrigo et al. 2006;  
305 Li et al. 2014; Narayan et al. 2018) is also a factor to consider when comparing results. Interestingly, the  
306 absence of iron-oxidizing bacteria (see below) can be related to the neutral pH which favors its abiotic  
307 oxidation, thus reducing the availability of electron donors for this type of bacteria (Earhart 2009; Johnson  
308 et al. 2012). This is one of the clearest differences in Bajo las Peñas compared to other hot springs that  
309 we have characterized in Costa Rica (Arce-Rodríguez et al. 2019, 2020), whose physicochemical  
310 characteristics are typical of ARD or VARD environments (with acidic pH), where the presence of iron  
311 oxidizers is very common.

312

313 The most abundant microorganisms in the origins of Bajo las Peñas (water samples BP1 and BP3) were  
314 phylotypes from the genus *Sulfurihydrogenibium*, particularly OTU\_439 (10.3-14.8%) (Fig. 5) and some  
315 putative members of the family Nitrospiraceae (OTU\_67, 10.0-26.8%), which closest relatives were  
316 bacteria of the genus *Thermodesulfobivrio* (92.2% identity). The genus *Sulfurihydrogenibium* has to our  
317 knowledge five reported species: *S. azorensis* (Aguilar et al. 2004), *S. kristjanssonii* (Flores et al. 2008), *S.*  
318 *rodmanii* (O'Neill et al. 2008), *S. subterraneum* (Takai et al. 2003), and *S. yellowstonense* (Nakagawa et  
319 al. 2005). These microorganisms are reported as aerobic, thermophilic, and sulfur-oxidizing bacteria.  
320 These bacteria isolated from hot springs grow at temperatures between 40-80°C and can use S<sup>0</sup>,  
321 thiosulfate and sulfite as electron donors. Some species from this genus have reported the ability to use  
322 CO<sub>2</sub> as carbon source. Moreover, these microorganisms are known to be important organic carbon  
323 producers in the environments where they were isolated (Nunoura 2014). Thus, we suggest that  
324 members of the genus *Sulfurihydrogenibium* could act as CO<sub>2</sub>-fixers in Bajo Las Peñas.

325

326 On the other hand, the relatively low sequence identity of OTU\_67 with members of the genus  
327 *Thermodesulfobivrio*, does not allow us to reliably affirm that they belong to this genus. However,

328 considering that (i) they belong to the family Nitrospiraceae, (ii) the moderate levels of available sulfate,  
329 and (iii) the temperature in Bajo las Peñas, presumably members of OTU\_67 could participate in sulfate  
330 reduction in a similar way as species of *Thermodesulfovibrio* genus do. Bacteria from the  
331 *Thermodesulfovibrio* genus have typically been isolated from hot springs and are generally known as  
332 anaerobic sulfate reducers that grow at temperatures between 40-70°C. *T. yellowstonii* has the ability to  
333 use sulfate, thiosulfate, sulfite and nitrate as terminal electron acceptors and oxidizes organic compounds  
334 like pyruvate and lactate to acetate. Autotrophic growth using hydrogen have also been reported in  
335 members of this genus (Daims 2014).

336

337 In the sediments taken from Bajo las Peñas sources it was possible to find the same OTU from the genus  
338 *Sulfurihydrogenibium* (OTU\_439, 1.6-5.2%) and the OTU\_67 (2.2-2.8%) found mostly in the waters.  
339 However, many other OTUs closely affiliated to *Thermoflexus* (OTU\_320, 1.9-2.9%) and *Thermus*  
340 (OTU\_5, 8.9-18.9%), to the families Dictyoglomaceae, Thermomicrobiaceae and Aquificaceae  
341 respectively, (OTU\_480, 1.9-5.1%; OTU\_492, 2.7-3.8% and OTU\_64, 2.5-4.6 %) or to the Clostridia  
342 (OTU\_128, 7.5-18.9%) were also detected with a high abundance (Figure 5).

343

344 Bacteria from the genera *Thermoflexus* and *Thermus* have been reported to be chemoheterotrophic  
345 microorganisms, that may consume byproducts of other microorganism or substrates deposited in the  
346 sediments of Bajo las Peñas. Bacteria of the phylum Chloroflexi were also detected in some of these  
347 sediment samples (6.1-10.6%) (Fig. 5), which is in accordance with a report by Uribe-Lorio et al. (2019)  
348 on the microbial mat. However, our data suggest that in Bajo las Peñas the first level of the trophic chain  
349 is not only associated with photosynthesis by Cyanobacteria and Chloroflexi but also by CO<sub>2</sub> fixation  
350 linked to sulfur metabolism by autotrophic bacteria which derives probably from the subdeposits of water  
351 in Bajo las Peñas.

