

1 **Fire effect and its legacy modulate soil bacterial and fungal communities in Chinese**
2 **boreal forests along a chronosequence**

3 Wei-qin Su^{a,b}, Caixian Tang^c, Jiahui Lin^{a,b}, Mengjie Yu^{a,b}, Yu Luo^{a,b}, Yong Li^{a,b}, Zhongmin
4 Dai^{a,b}, Jianming Xu^{a,b}#

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6 ^a College of Environmental and Resource Sciences, Zhejiang University, Hangzhou 310058,
7 China

8 ^b Zhejiang Provincial Key Laboratory of Agricultural Resources and Environment, Zhejiang
9 University, Hangzhou 310058, China

10 ^c Department of Animal, Plant and Soil Sciences, La Trobe University, Melbourne Campus,
11 Bundoora, VIC 3086, Australia

12

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15 * Corresponding author, email: jmxu@zju.edu.cn

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17

18 **Abstract**

19 Wildfire has increasingly profound and pervasive consequences for forest ecosystems via
20 directly altering soil physicochemical properties and modulating microbial community. In
21 this study, we examined the changes in soil properties and microbial community at different
22 periods after highly severe wildfire events (44 plots, 113 samples) in the Chinese Great
23 Khingan Mountains. We also separated charcoals from burnt soils to establish the
24 relationship between soil microbes and the microbes colonized on the charcoal. Wildfire
25 significantly altered bacterial and fungal community structures across a 29-year
26 chronosequence. The network analysis revealed that from 17 years after fire, the complexity
27 and connectivity of bacterial and fungal communities were significantly increased.
28 Differential abundance analysis suggested that bacterial and fungal OTUs were enriched or
29 depleted only during 0-4 years after fire. In addition, soil factors, including soil pH, total C
30 and N, soil water content, and dissolved C and N, are key determinants of soil bacterial and
31 fungal communities from 17 years after fire. The fire-derived charcoals provided a new and
32 unusual niche for microbial colonization and charcoal microbes had a significantly different
33 community structure from the burnt soil microbes. Our data suggest that soil bacterial and
34 fungal communities changed dramatically during the recovery from fire events in terms of
35 the abundance and co-occurrence networks in the boreal forest ecosystems.

36 **Importance**

37 Pervious research has reported fire altered soil microbial community composition and
38 function during short-term succession in boreal forests. However, the long-term effect of fire

39 and fire-derived charcoals which are regarded as fire legacy effect on soil bacterial and
40 fungal communities composition and structure have not previously been shown.
41 Understanding how soil microbes particularly the keystone taxa and determinative soil
42 factors, respond to fire and its legacy matter charcoal, is critical for predicting how future
43 fire influences soil nutrient transformations and biological processes. We accessed time
44 chronosequence to examine the effect of fire history on soil microbial abundance and
45 co-occurrence network. These findings suggest that soil microbes can be reshaped by fire
46 and its legacy effect of fire-derived charcoal even in the long periods after fire and provide
47 further insights into fire and its legacy effect.

48 **Keywords:** Wildfire; Fire history; Legacy effect; Microbial community; Charcoal; Soil
49 factors; Forest ecosystems

50 **Introduction**

51 Fire is one of the most important disturbance agents in terrestrial ecosystems and a
52 worldwide phenomenon in our earth (1). Ecosystems such as boreal forests, shrublands,
53 grasslands and savannas, which have been regarded as “flammable systems”, are often
54 disturbed by wildfires over millions of years (2). Triggered by heat, fire affects soil
55 microbiota and nutrient cycling directly and indirectly (3, 4). Basically, severe wildfire
56 changes critical biotic and abiotic processes to cause complex consequences, including
57 significant removal of plant shoot and organic matter, destruction of the soil physical
58 structure and porosity, increase in nutrient losses through leaching and volatilization and
59 considerable shift of chemical properties (3, 5–7). Fire breaks biomolecules (8) by exposing
60 soil microbes directly to extremely high temperatures and lowers bacterial and fungal
61 biomass (by 33 - 49%) (4, 9). Meanwhile, fire-based disturbance also modifies microbial
62 community diversity and phylogenetic structure, which link to carbon dynamics and thawing
63 of permafrost (10, 11).

