

1 Reconstructing genomes of carbon monoxide oxidisers in volcanic 2 deposits including members of the class Ktedonobacteria

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4 **Marcela Hernández**^{1,2,6,*}, **Blanca Vera-Gargallo**³, **Marcela Calabi-Floody**⁴, **Gary M King**⁵, **Ralf**
5 **Conrad**⁶ and **Christoph C. Tebbe**¹

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7 ¹ Thünen Institut für Biodiversität, Braunschweig, Germany, christoph.tebbe@thuenen.de;

8 ² School of Environmental Sciences, Norwich Research Park, University of East Anglia, Norwich, UK,
9 marcela.hernandez@uea.ac.uk;

10 ³ Department of Microbiology and Parasitology, Faculty of Pharmacy, University of Seville, Seville, Spain,
11 vera@us.es;

12 ⁴ BIOREN-UFRO, Universidad de La Frontera, Temuco, Chile; marcela.calabi@ufrontera.cl;

13 ⁵ Department of Biological Sciences, Louisiana State University, Baton Rouge, LA, USA;
14 gkingme@gmail.com;

15 ⁶ Max Planck Institute for Terrestrial Microbiology, Marburg, Germany, conrad@mpi-marburg.mpg.de;

16
17 * Correspondence: marcela.hernandez@uea.ac.uk (M.H.)

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19 **Abstract:** Microorganisms can potentially colonize volcanic rocks using the chemical energy in
20 reduced gases such as methane, hydrogen (H₂) and carbon monoxide (CO). In this study, we
21 analysed soil metagenomes from Chilean volcanic soils, representing three different successional
22 stages with ages of 380, 269 and 63 years, respectively. A total of 19 metagenome-assembled
23 genomes (MAGs) were retrieved from all stages with a higher number observed in the youngest soil
24 (1640: 2 MAGs, 1751: 1 MAG, 1957: 16 MAGs). Genomic similarity indices showed that several
25 MAGs had amino-acid identity (AAI) values >50% to the phyla Actinobacteria, Acidobacteria,
26 Gemmatimonadetes, Proteobacteria and Chloroflexi. Three MAGs from the youngest site (1957)
27 belonged to the class Ktedonobacteria (Chloroflexi). Complete cellular functions of all the MAGs
28 were characterised, including carbon fixation, terpenoid backbone biosynthesis, formate oxidation
29 and CO oxidation. All 19 environmental genomes contained at least one gene encoding a putative
30 carbon monoxide dehydrogenase (CODH). Three MAGs had form I *coxL* operon (encoding the large
31 subunit CO-dehydrogenase). One of these MAGs (MAG-1957-2.1, Ktedonobacterales) was highly
32 abundant in the youngest soil. MAG-1957-2.1 also contained genes encoding a [NiFe]-hydrogenase
33 and *hyp* genes encoding accessory enzymes and proteins. Little is known about the
34 Ktedonobacterales through cultivated isolates, but some species can utilize H₂ and CO for growth.
35 Our results strongly suggest that the remote volcanic sites in Chile represent a natural habitat for
36 Ktedonobacteria and they may use reduced gases for growth.

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39 **Keywords:** Ktedonobacteria; carbon monoxide dehydrogenase; metagenome-assembled-genome;
40 volcanic soils; shotgun sequencing

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45 1. Introduction

46 Volcanic eruptions provide a model for understanding soil-forming processes and the roles of
47 pioneer bacteria during early biotic colonization. Recently, it has been demonstrated that the
48 structure of microbial communities can play a key role in the direction of plant community succession
49 pathways [1]. This is due in part to bacterial contributions to weathering of volcanic rocks, which
50 releases nutrients resulting in some of the most fertile soils in the world.

51 After lava and other volcanic deposits (i.e. ash and tephra) cool sufficiently, mineral surface
52 areas become accessible for microbial colonization [2-5]. In fact, microbes, and especially bacteria, are
53 among the first colonizers of volcanic deposits and thereby initiate soil formation during the early
54 stages of terrestrial ecosystem development [6-12]. While methane (CH₄), hydrogen sulfide (H₂S),
55 hydrogen (H₂), and carbon monoxide (CO) have been proposed to promote bacterial colonization
56 and support microbial life in these organic-carbon deficient environments [13,14], the actual carbon
57 and energy sources of the first colonizers remain elusive, but likely include a range of endogenous
58 and exogenous sources, including reduced minerals and gases. The microbial ability to utilize these
59 substrates for growth, and thereby initiate the formation of soil organic material, depends on
60 specialized enzymes which may not be prevalent in many different microbial groups, thus
61 representing a limited phylogenetic distribution. As a consequence, this could constrain the
62 composition of pioneering microbial communities.

63 CO is a potential source of carbon and energy for microbes pioneering the colonization of
64 volcanic substrates. CO utilization under oxic conditions requires a molybdenum-dependent carbon
65 monoxide dehydrogenase (Mo-CODH), which catalyses the oxidation of CO to CO₂ [15]. Surveys of
66 genome databases (e.g., Integrated Microbial Genomes) reveal that Mo-CODHs occur in
67 Proteobacteria, Actinobacteria, Firmicutes, Chloroflexi, Bacteroidetes, Deinococcus-Thermus,
68 Halobacteria, and Sulfolobales among others. They were originally described as inducible enzymes
69 and have subsequently been shown to be up-regulated by carbon limitation during growth of
70 Actinobacteria (e.g., *Rhodococcus* and *Mycobacterium smegmatis* [15-17], and Chloroflexi (e.g.,
71 *Thermogemmatispora* and *Thermomicrobium* [18]). Mo-CODH has been previously targeted in
72 molecular ecological studies of CO oxidizers in volcanic systems and other extreme environments
73 [15]. These studies have revealed changes in community composition with the age and
74 developmental status of individual sites [7,15]. The results suggest that CO-oxidizing communities
75 are not static, but that they change in response to changing environmental conditions, and possibly
76 affect the direction of changes.

77 In a preceding study, we identified the microbial communities involved in volcanic soil
78 formation in different sites on Llaima Volcano (Chile). The bacterial communities of soils from 3 sites
79 affected by lava deposition in 1640, 1751 and 1957 were analysed using 16S rRNA gene amplicon
80 sequencing [19] and it was demonstrated that microbial diversity increased with the age of the soil
81 deposits. Interestingly, bacterial phylotypes of the poorly studied Ktedonobacterales were among the
82 predominant community members in the 1957 soil, representing 37% of all OTUs, as compared with
83 18% in the 1751 and 7% in the 1640 soils. Thus, we suspected that bacteria of this order could be
84 instrumental for the initiation of soil formation, paving the way for soil organic carbon formation and
85 preparing a substrate for microbial colonisation and plant growth. Some already cultivated
86 Ktedonobacterales were found to be carboxydrotrophs and hydrogenotrophs (i.e. carbon monoxide
87 (CO) and hydrogen (H₂) oxidisers/consumers) [14]. Thus, this leads to the hypothesis that CO and H₂
88 are important carbon and energy sources for early stages of microbes colonizing the Llaima Volcano
89 soil. However, the ecophysiology of the few bacterial isolates assigned to Ktedonobacterales limits
90 predictions about metabolic functions based on 16S rRNA gene sequences alone.