352

353 A few meters downstream of the origin (BP2 and BP4) the organics produced in the first trophic level  
354 (described above) as well as the organic matter input may generate the conditions necessary for  
355 thermotolerant chemoheterotrophic bacteria to colonize the place. This would a drastic shift in the  
356 microbial community structure with respect to the samples obtained at the origin, although spatially the  
357 samples are separated by only 5-20 meters. In the water samples BP2 and BP4, bacteria of the family  
358 Nitrososphaera (8.2-25.8%), Ignavibacteriaceae (7.3-7.9%) and of the genus *Nitrospira* (6.2-9.7%) and  
359 *Meiothermus* (5.6-10.6%) were dominating. Most of these phlotypes were also dominating in the  
360 sediment samples from sampling points 2 and 4 (see Fig. 5). We propose that many of these  
361 microorganisms found downstream the origins (sampling points 2 and 4), use fermentation products

362 generated by other microorganisms found at the origin (sampling points 1 and 3). For instance,  
363 fermentation products like acetate, have been reported to be produced by members of the genera  
364 *Thermodesulfobrio* and *Ignavibacterium* (Lino et al. 2009) which were closely related to the  
365 microorganism identified in our samples. These byproducts could be used as a carbon source by some  
366 acetate-oxidizing bacteria. These observations suggest a continuous cross-feeding between the different  
367 niches of this hot spring.

368

369 Besides the possible carbon cross-feeding, a well-characterized relationship between ammonium-  
370 oxidizing and nitrite-oxidizing microorganisms may be occurring. Members of the family Nitrososphaera  
371 (2.1-25.7%), may oxidize ammonium to nitrite, which is subsequently utilized by nitrite-oxidizing Nitrospira  
372 (6.2-14.0%), for growth (Hovanec et al. 1998; Stieglmeier et al. 2014; Daims et al. 2015; Lehtovirta-  
373 Morley 2018). Such a nitrification process has been reported in other hot springs and is known to be an  
374 essential process for the maintenance of the community (Reigstad et al. 2008). Thus, our data are  
375 aligned with the existence of multiple cross-feeding processes, which are probably the product of  
376 microbial adaptation to the low abundance of organic carbon producers in the environment.

377

378 As mentioned above, Bajo las Peñas hot spring is characterized for having a red-gelatinous microbial  
379 tappet (Fig. 1C) which, we propose, is made of a complex bacterial community in association with  
380 ferrihydrite minerals. Uribe-Lorio et al. (2019) reported that the microbial mat in Bajo las Peñas is  
381 composed mainly by Cyanobacteria (*Limnothrix*, *Leptolyngbya*, *Phormidium* and *Fisherella*) and to a  
382 lower extent by *Meiothermus*. We were also able to detect *Meiothermus* which is considered a key  
383 microorganism in hot springs due to its ability to form a strong biofilm by specialized structures  
384 (Spanevello and Patel 2004; Masurat et al. 2005; Wu et al. 2009; Mori et al. 2012).

385

## 386 **Conclusions**

387

388 This study sheds light on the ecology of microbial communities of a rare type of hot springs with neutral  
389 conditions located in Bajo las Peñas, Costa Rica. In this hot spring, abiotic precipitation of ferric  
390 oxyhydroxides in an extensive bacterial mat contribute to the red-colored appearance of the habitat. The  
391 analyses of 16S rRNA sequences revealed that Bajo las Peñas harbors a very complex environment  
392 where the microbial community changes along the water flow from a chemoautotrophic to a  
393 chemolithotrophic metabolism. CO<sub>2</sub> fixation occurs through photosynthetic microorganisms (Cyanobacteria  
394 and *Chlorella*) but also bacteria that fix CO<sub>2</sub> linked to sulfur metabolism (i.e. *Sulfurihydrogenibium*). In

395 this first autotrophic level, organic molecules are produced that then can be used by heterotrophic  
396 microorganisms such as those closely related to *Thermodesulfovibrio* and *Ignavibacterium*. We propose  
397 that the microbial community in Bajo Las Peñas may have a complex metabolism due to niche  
398 specialization where cross-feeding processes between phylotypes plays an essential role. Our data  
399 demonstrate the quick change in the microbial community of an ecosystem in response to environmental  
400 variables such as organic matter and oxygen availability. In Bajo las Peñas it was possible to observe  
401 how in just ~5 meters (e.g. BP1 vs BP2) the global microbial community and the dominant species  
402 change from autotrophic to heterotrophic microorganisms. This work also helps to understand  
403 extremophile communities present in neutral hot springs, which are considerably less explored compared  
404 to acidic hot springs, and shows that they are an interesting environment for the bioprospecting of  
405 microorganisms and extremozymes with promising biotechnological applications.