64 Boreal forest is a major fire-prone biome, covering approximately 30% area of the
65 global forest area. Distributed in high-latitude regions of Eurasia and North America, boreal
66 forests have low productivity and are easily forced by fire and climate change (12). During
67 long-term persistence and repeated succession, boreal forest ecosystems are tied to
68 corresponding fire patterns and post-fire effects. Future climate change (such as global
69 warming and drought) is likely to increase fire frequency and severity (13–15). Fire regimes
70 have brought international concerns in the global boreal forests.

71 In China, boreal forest landscapes encompass the most southern parts of the global
72 boreal forest (16). Most fires occur as a surface fire which is usually of moderate to high
73 severity, with an estimated historical fire interval period of 30-120 years (17). Chinese boreal
74 forests store approximately 24-31% of total carbon storage (1.0-1.5 Pg C) in China (14) and
75 the average of annual carbon emissions from boreal forest fire up to 0.54 Tg (18). Moreover,
76 strict fire suppression has been implemented in this region for over a half of century (19),
77 since a catastrophic wildfire (20) occurred on 6 May 1987. Such a human control has already
78 changed the fire regime—the patterns of fire spread, intensity, severity, frequency,
79 seasonality and ecological effects (2, 21).

80 The formation of charcoals is another important phenomenon during vegetation
81 burning, especially in the boreal forest region. Most carbon is converted to CO₂ releasing to
82 the atmosphere while a small portion of forest fuel (plant biomass and soil organic matter) is
83 pyrolyzed into other forms of pyrogenic carbon (PyC) (22). The PyC is a term synonymous
84 with black carbon (23) and is mainly produced as solid charred residue including
85 visually-defined charcoal and much lower proportions of volatile soot (24). Charcoal is part
86 of the PyC continuum. Because of its resistance to decomposition, it serves as (1) historic
87 records of fire and estimates for fire activities (1, 25), (2) an important long-term carbon sink
88 (22, 23, 26) and (3) different habitat patches for soil microorganisms (27–29). In addition,
89 incorporation of fire-derived charcoal into forest soils has an impact on soil C and nutrient
90 cycles, and vegetation regeneration (30–32). Owing to the presence of charcoal, burning *per*
91 *se* contributes to the shifts in ecological properties (including soil biological properties).

92 Therefore, charcoal is thought to function as fire's black legacy. However, some studies have
93 revealed that charcoal produced in fire would re-mineralize or degrade more rapidly than
94 expected (33, 34). Charcoal persistence directly influences soil properties in decades and
95 indirectly affect microbial community succession.

96 The effect of fire on soil biogeochemical and microbial processes can last for decades
97 and even centuries (35). The short-term effects of fire on soil pH, water content, organic
98 matter content, soil autotrophic respiration, and concentrations of total N, ammonium and
99 cations (K, Na, Ca, and Mg) are also observed in the boreal region (7, 37, 38). These effects
100 on soil properties in turn influence soil microbial community diversity and structure (7).
101 However, recovery patterns of soil microbial community and the functioning from fire
102 events are currently unclear in boreal ecosystems. The information is limited about the direct
103 and indirect impacts of fire, associated changes in edaphic factors and fire-derived charcoal
104 on soil microbes over long-term scales. The knowledge gap still exists in understanding
105 complex and diversified interactions among soil microorganisms with the soil environment
106 (such as in the soil-charcoal system). In addition, we use "space-for-time" substitution (39)
107 as a judicious strategy for chronosequence research. On the other hand, it is difficult to
108 separate charcoal particles from the soil and hence to identify specific taxa colonized in the
109 charcoal or soil samples. Consequently, we are not able to quantitatively assess the effect of
110 each individual factor even if charcoal could affect microbial community assembly through
111 several known pathways (40).

112 In this field study, we used a 29-year chronosequence to examine the effect of fire
113 history on soil microbial community recovery processes. Specifically, we quantified the
114 responses of soil microbes to fire and its legacy effect of fire-derived charcoal, and
115 determined the key determinants of bacterial and fungal community structure. We
116 hypothesized that (i) bacterial and fungal co-occurrence networks would become more
117 complex and connected with increasing time since fire (TSF), and (ii) the fire-derived
118 charcoal provided a unique niche for soil microbes and had microbial communities distinct
119 from the soil, and (iii) some soil properties would be the key determinants of soil microbial
120 community structure in soils with different fire histories.