91 Therefore, in this study a metagenomic approach was chosen to identify microbial traits
92 associated with early stages of colonization and soil formation in a volcanic ecosystem. In particular,
93 the presence of functional genes implicated in CO-oxidation (*coxL* genes) and H₂-oxidation (*hyd* and
94 *hup* genes) was assessed from metagenome-assembled genomes (MAGs). These MAGs were
95 retrieved from volcanic soils of different ages, representing sites 1640, 1751 and 1957. During this
96 period, the soil formation evolved as indicated by their different levels of soil organic matter ranging

97 from 65.33 ± 2.31 in the most recent soil (1957) to 9.33 % in both medium (1751) and oldest soils (1640)
98 [20].

99 Especially the most recent soil (youngest soil) was suspected to reveal microbial adaptations to
100 the challenging environmental conditions and thus to unveil the metabolic processes which initiate
101 microbial colonisations. Therefore, functional metabolic modules annotated in the environmental
102 genomes were analysed, with a main focus on the poorly characterised class of the Ktedonobacteria
103 (Chloroflexi). Three Ktedonobacteria MAGs were obtained and all contained genes encoding CO and
104 H₂ oxidation. Additional MAGs from other phyla were also found to contain these genes. Our study
105 advances the understanding of the ecology of Ktedonobacteria and their potential to act as early
106 colonizers in volcanic soils.

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108 2. Materials and Methods

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110 2.1 Sequencing

111 The DNA from the volcanic soils used in our study had been previously extracted [20]. The soil
112 physico-chemical characteristics have been published [20] showing a pH of 5.6 in both medium and
113 oldest soil and 4.7 in the youngest soil, and nitrogen (mg/kg) of 25 (1640), 26 (1751) and 36.33 (1957).
114 Briefly, the soil samples originating from three different sites of different ages according to the latest
115 lava eruption (1640, 1751, 1957, map in [20]). A total of nine samples (triplicate per site) were
116 sequenced on an Illumina MiSeq at the Max-Planck-Genome Centre, Köln, Germany. The
117 metagenome was analysed on a high-performance computer using 650 GB RAM and 64 cores at the
118 Thünen Institute of Biodiversity, Braunschweig, Germany.

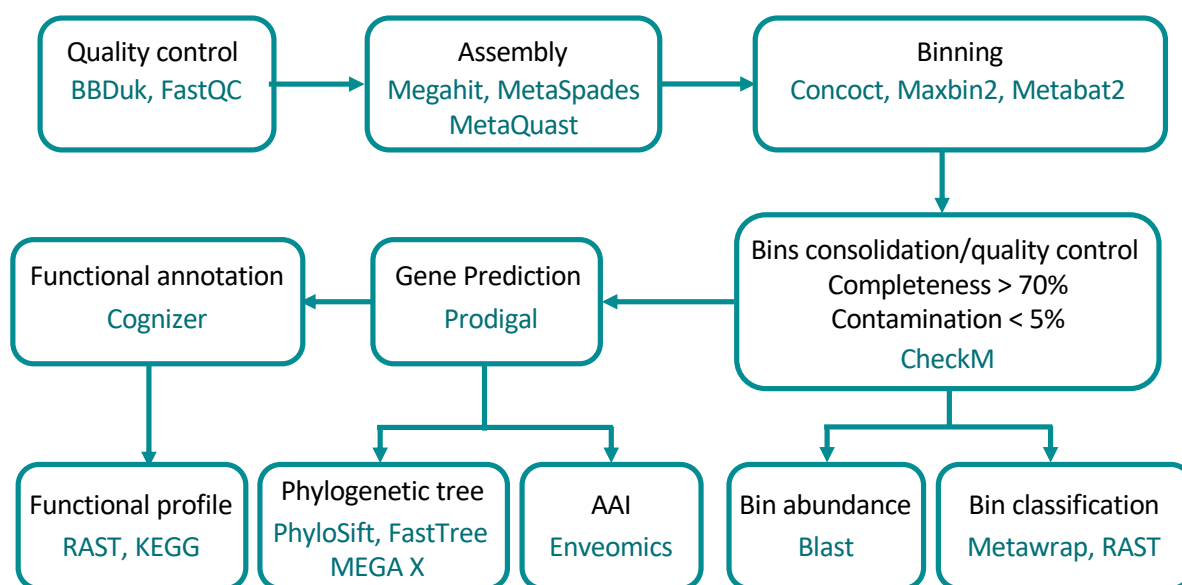
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120 2.2. Quality control

121 The sequence reads were checked using FastQC version 0.11.8 [21]. Low quality reads were
122 discarded using BBDuk version 38.68, quality-trimming to Q15 using the Phred algorithm [22]. A
123 schematic overview of the steps and programs used are shown in Figure 1.

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Figure 1. Workflow of the metagenome analysis and programs used in the present study. Binning was run by using MetaWrap package. AAI: Amino Acid Identity.

2.3. Metagenome assembly and binning

131 All trimmed Illumina reads were merged into longer contiguous sequences (scaffolds) using *de*
132 *novu* assemblers Megahit version 1.2.8 [23] with k-mers 21,29,39,59,79,99,119,141 and MetaSPAdes
133 (SPAdes for co-assembly) version 3.13.1 [24,25] with k-mers 21,31,41,51,61,71,81. Triplicate samples
134 were co-assembled in order to improve the assembly of low-abundance organisms. Assembly quality
135 was checked with MetaQuast version 5.0.2 [26] showing that the best quality was obtained with
136 MetaSPAdes for our samples (data not shown). Downstream analysis was carried out using the
137 scaffolds retrieved from MetaSPAdes. Krona charts [27] were recovered from MetaQuast runs to
138 identify taxonomic profiles. Downstream binning analysis was performed with two sets of scaffolds:
139 full size scaffolds and scaffolds larger than 1000 bp.

140 Metagenomic binning of the assembled scaffolds was carried out with the metaWRAP version
141 1.2.1 pipeline [28], which binning module employs three binning software programs: MaxBin2 [29],
142 metaBAT2 [30], and CONCOCT [31]. Completion and contamination metrics of the extracted bins
143 were estimated using CheckM [32]. The resulting bins were collectively processed to produce
144 consolidated metagenome-assembled genomes (MAGs) using the bin_refinement module (criterion:
145 completeness > 70%; contamination < 5%). Both sets of MAGs (18 from scaffolds larger than 1000 bp
146 and 17 from full size scaffolds) were aggregated, visualized with VizBin [33] and then dereplicated
147 using dRep [34]. Only the highest scoring MAG from each secondary cluster was retained in the
148 dereplicated set. The abundance of each MAG in the different sites was calculated using BLASTN
149 version 2.5.0+ [35] keeping only hits with >95% identity and e-value 1e-5 for the analysis [36]. A final
150 heatmap was constructed using the function heatmap.2 from the gplots package version 3.0.4 [37] in
151 R version 4.0.2 (<https://www.r-project.org>).