406

#### 407 **Acknowledgements**

408

409 We thank Carlos Rodriguez of Centro de Investigación en Contaminación Ambiental (CICA) for the help  
410 in the chemical analysis. We also are grateful to Solange Voysest and Jose Jimenez for help in the  
411 design of some figures.

412

#### 413 **Author contributions**

414

415 AAR, FPS, EL, MC conceived and designed the experiments; AAR, RA, RMA, DRG, PF, GCB performed  
416 the experiments; AAR, FPS, DRG, KR, MC analyzed the data; PF, MC, DHP contributed reagents or  
417 materials or analytical tools; DRG, EL, KR, DHP, MC wrote the paper. All authors reviewed and approved  
418 the final version of the manuscript.

419

#### 420 **Funding**

421

422 This work was supported by the Vice-rectory of Research of Universidad de Costa Rica (VI 809-B6-524),  
423 the Costa Rican Ministry of Science, Technology and Telecommunication (MICITT) and Federal Ministry  
424 of Education and Research (BMBF) (project VolcanZyme contract No FI-255B-17) and by the ERC grant  
425 IPBSL (ERC250350-IPBSL). F.P-S. is supported by the Spanish Economy and Competitiveness Ministry  
426 (MINECO) grant CTM2016-80095-C2-1-R.

427

428 **Compliance with ethical standards**

429

430 **Conflict of interest** The authors declare that there are no conflicts of interest.

431 **Ethical approval** This study does not describe any experimental work related to human.

432

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731

732 **Tables**

733

734 **Table 1. Physical properties and chemical composition of Bajo las Peñas hot spring.**

Property/element/ion	BP-1	BP-2	BP-3	BP-4
Coordinates	9.993492 N 83.801597 W	9.993558 N 83.801639 W	9.993603 N 83.801677 W	9.993658 N 83.801716 W
Distance from the first origin (m)	0	10	15	20
Temperature (°C) / ± 0.1	68.0	65.6	67.0	62.0
pH / ± 0.1	6.5	7.4	6.6	7.3
Dissolved oxygen (measured to 22°C)	2.01	2.05	2.23	2.53
Aluminium (µg L <sup>-1</sup> ) ± 3	15	25	13	14
Arsenic (µg L <sup>-1</sup> )	<2	<2	<2	<3
Cadmium (µg L <sup>-1</sup> )	< 0.11	< 0.17	< 0.12	< 0.11
Calcium (mg L <sup>-1</sup> ) ± 2	15	11	20	11
Chloride (mg L <sup>-1</sup> ) ± 1	< 0.20	1	1	1
Cooper (mg L <sup>-1</sup> )	< 0.10	< 0.10	< 0.10	< 0.10
Chromium (µg L <sup>-1</sup> )	< 1.2	< 1.2	< 1.2	< 1.2
Fluoride (mg L <sup>-1</sup> )	<0.09	<0.09	<0.09	<0.09
Iron (mg L <sup>-1</sup> ) / ± 0.06	0.60	0.16	0.16	<0.08
Magnesium (mg L <sup>-1</sup> ) ± 0.7	11.6	12.2	12.6	13.0
Manganese (mg L <sup>-1</sup> ) ± 0.06	0.36	0.36	0.41	0.40
Nickel (µg L <sup>-1</sup> )	< 1	< 1	< 1	< 1
Nitrate (mg L <sup>-1</sup> )	< 0.10	< 0.10	< 0.10	< 0.10
Lead (µg L <sup>-1</sup> )	< 1.2	< 1.2	< 1.2	< 1.2
Potassium (mg L <sup>-1</sup> ) ± 0.3	9.8	10.0	10.6	10.6
Sodium (mg L <sup>-1</sup> ) ± 0.8	38.5	39.3	38.1	38.5
Sulfate (mg L <sup>-1</sup> ) ± 20	264	281	284	243
Zinc (µg L <sup>-1</sup> )	< 25	< 25	< 25	< 25

735

736 **Legends of figures**

737

738 **Fig. 1 Bajo las Peñas hot spring** A) Bajo las Peñas hot spring is located in Cartago Costa Rica near  
739 Turrialba Volcano and Toro Amarillo River. B) It is a small lotic hot spring, with two springs (BP1 and BP3)  
740 where water emerges from underground deposits. C) The water flow around Bajo las Peñas is very slow  
741 and it flows for about 20 meters. D) The entire hot spring is covered by a dense reddish gelatinous  
742 microbial mat where ferrihydrite precipitate constantly.