121

122 **Results**

123 *Effect of fire on soil properties and microbial communities*

124 Fire enhanced the mean concentration of soil available P from 5.4 to 19.3 $\mu\text{g g}^{-1}$ ($p < 0.01$)
125 (Table S2) in A horizon. Compared with the unburnt controls, middle-term fires (MF)
126 increased available P by 224% and old fires (OF) increased available P by 257%. However,
127 no significant changes ($p > 0.05$) in soil properties were observed in O horizon. Severe
128 wildfires significantly increased soil C/N ratio and the concentrations of available P in both
129 horizons ($p < 0.05$) (Fig. S2). There was no statistical difference in other soil properties (Fig.
130 S2).

131 In total, 10480 bacterial and 56 archaeal OTUs were identified from 16S sequences,
132 while 2711 classified fungal OTUs assigned from 18S sequences. Proteobacteria

133 (30.6%-40.4%) were the most abundant phylum of bacteria overall in the soil samples (Fig.
134 1d), followed by Actinobacteria (22.2%-32.0%), Acidobacteria (12.6%-16.6%) and
135 Verrucomicrobia (4.2%-10.9%) (Table S3). Relative abundances of Proteobacteria,
136 Actinobacteria and Acidobacteria were generally constant across different fire histories
137 (Table S3), irrespective of soil horizon. Most of fungal OTUs were in Ascomycota
138 (41.3%-60.1%), with fewer in Basidiomycota (26.1%-49.4%) and Mucoromycota
139 (8.3%-13.2%) (Table S3). The proportions of dominant phyla were variable in the post-fire
140 soils. Overall, there is no significant temporal variation on taxonomic composition after fire
141 (Fig. 1d). Alpha diversity indices (includes Shannon index and OTU richness) of bacterial
142 and fungal communities did not vary significantly between the fire-affected (RF, MF, OF)
143 and unaffected groups (UF) (Fig. 1c). However, wildfire caused significant differences ($p <$
144 0.001, PERMANOVA) in β -diversity of soil microbial community (Fig. 1a & b) across
145 different fire histories (RF, MF, OF and UF). The NMDS of Bray-Curtis dissimilarities also
146 displayed that bacterial and fungal β -diversities in O horizon were significantly ($p <$ 0.005,
147 PERMANOVA) different from those in A horizon (Table S4).

148 ***Effect of fire on differential OTUs***

149 Differential abundances in bacterial and fungal community composition after wildfire across
150 a chronosequence were identified (Fig. 2). The enrichment and depletion effects on bacterial
151 communities in O (20 vs. 0) and A (13 vs. 1) horizons were the most remarkable in the
152 recent burnt soils (RF). There are no differential abundances in bacterial community
153 composition after middle-term and old fires, compared with the unburnt soil (Table S5).

154 After fire (RF vs. UF), the changes of bacterial composition were mostly positive. In the
155 period of 0-4 years after fire, differential OTUs had at least 3.5-fold enrichment in bacteria in
156 both O and A horizons (Fig. 2a). There were a number of spore-forming bacterial taxa,
157 including the group of Terrabacteria (Battistuzzi and Hedges, 2009). Terrabacterial phyla
158 (Order Bacillales and Order Micromonosporales) were only significant in O horizon but not
159 found in A horizon. However, the abundance of many fungal OTUs (e.g. Ascomycota and
160 Basidiomycota) were declined significantly after recent fires, especially in A horizon (Fig.
161 2b). Similarly, the RF group also had the most enriched and depleted fungal OTUs in O (17
162 vs. 7) and A (11 vs. 20) horizons compared with the unburnt soils (Table S6). There were
163 more fungal differential OTUs in A horizon than in O horizon (Fig. 2b). Moreover, the MF
164 group had only one significantly enriched OTU while the OF group had only one depleted
165 OTU (Table S6).