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153 2.4. Functional annotation

154 The open reading frames (ORFs) in all scaffolds of each MAG were predicted using Prodigal
155 (v2.6.3) [38]. Functions were annotated using Cognizer [39] and KEGG annotation framework [40].
156 The annotations of the predicted proteins from the Kyoto Encyclopedia of Genes and Genomes
157 (KEGG) were used to confirm protein functional assignment and identify pathways. Complete
158 pathways were identified using KEGG BRITE pathway mapping [40]. Aerobic carbon-monoxide
159 dehydrogenases and hydrogen dehydrogenase were also identified using KEGG ortholog
160 annotations; CODH was further distinguished as form I and form II (putative CODH) based on active
161 site motifs present in *coxL* genes (e.g., [41]).

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163 2.5. Phylogenomic analysis

164 Taxonomic classification of MAGs was performed using the classify_bins module from
165 metaWRAP which relies on NCBI_nt database. MAGs were also screened using the RAST Server
166 (Rapid Annotations using Subsystems Technology; [42,43]), which also allowed to retrieve
167 information regarding close relative genomes in order to construct the phylogenetic tree.

168 To estimate intergenomic similarity, amino-acid comparisons between MAGs and their closest
169 relative genomes present in the databases were calculated based on reciprocal best hits (two-way
170 AAI) using the enveomics collection (<http://enve-omics.gatech.edu/> [44]).

171 The phylogenetic affiliation of MAGs was determined by constructing a genomic tree using
172 FastTree version 2.1.11 [45]. Reference genomes were manually downloaded from the National
173 Center for Biotechnology Information (NCBI) Refseq database (Table S1). Conserved genes from the
174 extracted bins and the reference genomes were concatenated using Phylosift version 1.0.1 [46].

175 Phylogenetic analysis of the large sub-unit CO dehydrogenase gene (*coxL*) using the Maximum
176 Likelihood method with a JTT matrix-based model [47] was performed. Bootstrap values (100
177 replicates) are shown where support ≥ 70 percent. The scale bar indicates substitutions per site. All
178 gapped positions were deleted resulting in 420 positions in the final dataset. Evolutionary analyses
179 were conducted in MEGA X [48,49].

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181 2.6. Accession number

182 Raw metagenomic data and environmental genomes derived from binning processes were
183 deposited in the Sequence Read Archive (SRA) under the bioproject accession number PRJNA602600
184 for raw data and PRJNA602601 for metagenome-assembled genomes.
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186 3. Results

187 3.1. MAGs recovery

188 A total of ~3-4 million scaffolds were recovered from the soil metagenomes in each site. Even
189 though all the sites underwent similar sequencing efforts (between 2.3 GB in 1640 and 1957 to 2.4
190 Gb in 1751), the youngest soil had the largest number of scaffolds with 499 sequences >50 kb
191 compared to the oldest soil with only 12 scaffolds with a size >50 kb (Table 1). The youngest soil
192 also had a larger N50 length of scaffolds and L50 compared to the other soils (Table 1). A total of 19
193 MAGs with a completeness of >70% and a contamination <5% (2 from 1640, 1 from 1751 and 16
194 from 1957) were retrieved and characterised.

195 **Table 1. Summary report for the assembly quality assessment using MetaQuast**

Statistics	1640	1751	1957
Number of scaffolds (>= 0 bp)	3631380	4047900	3138527
Number of scaffolds (>= 500 bp)	1488437	1609020	1578887
Number of scaffolds (>= 1000 bp)	326980	333885	447418
Number of scaffolds (>= 25000 bp)	172	195	1551
Number of scaffolds (>= 50000 bp)	12	40	499
Total length (>= 0 bp)	2320437372	2460323814	2396069091
Total length (>= 500 bp)	1385386471	1404094641	1717252425
Total length (>= 1000 bp)	631179105	570404360	964085094
Total length (>= 25000 bp)	5697241	8152100	82365310
Total length (>= 50000 bp)	686375	2983958	46419897
N50	914	850	700
L50	391387	477193	348604

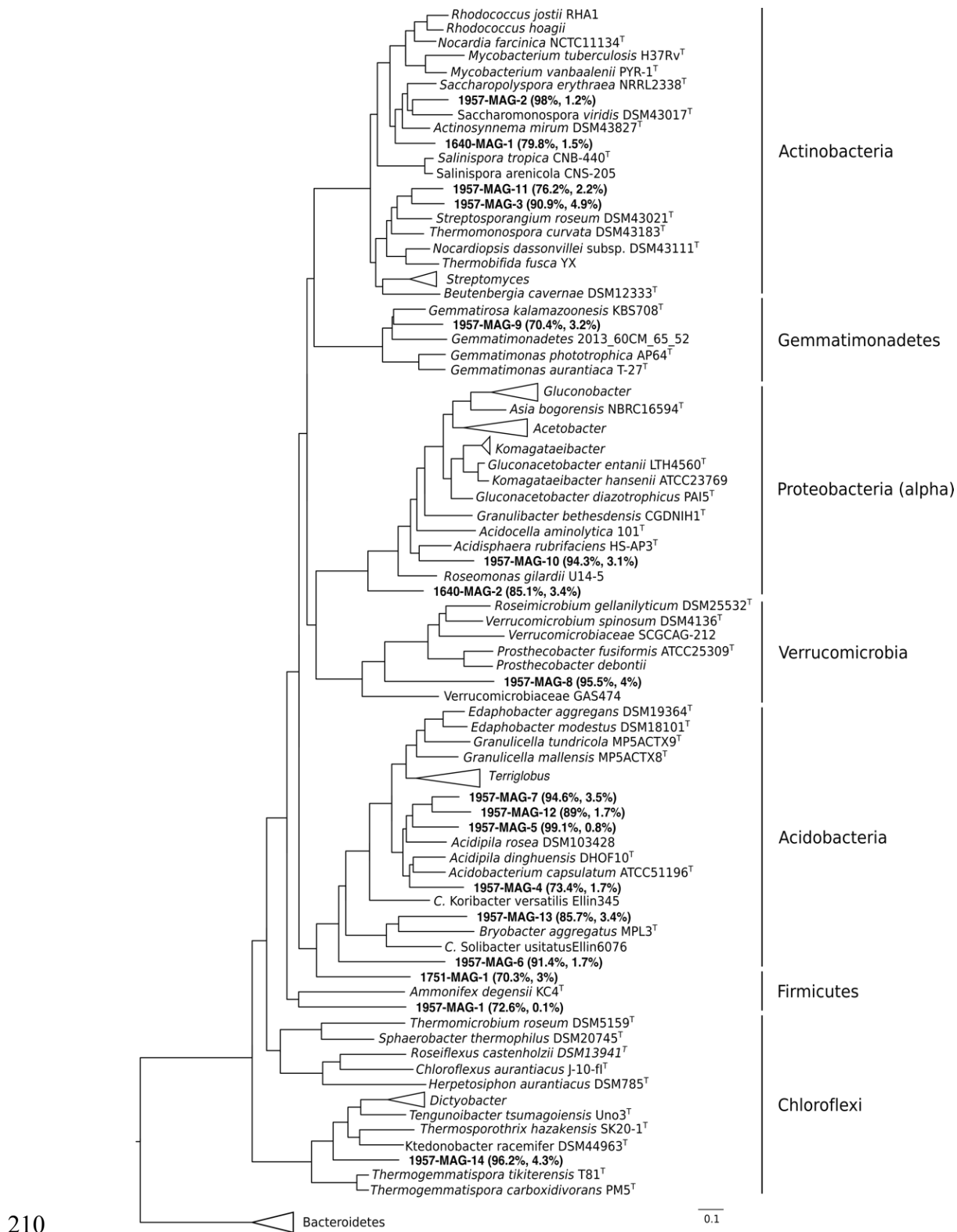
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197 N50 - length such that scaffolds of this length or longer include half the bases of the assembly; L50 -
198 number of scaffolds that are longer than, or equal to, the N50 length and therefore include half the
199 bases of the assembly (<https://www.ncbi.nlm.nih.gov/assembly/help/#globalstats>)