743

744 **Fig. 2 SEM-EDS analysis of Bajo las Peñas sediments.** A) SEM micrographs to 2  $\mu\text{m}$  (15.0 kV 7.6 mm  
745  $\times$  60) of the reddish gelatinous sediments at the bottom of Bajo las Peñas. Micrographs indicate that the  
746 sediment consists of lumps of mainly amorphous material of nanometer size. B) EDS analysis of the  
747 sediments showed that it is composed of mainly iron and oxygen. Other species such as silicon,  
748 phosphate and chlorine were found in the sediment in minor proportions.

749

750 **Fig. 3 Taxonomic composition of Bajo las Peñas microbial communities.** Relative abundance of  
751 bacterial and archaeal organisms to phylum. The OTUs were taxonomically classified using RDP  
752 database, as described in Materials and Methods. Waters of Bajo las Peñas are termed BP1 to BP4 and  
753 sediments are termed BP1S1 to BP4S2.

754

755 **Fig. 4 Principal Coordinate Analysis (PCoA) of the prokaryotic communities in Bajo las Peñas.** A  
756 clustering of communities according to the sampling point is shown. The PCoA and the PERMANOVA  
757 analyses were performed with the package Vegan.

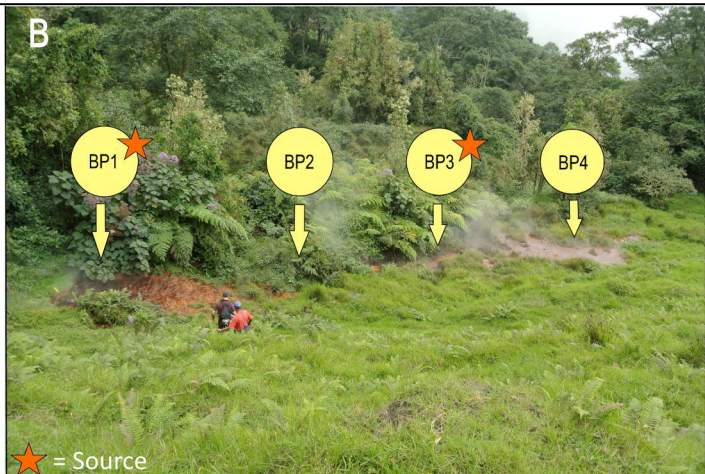
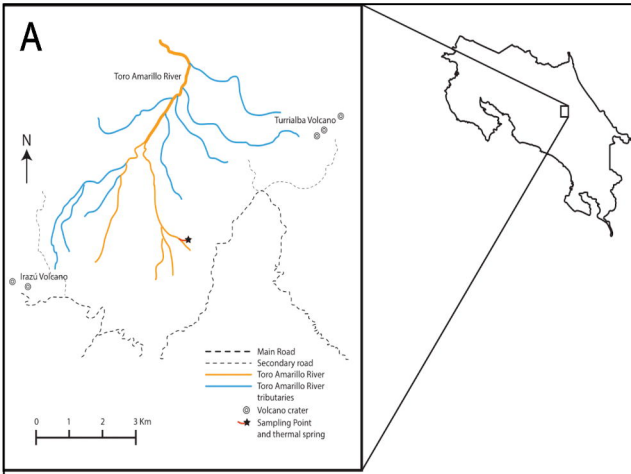
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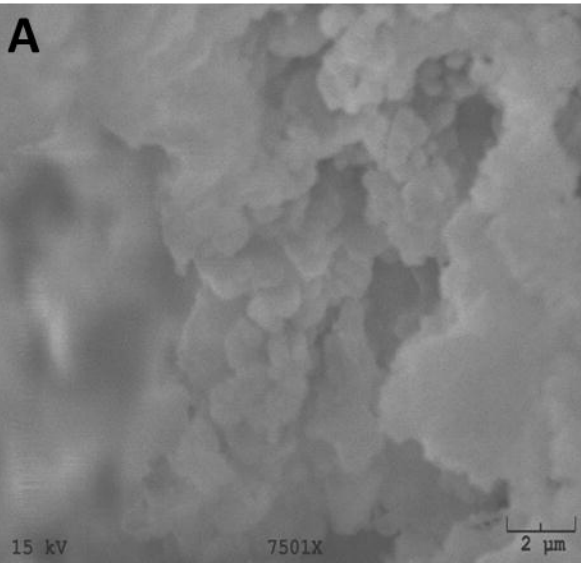
759 **Fig. 5 Heat map of the most abundant OTU in each sample.** The heat map depicts the relative  
760 percentage of 16S rRNA gene sequences assigned to each OTU (y axis) across the 11 samples analyzed  
761 (x axis) with an abundance higher than 2%.

762

763





**A****B**