166 ***Bacterial and fungal co-occurrence networks***

167 To identify the co-occurrence patterns of bacterial and fungal communities in soils with
168 different fire histories, we constructed bacterial and fungal networks for the two horizons.
169 The networks displayed remarkable differences in their structure and topology (Figs 3 & 4).
170 The bacterial network of the UF group in O horizon consisted of 64 nodes (i.e. taxa) and 57
171 edges (associations between taxa), while the RF group in O horizon consisted of 60 nodes
172 and 30 edges. In stark contrast, the OF group in O horizon consisted of 222 nodes and 380
173 edges (Table 1). The average degree and the average number of neighbors of bacterial
174 network in the OF group were also considerably higher than the UF and other fire-affected

175 groups. The higher complexity and connectivity in the OF bacterial network showed that the
176 special and complicate modules had been formed during long periods after wildfire (Fig. 2d).
177 There were similar results in A horizon that more complex coupling among bacteria occurred
178 during the post-fire succession process (Figs S5a & d).

179 Multiple topological properties of the fungal co-occurrence networks pronouncedly
180 varied in different fire histories (Table S7). The fungal network of the OF was largest with
181 152 nodes and 253 edges and had the highest average degree in O horizon (Table 1). Higher
182 complexity in the OF group was visible, indicating that many fungi had developed a number
183 of associations (Fig. 3). We observed that both modularity and average path length of fungal
184 network in the OF group were highest in O horizon, showing more structured fungal
185 communities within the network. Unlike bacterial patterns, fungal co-occurrence patterns of
186 the OF group were more modular than the corresponding unburnt soil networks. Moreover,
187 the connections between fungi in A horizon were gradually increased after wildfire
188 disturbance (Fig. S6).

189 ***The effects of fire-derived charcoal***

190 To gain insight into the fire legacy effect, we compared microbial properties in charcoal
191 samples with soils. Despite common patterns in the relative abundance of main taxa at the
192 phylum level, microbial community composition, especially for fungi, had some apparent
193 divergences between charcoal and soil, and between different fire histories (Fig. 1d). The
194 predominant phyla (ascomycota, basidiomycota, and mucoromycota) significantly differed
195 in their relative abundances in the charcoal and unburnt soils (Table S3). Unlike burnt soils,

196 charcoal samples had an enhanced colonization of differential bacterial and fungal OTUs
197 from the unburnt controls (Fig. 2c & d). On the other hand, the burnt soil samples of O
198 horizon had more similar microbial composition to the unburnt sites than to charcoal samples,
199 as indicated by nearly zero differential OTUs (Fig. 2c & d). Bray-Curtis dissimilarity was
200 used to assess the structure of microbial community colonized on the surface of fire-derived
201 charcoal or inhabited within the soil. NMDS ordination revealed significant differences ($p <$
202 0.001 , PERMANOVA) among charcoal, unburnt, and burnt soil samples (Fig. S4) for
203 bacteria ($R^2 = 0.30$, *pseudo F* = 23.5) and fungi ($R^2 = 0.19$, *pseudo F* = 13.2). These results
204 showed that the charcoal colonized distinct microbial community, as fire legacy effect. The
205 explanation of variance to 40.7% in the soil fungal community structure on SEM (including
206 charcoal microbial community) while explanation of variance decreased slightly to 41.4%
207 for bacterial community structure (Fig. 6a & b). Based on SEM, fire significantly affected
208 bacterial and fungal community in charcoal and bacterial community in charcoal also
209 significantly affected bacterial community in soil.

210 ***The influence of soil properties on microbial community***

211 The relationships between soil properties and soil microbial community structure varied with
212 different fire histories and horizons (Fig. 5). Soil pH, soil organic matter, total N, total P,
213 dissolved organic C, and dissolved organic N had significant relationships with bacterial
214 community structure in both O and A horizons of the OF group. However, in the RF and MF
215 groups, most of soil factors had no significant correlations with bacterial community
216 structure. Soil factors, such as pH, soil organic matter, total P, dissolved organic C and

217 dissolved organic N, significantly correlated with the structures of fungal community in O
218 horizon (Fig. 5). More soil factors had significant relationships with bacterial than fungal
219 community structure. In O horizon of the OF group, bacterial community structure correlated
220 highest with pH ($r = 0.8095$, $p = 0.0001$) while fungal community structure correlated
221 highest with dissolved organic C ($r = 0.6691$, $p = 0.0001$) across all the environmental
222 variables (Table S8).