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201 3.2. MAG identification

202 MAGs were affiliated to the phyla Actinobacteria, Proteobacteria, Acidobacteria,
203 Gemmatimonadetes, Chloroflexi, Firmicutes and Verrucomicrobia (Figure 2). In the oldest soil, two
204 environmental genomes were retrieved related to Actinomycetales (Actinobacteria) and
205 Rhodospirillales (Proteobacteria). The only MAG retrieved from the middle soil was related to
206 Acidobacteria. MAGs binned from the youngest soil included seven assigned to Acidobacteria, one
207 to Proteobacteria, three to Actinobacteria, one to the phylum Gemmatimonadetes, one to
208 Verrucomicrobia and three to the phylum Chloroflexi (Figure 2).

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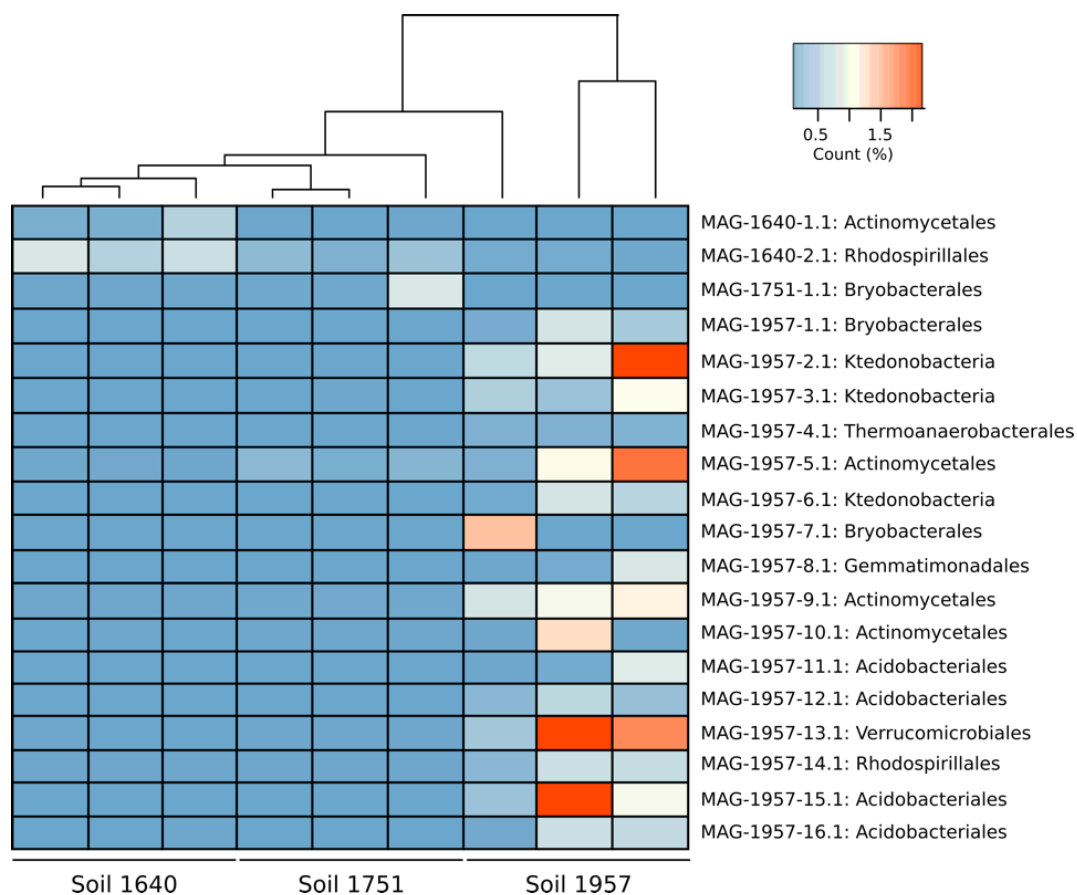
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211 Figure 2. Phylogenomic tree of the bacterial genomic bins. The tree was built with PhyloSift against
 212 reference genomes downloaded from NCBI. MAGs are indicated in bold together with their
 213 respective completeness and contamination. FastTree confidence values of MAG branches are
 214 shown. The horizontal bar represents 10% sequence divergence.

215 The abundance of the MAGs in each site was calculated by using BLASTN (Figure 3). MAGs were
 216 more abundant from the soil they were recovered. MAGs with a total abundance >1% were found
 217 only in the young soil (1957). MAG 1957-2.1 (Ktedonobacteria, 1.21% ± 0.82), MAG 1957-5.1

218 (Actinomycetales, $1.02\% \pm 0.84$), MAG 1957-13.1 (Verrucomicrobiales, $1.58\% \pm 1.1$) and MAG 1957-
 219 16.1 (Acidobacteria, $1.14\% \pm 0.87$) were the most abundant MAGs (Figure 3).

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224 Figure 3. Heatmap representing the abundance of MAGs in each metagenome. The analysis was done
 225 by blast and only hits greater than 95% identity and e-value $1e-5$ were used.
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227 3.3. Metabolic characterisation of MAGs

228 Genes encoding enzymes involved in carbohydrate and energy metabolism, such as carbon fixation,
 229 sulfur metabolism, ATP synthesis and nitrogen metabolism, as well as terpenoid backbone
 230 biosynthesis were found in all the MAGs (Table 2). Other functions, including xenobiotic
 231 biodegradation, fatty acid metabolism, nucleotide metabolism, and vitamin metabolism, among
 232 others, were also found (Table S2).

233 Table 2. Summary table of complete cellular functions and other high-level features in the MAGs
 234 recovered from sites 1640, 1751 and 1957 retrieved from KEGG analysis. Reference genomes for
 235 Ktedonobacteria: DSM45816T [50], DSM44963T [51] and NBRC 113551T [52], (K: KEGG orthology;
 236 M: KEGG Mode). Asterisks indicate the MAGs isolated from the class Ktedonobacteria. "X" indicates
 237 MAGs containing genes encoding form I of the CoxL.

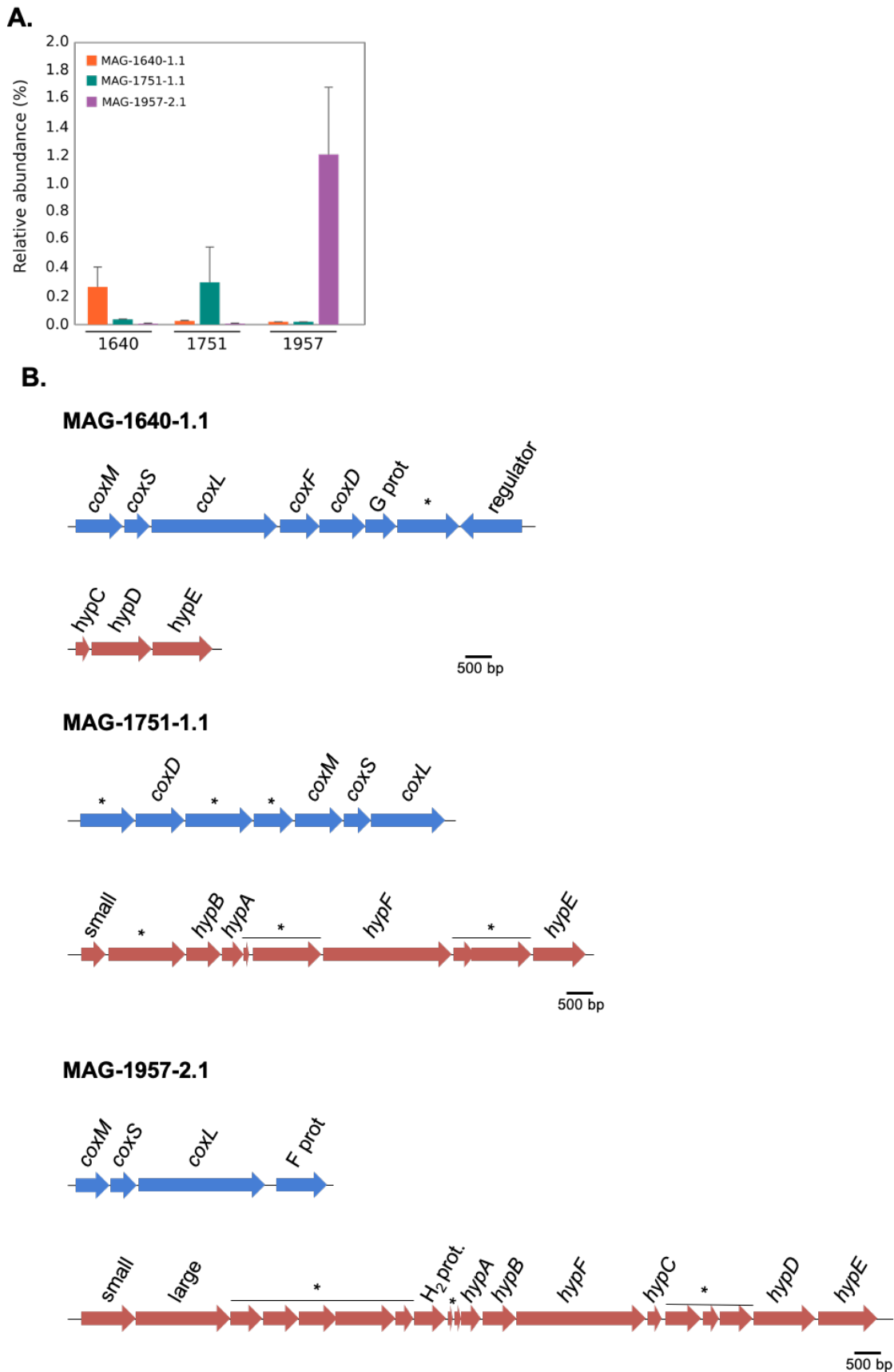
	1640	1751	1957													Reference							
Cellular functions	MAG-1.1	MAG-2.1	MAG-1.1	MAG-1.1	MAG-2.1 *	MAG-3.1 *	MAG-4.1	MAG-5.1	MAG-6.1 *	MAG-7.1	MAG-8.1	MAG-9.1	MAG-10.1	MAG-1.1	MAG-1.2.1	MAG-1.3.1	MAG-1.4.1	MAG-1.5.1	MAG-1.6.1	DSM45816 [†]	DSM44963 [†]	NBRCT13551 [†]	
Carbon monoxide oxidation																							
K03518 Carbon monoxide dehydrogenase small subunit <i>coxS</i>																							
K03519 Carbon monoxide dehydrogenase medium subunit <i>coxM</i>																							
K03520 Carbon monoxide dehydrogenase large subunit <i>coxL</i> (Form-II)																							
K03520 Carbon monoxide dehydrogenase large subunit <i>coxL</i> (Form-I)	X	X	X																				
Formate oxidation																							
K00122 Formate dehydrogenase																							
K00123 Formate dehydrogenase major subunit																							
Hydrogen oxidation: H₂ dehydrogenase																							
K00436: NAD-reducing hydrogenase large subunit																							
Central carbohydrate metabolism																							
M00001 Glycolysis (Embden-Meyerhof pathway), glucose => pyruvate																							
M00002 Glycolysis, core module involving three-carbon compounds																							
M00003 Gluconeogenesis, oxaloacetate => fructose-6P																							
M00004 Pentose phosphate pathway (Pentose phosphate cycle)																							
M00005 PRPP biosynthesis, ribose 5P => PRPP																							
M00006 Pentose phosphate pathway, oxidative phase, glucose 6P => ribulose 5P																							
M00007 Pentose phosphate pathway, non-oxidative phase, fructose 6P => ribose 5P																							
M00009 Citrate cycle (TCA cycle, Krebs cycle)																							
M00010 Citrate cycle, first carbon oxidation, oxaloacetate => 2-oxoglutarate																							
M00011 Citrate cycle, second carbon oxidation, 2-oxoglutarate => oxaloacetate																							
M00307 Pyruvate oxidation, pyruvate => acetyl-CoA																							
M00580 Pentose phosphate pathway, archaea, fructose 6P => ribose 5P																							
Other carbohydrate metabolism																							
M00012 Glyoxylate cycle																							
M00061 D-Gluconate deg., D-gluconate => pyruvate + D-glyceraldehyde 3P																							
M00081 Pectin degradation																							
M00549 Nucleotide sugar biosynthesis, glucose => UDP-glucose																							
M00552 D-galactonate deg., D, L-Doudoroff pathway, D-galactonate => glycerate-3P																							
M00554 Nucleotide sugar biosynthesis, galactose => UDP-galactose																							
M00631 D-Galacturonate deg., D-galacturonate => pyruvate + D-glyceraldehyde 3P																							
M00632 Galactose degr., Leloir pathway, galactose => alpha-D-glucose-1P																							
M00741 Propanoyl-CoA metabolism, propanoyl-CoA => succinyl-CoA																							
M00761 Undecaprenylphosphate & L-Ara4N biosynthesis																							
M00854 Glycogen biosynthesis, glucose-1P => glycogen/starch																							
M00855 Glycogen degradation, glycogen => glucose-6P																							
M00909 UDP-N-acetyl-D-glucosamine, prokaryotes, glucose => UDP-GlcNAc																							
Carbon fixation																							
M00166 Reductive pentose phosphate cycle, ribulose-5P => glyceraldehyde-3P																							
M00168 CAM (Crassulacean acid metabolism), dark																							
M00169 CAM (Crassulacean acid metabolism), light																							
M00172 C4-dicarboxylic acid cycle, NADP - malic enzyme type																							
M00579 Phosphate acetyltransferase-acetate kinase pathway, acetyl-CoA => acetate																							
ATP synthesis																							
M00144 NADH:quinone oxidoreductase, prokaryotes																							
M00149 Succinate dehydrogenase, prokaryotes																							
M00151 Cytochrome bc1 complex respiratory unit																							
M00155 Cytochrome c oxidase, prokaryotes																							
M00156 Cytochrome c oxidase, ebb3-type																							
M00157 F-type ATPase, prokaryotes and chloroplasts																							
Terpenoid backbone biosynthesis																							
M00096 C5 isoprenoid biosynthesis, non-mevalonate pathway																							
M00364 C10-C20 isoprenoid biosynthesis, bacteria																							
M00365 C10-C20 isoprenoid biosynthesis, archaea																							