223

224 **Discussion**

225 *Temporal patterns of bacterial and fungal communities in response to fire*

226 This study demonstrates that wildfire has an impact on bacterial and fungal composition, and
227 thereby modulates the patterns of soil microbial community succession over three decades
228 after fire in the boreal forest region. Translated spatial differences into temporal changes, our
229 results clearly show that temporal patterns of bacterial and fungal communities occur across
230 a chronosequence. Supported by previous findings about bacterial diversity shifts under fire
231 disturbance (7), wildfire does not show significant impacts on microbial α -diversity
232 (Shannon index and OTU richness) both in O and A horizons. Only fungal α -diversity in A
233 horizon significantly declined by the recent fires, corresponding with fungal co-occurrence
234 networks in this horizon (Fig. S6b). This is in agreement with previous studies that fungal
235 diversity changed significantly in 2 years after fire across a 152-year chronosequence (41). In
236 addition, the fire-related effect on species richness of ectomycorrhizal fungi has not been
237 observed on many fire-affected soils, and mycorrhizal colonization of roots may increase

238 after a fire event (42). However, in Alaskan boreal forests, soil fungal community diversity
239 (Shannon index) varied significantly only in the 90-year-old site, not in sites with
240 shorter-term fire history across a 100-year chronosequence (43). Therefore, soil bacterial and
241 fungal community diversities have concomitantly and rapidly recovered, suggesting that fire
242 only has a limited impact on bacterial and fungal diversity across a chronosequence.

243 In general, wildfire is reported to have a distinct effect on soil microbial community
244 composition. Our results demonstrate that the bacterial and fungal community compositions
245 differ among fire history. Bacterial and fungal community structures have an underlying
246 trend clustered with increasing TSF (Fig. S3), which means fire has a significant effect on
247 soil microbial structure across a chronosequence. This phenomenon goes beyond previous
248 results reporting that bacterial community altered dramatically after 16 weeks since fire (44)
249 and could return to similar levels of the controls after 11 years (45). Yet, there are different
250 temporal patterns for bacteria and fungi under fire disturbance. Bacteria under the MF and
251 UF groups were closely clustered in the NMDS plot. Their co-occurrence networks and
252 topological networks are similar to those of other fire-affected groups in both O and A
253 horizons (Fig. S5). However, fungal co-occurrence always differs between the fire-affected
254 groups and the unburnt group, suggesting that the fungal communities still change after 29
255 years since fire. Our partial Mantel data show that bacterial community structure is
256 correlated with many soil factors in the OF group but correlated with few soil factors in the
257 RF and MF groups. These results indicate that fire-derived changes in soil properties become
258 the major determinants of soil microbial community composition in the post-fire succession

259 process. Our analysis further suggests that soil properties have significant effects on soil
260 microbial community especially for bacterial community as environmental filters from 17
261 years after fire. Furthermore, bacterial community was complexly modulated by soil factors
262 more significantly than fungal community.

263 Soil microbial co-occurrence networks, that provide insight into ecological interactions
264 among taxa, were influenced by fire history. The major changes in topological properties of
265 the bacterial and fungal networks have more nodes and edges under the OF group than the
266 UF group (Figs 2, 3, S5 & S6), predicting that microbial communities have more
267 connections and complex functional modules under fire disturbance. Fire reshapes soil
268 microbial network and makes microbial interaction stronger in both O and A horizons after
269 17 years since fire. In addition, during 1-15 years since fire (RF and MF groups), taxa only
270 form weak relationships and lower-complexity modules in most of the samples. Clearly,
271 bacteria and fungi need more than 15 years to recover and reconstruct ecological functioning
272 after fire events. Our results are in contrast to the short-term experiments (within 12 months
273 after fire) showing that fire did not alter co-occurrence network generally (46).