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3.3.1 Characterization of CODH and hydrogenase genes in MAGs

Three MAGs (MAG-1640-1.1, MAG-1751-1.1 and MAG-1957-2.1) encoded form I of the CO-dehydrogenase large subunit (*coxL*). These MAGs were each associated with a particular soil, with low abundance in the metagenomes of the other sites (Figure 4A). In addition to these three form-I *coxL* encoding MAGs, 15 other scaffolds from MAGs containing form II *coxL*-like genes were recovered (data not shown), but the function of form II *CoxL* is not yet known. The arrangement of genes encoding form I CODH in each of the MAGs is shown in Figure 4B. It should be noted that all of these three MAGs show the canonical arrangement for the three structural genes of CODH, that is the MSL genes. The genes encoding the [NiFe]-hydrogenase and its accessory proteins were only identified in MAG-1957-2.1, and instead only some of the accessory *hyp* genes were found in the other two MAGs (Figure 4B). A phylogenetic analysis of the form I *coxL* genes was performed, showing they are affiliated with Actinobacteria (MAG-1640-1.1), Nitrospirae *Candidatus* Manganitrophus noduliformans (MAG-1751-1.1) and Chloroflexi (MAG-1957-2.1) (Figure 5). This grouping is consistent with the results of PhyloSift (Figure 2), except for MAG-1751-1.1 where it was loosely associated with Acidobacteria (although with an amino-acid identity of only 40%) rather than Nitrospirae.

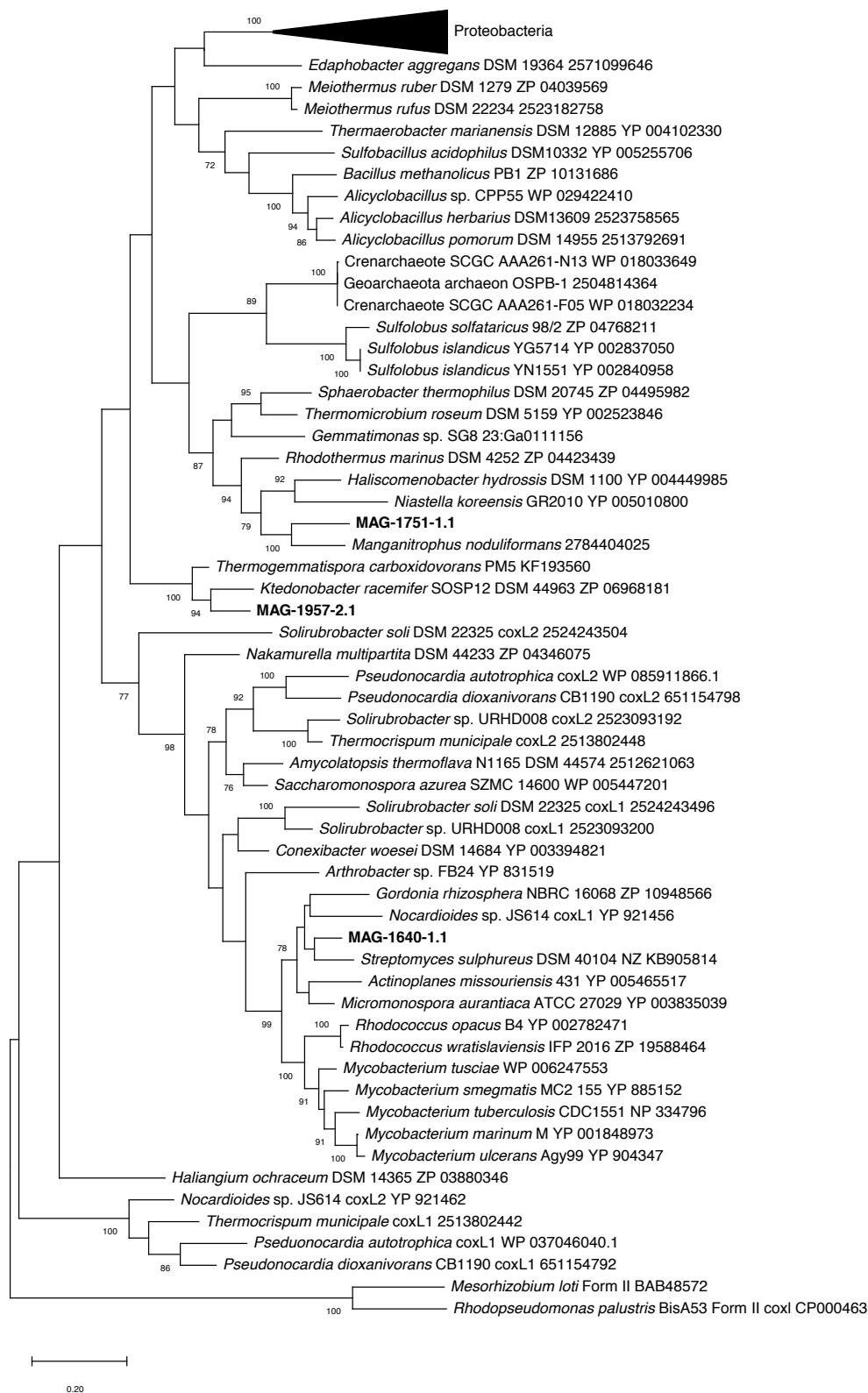
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Figure 4. A. Relative abundance of MAG-1640-1.1, MAG-1751-1.1 and MAG-1957-2.1 in sites 1640, 1751, 1957. Bars indicate standard error of triplicates. B. Gene arrangement of the carbon monoxide dehydrogenase (CODH) and membrane-bound [NiFe]-hydrogenase in MAG-1640-1.1, MAG-1751-

264 1.1 and MAG-1957-2.1 (* indicates hypothetical proteins; small indicates NiFe-hydrogenase small
265 subunit and large indicates NiFe-hydrogenase large subunit).
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267
268 Figure 5. Phylogenetic tree of form I carbon monoxide dehydrogenase large subunit (CoxL) of
269 metagenome-assembled genomes retrieved from Llaima volcano (MAG-1640-2.1, MAG-1751-1.1 and
270 MAG-1957-2.1) against a reference sequences (with accession numbers included in the tree). The tree

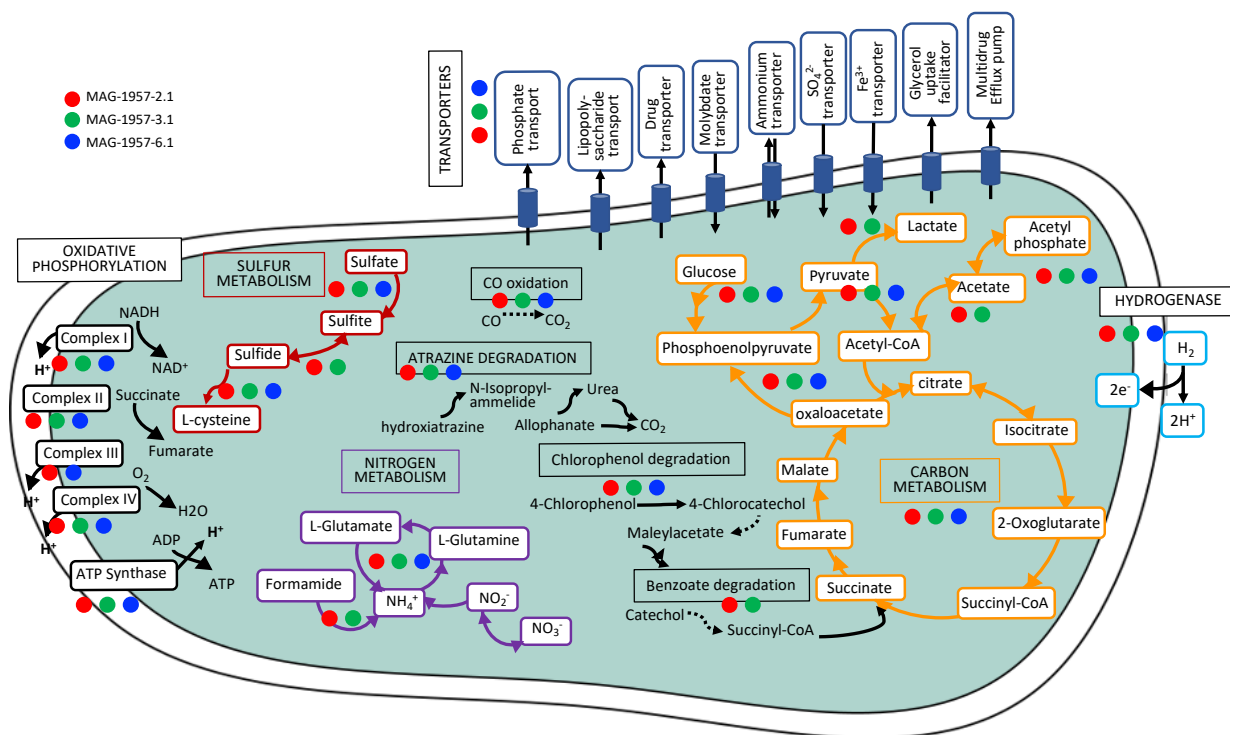
271 was drawn using the Maximum Likelihood method using MEGA X [48]. Bootstrap values (500
272 replications) are shown at the nodes. MAG-1751_1.1 is a partial sequence of the *coxL* gene.

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274 3.3.2 Complete metabolic characterization of Ktedonobacterales MAGs

275 Here we focused on the Ktedonobacterales MAGs because of their apparent importance in early soil
276 formation. Three Ktedonobacterales MAGs were identified in the 1957 soil metagenomes, but were
277 not found in the older soils (Figure 2 and 3). Two of the MAGs (MAGs 1957-2.1 and 1957-3.1),
278 related/affiliated to the class Ktedonobacteria (phylum Chloroflexi), contained genes for the complete
279 electron transport chain, citric acid metabolism, nitrogen metabolism, sulfur metabolism, several
280 transporters, the complete gene set for carbon monoxide oxidation (CO dehydrogenase), herbicide
281 degradation and degradation aromatics as well as the major subunit of the formate dehydrogenase,
282 and also a hydrogenase. MAG 1957-6.1 (Ktedonobacteria) had very similar pathways as the other
283 Chloroflexi MAGs, except a step for CO-oxidation and the electron transport chain were absent (Table
284 2, Table S2, Figure 6).

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287 Figure 6. Metabolic reconstruction for some of the most important functions in the
288 Ktedonobacteria MAGs isolated from the youngest soil using KEGG (More information in Table 2
289 and Table S2), dashed lines were used for incomplete pathways. Multi-arrows lines indicate several
290 steps of a pathway.