274 ***Specific bacterial and fungal taxa in response to fire***

275 By using differential OTU abundance analysis across a chronosequence, we observed that
276 only RF groups serve an enrichment/depletion role for microbial OTUs relative to the
277 unburnt soil. There are few differential OTUs found in the MF and OF groups although
278 microbial co-occurrence network in the MF and OF groups is significantly different from
279 that in the UF group. These results indicate the number of most OTUs in post-fire boreal soil

280 back to pre-fire level after 7 years. Many bacterial species are able to cope with extreme
281 environments such as high temperature by forming resistant substances such as spores and
282 endospores (10). We observed that the RF group serves an enrichment role for a subset of
283 bacterial OTUs relative to the unburnt soil. The recent fires increased the number of
284 spore-forming OTUs in the order Micromonosporales (phylum Actinobacteria) and the
285 Gram-positive Bacillales (phylum Firmicutes). The Bacillales can generate dormant and
286 resistant spores that make them survive under harsh conditions for years (47). The Bacillales,
287 that consist of many spore-formers, allow themselves to recover better and faster from
288 stresses. Micromonosporales (genus Actinoplanes) also form spores by fragmentation to
289 reproduce (48). These post-fire increases in the number of bacterial OTUs are more in O
290 horizon than in A horizon, which could be linked to higher fire severity or upper fire loading
291 density in O horizon. The results of these differential bacterial OTUs are consistent with the
292 findings of many previous studies showing pronounced effects on soil bacterial community
293 in the short-term, mostly within 36 months after fire (6, 10, 49). Furthermore, apart from fire
294 severity, environmental factors (soil pH, soil water content, total C, dissolved C and N) are
295 also key determinants driving soil bacterial community. Some bacterial OTUs depend on
296 plant-soil feedback (50) and are strongly affected by aboveground community during forest
297 regeneration (51, 52).

298 We observed that there are many depleted ectomycorrhizal OTUs (such as Family
299 Tricholomataceae) in O and A horizons, which is consistent with a previous study showing
300 that forest fire caused a significant loss of ectomycorrhizal fungi biomass in the litter and

301 organic horizons (53). Severe fire consumes the most portion of humus on soil surface and
302 kills most of trees, and hence the death of ectomycorrhizal fungi although some
303 ectomycorrhizal fungi have sporophytes (54). Moreover, during the early successional phase
304 after fire disturbance, the regeneration has not reached pre-wildfire levels in the Great
305 Khingan Mountain. The impact of vegetative recovery on fungal community may play an
306 important role in enriched OTUs in O horizon. Furthermore, most of differential OTUs in A
307 horizon are depleted probably because vegetation including herbaceous species and shrubs
308 still cannot influence soil microbes in A horizon. However, many fungal species can still
309 survive and largely remain intact even after a high-intensity wildfire. The regeneration of
310 post-fire pyrophilous ascomycetes is well documented for the reasons such as the tolerance
311 of high pH and other physicochemical consequences of wildfire (42).

312 *Fire legacy effect on soil microbial community*

313 Our study, for the first time, examined the effect of fire-derived natural charcoal on soil
314 microbial community along a chronosequence in a boreal forest region, and showed that the
315 microbial community colonized on the charcoal can influence the responses of soil microbial
316 community to wildfire. Due to specific porous structure and absorbed nutrients, charcoal is a
317 habitat for soil microbes and these microbes are significantly different from those in the
318 burnt soils (Fig. S4). The major taxa and α -diversity of microbes in the charcoal (Figs 1 &
319 S4) are also distinct with the burnt forest soils, highlighting a new and unique habitat
320 produced by fire (28). Moreover, there are more than 680 bacterial OTUs and more than 120
321 fungal OTUs in the charcoal samples (Fig. 2c & d). Together these observations provide

322 strong evidence that fire-derived charcoal create distinct environments (including abiotic
323 factors and biotic factors) as the habitat for fire-responsive microbes. Because of its higher
324 pH value, charcoal causes a localized increase in soil pH (55, 56) and shifts the way of soil
325 pH effect on soil microbial community to some extent. Charcoal microbial community has a
326 significantly positive relationship with soil bacterial community (Fig. 6a), indicating that
327 microbes in soil are more likely impacted by post-fire charcoal. Charcoal microbial
328 community is also related to soil pH, dissolved organic C and total P which regulate
329 microbial activity with nutrient cycling. Previous studies have shown that charcoal absorbs
330 some labile C compounds to inhibit or activate some microbial processes (29, 40, 57).
331 However, our present study has not found that dissolved organic C has any specific
332 relationship with bacteria or fungi in the charcoal. Based on SEM, fire (represented by TSF)
333 affects soil microbial community via the indirect effect of charcoal. Furthermore, charcoal
334 community structure varied with TSF showing that the ecological recovery processes after
335 fire occur in the charcoal particles and the fire legacy effect gradually changes over time.