291

292 4. Discussion

293 4.1. Characterisation of MAGs

294 In this study, a characterisation of metagenome-assembled genomes retrieved from Llaima volcano
295 was performed. This study builds from a previous study [19] in which 16S rRNA gene amplicon-
296 based sequences from those soils were analysed. The main objective of this study was to characterise
297 genomes from those sites and to analyse the functions of the abundant but the poorly characterised
298 Ktedonobacteria (phylum Chloroflexi) present at Llaima volcano. The relative abundance of the main
299 phyla based on classification of scaffolds larger than 500 bp showed that microbial communities

300 change as the soils age (Figure S1). This corroborates findings from a previous study [19]. For
301 example, the relative abundance of Chloroflexi is higher in the younger soils (28% in the youngest
302 soil to 7% in the oldest soil) and the opposite trend is observed for members of the phylum
303 Proteobacteria, as their abundance increases as the soil ages (from 42% in the youngest soil to 59% in
304 the oldest soil) (Figure S1). Except for those related to Firmicutes and Verrucomicrobia, and to a lesser
305 extent Acidobacteria and Proteobacteria, the extracted environmental genomes had an amino acid
306 identity > 50% with their closest reference genome (Table S3), which suggest that they belonged to
307 those genera [53].

308 A total of 16 MAGs were recovered from the youngest soil (1957) (Figure 3). This soil is only
309 partially vegetated (about 5%) by mosses and lichens. The microbial community in this area likely
310 harbours populations able to grow as facultative chemolithoautotrophs or mixotrophs on carbon
311 monoxide, hydrogen or methane. This high relative abundance of MAGs with genes for CO and
312 hydrogen utilization in the youngest soils is consistent with reports by King and colleagues for
313 Hawaiian and Japanese volcanic deposits (21- to 800-year old sites). For some of those sites, microbial
314 community structure changed as the soil matured with members of the phylum Proteobacteria
315 dominating vegetated sites while younger sites were enriched with Ktedonobacteria within the
316 Chloroflexi and characterized by relatively high rates of atmospheric CO uptake [7,14,54].

317 MAGs were most abundant in the soil site from where they were retrieved (Figure 3). Relatively
318 few MAGs were retrieved from the two older soils, which can be explained by the higher diversity
319 in these soils and the decreased likelihood of recovering MAGs from groups such as Actinobacteria,
320 Acidobacteria and Chloroflexi that were less common in them. In fact, several of the MAGs retrieved
321 had a low relative abundance within the soils (Figure 3), which is consistent with their relative
322 abundance of 16S rRNA genes in these soils [19]. Binning at the strain level remains a technical
323 challenge [55], with the chances of retrieving MAGs at a given sequencing effort being reduced with
324 increasing microdiversity (intra-population genetic diversity) and overall community diversity [56].
325 We previously reported that as the soil recovered and vegetation established, the microbial
326 population appeared to enlarge and become more diverse [19], which explains the lower number of
327 MAGS retrieved from more mature soil (1640 sample), compared to the younger sites (1957).

328 4.2. Metabolic characterisation of MAGs

329 The three MAGs containing form I *coxL* genes were found in an operon structure (Fig. 4B) typical
330 of known CO oxidizers [41]. Form I *coxL* has been definitively associated with CO oxidation at high
331 concentrations and also at sub-atmospheric levels [41]. Thus, even at low abundance, the presence of
332 these *cox*-containing MAGs strongly suggests a capacity for atmospheric CO uptake at all the sites.

333 Most of the complete functions found from the Ktedonobacteria MAGs were also found in three
334 reference genomes: *Ktedonobacter racemifer* DSM 44963 [51], *Thermogemmatispora carboxidivorans* PM5,
335 isolated from a geothermal biofilm on Kilauea Volcano, Hawaii (USA) [50] and *Dictyobacter vulcani*
336 W12 [52]. According to our genomic analyses, all of these reference strains possess formate-, H₂-, and
337 CO-dehydrogenases as do the MAGs recovered in the present study. *Burkholderia* strains (phylum
338 Proteobacteria) [57], members of the phylum Chloroflexi [14] and other members of the phyla
339 Proteobacteria and Actinobacteria [58] have also been reported as CO-oxidisers in Hawaiian volcanic
340 deposits. *coxL* genes encoding the large subunit of the CO dehydrogenase have been found in
341 Proteobacteria species from Kilauea and Miyake-jima volcanoes [10,14,54].

342 The taxonomies of MAGs 1640-1.1 and 1957-2.1 were consistent for *coxL* (Figure 5) and
343 phylogenomic analyses (Figure 2). In contrast, MAG-1751-1.1 clustered weakly with Acidobacteria
344 based on genomic analysis (40% amino acid identity with a reference genome, see Table S3) but did
345 not cluster with *Candidatus Manganitrophus noduliformans* as did the *coxL* sequence from this MAG.

346 Several strains from the class Ktedonobacteria have been isolated from different environments
347 (Table S4), but only *Dictyobacter vulcani* W12 [52] and *Thermogemmatispora carboxidivorans* PM5 [50]
348 have been isolated from volcanic environments. So far, the class Ktedonobacteria contains only six

349 genera and fifteen formally proposed species. Out of the 15 type strains, 11 genomes are available on
350 RefSeq (Table S2). The order Ktedonobacterales contains the type strains *Ktedonobacter racemifer*
351 SOSP1-21^T [51,59], *Dictyobacter aurantiacus* S-27^T [60], *Dictyobacter vulcani* W12^T [52], *Thermosporothrix*
352 *hazakensis* SK20-1^T [61], *Thermosporothrix narukonensis* F4^T [62], *Ktedonosporobacter rubrisoli* SCAWS-
353 G2^T [63], *Tengunoibacter tsumagoiensis* Uno3^T, *Dictyobacter kobayashii* Uno11^T and *Dictyobacter alpinus*
354 Uno16^T [64]. The order *Thermogemmatispora* contains the species *Thermogemmatispora aurantia* A1-2^T
355 [65], *Thermogemmatispora argillosa* A3-2^T [65], *Thermogemmatispora onikobensis* NBRC 111776
356 (unpublished, RefSeq Nr NZ_BDGT00000000.1), *Thermogemmatispora onikobensis* ONI-1^T [66],
357 *Thermogemmatispora foliorum* ONI-5^T [66] and *Thermogemmatispora carboxidivorans* PM5^T [50]. All of
358 those genomes contain the complete gene set for carbon monoxide oxidation (CO dehydrogenase),
359 as well as formate dehydrogenases, and a H₂ dehydrogenases (Table S4). Our study particularly
360 brings more insights into the role that early colonisers of this group from volcanic soils may have in
361 the development of soils.

362 The large subunit of the NAD-reducing hydrogenase was also found in several MAGs (Table 2).
363 Hydrogen metabolism has been shown to provide an additional energy source for some
364 microorganisms and has been observed in bacteria and archaea [67]. Hydrogen dehydrogenases have
365 also been found in members of the genus *Cupriavidus* (phylum Proteobacteria) from volcanic
366 mudflow deposits in the Philippines suggesting their potential contribution to hydrogen uptake [68].
367

368 4. Conclusions

369 This study is further evidence that poorly characterised groups such as Ktedonobacteria,
370 establish in remote volcanic sites and may use reduced gases for growth. Further studies are needed
371 to demonstrate the activity of these pathways and their significance in volcanic deposits.
372

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