336 ***Conclusions***

337 This study provides the key evidence how large-scale wildfire events and the derived
338 charcoal (fire legacy effect) affect soil microbial community along a 29-year chronosequence
339 of the boreal forest region in northeast China. It reveals that wildfire significantly impacts on
340 bacterial and fungal community structures. The complexity and connectivity of bacterial and
341 fungal communities were significantly enhanced from 17 years after fire when soil pH, total
342 C and N, soil water content, and dissolved C and N are key determinants of soil bacterial and

343 fungal communities. Furthermore, the charcoal formed at fire events and its colonized
344 microbes have important ecological functions of mediating post-fire successional boreal
345 forests although our current knowledge of the succession patterns about soil microbes
346 post-fire and charcoal functions in flammable boreal forest ecosystems is limited. Our
347 research highlights an important step forward to clarifying the effects of fire and its charcoal
348 on soil microbial community. Further research may aim at (i) identifying specific
349 physiological and biochemical processes in response to fire events and fire history and (ii)
350 distinguishing charcoal microbial community from post-fire soil microbial community in
351 finer details.

352

353 **Materials and methods**

354 *Study area and experimental design*

355 This study was carried out in the boreal forest region, the Great Khingan Mountains
356 (50°10'N-53°33'N, 121°12'E-127°00'E), north-eastern China (Fig. S1). The mean annual air
357 temperature in this region is -2 to -4 °C and the mean annual rainfall is 350~500 mm (18).

358 The soil at the sampling sites is classified as dark brown forest soil (Inceptisol) and the
359 parent material is granite bedrock (45, 58). The vegetation of this region is representative of
360 boreal coniferous forests, forming the southern extension of the Siberian boreal forests.

361 Transition plant species, dominated by larch (*Larix gmelini*), are late-successional and
362 widely distributed from a wildfire (14) or harvesting. Broadleaf trees as the pioneer species,

363 mainly birch (*Betula platyphylla*), are the earliest re-generation in the post-fire soil. They are
364 succeeded by boreal conifer tree species in the wildfire-disturbed boreal forest ecosystems.

365 We selected a total number of 44 plots (50 × 50 m) at altitudes ranging from 209 m to
366 466 m a.s.l. (Table S1) in July 2015. Plots including fire-affected and unaffected groups,
367 were located at the National Reserve and State-owned Forest with minimal human
368 interventions (e.g. prescribed fire, and logging). Due to the difficulty of long-term studies
369 and specificity of fire researches, we followed a classical space-for-time substitution method
370 to assess the impacts of fire and its legacy charcoal. Fire-affected groups consisted of a
371 chronosequence of burnt forests that were representative of different TSF (59), considered as
372 different fire histories. These affected groups were divided into Recent Fires (RF: 0-4 years
373 since fire, n=9), Middle-term Fires (MF: 7-15 years since fire, n=12), and Old Fires (OF:
374 17-29 years since fire, n=12) based on TSF. Each fire-affected site was considered as an
375 independent replicate and more information about TSF is given in the Supplementary
376 Information.

377 According to the historical records from the State Forest Administration and remote
378 sensing images, each of the fire-affected sampling sites had suffered one highly severe
379 wildfire during 1915-2015. The highly severe wildfire here means destroying at least one
380 hundred hectares of forests and removing litter and organic soil layer (60). To save
381 expenditure, we only sampled fire-unaffected controls (UF, n=11) nearby the burnt plots
382 with distinct burnt borders. The minimum distance between each plot was 100 meters. For
383 reducing spatial heterogeneity, all plot samples were comprised of more than five

384 subsamples (10×10 m) and each subsample was homogenized by at least ten soil cores.

385 Given that plots were structurally analogous, all samples followed by unified strict selection

386 criteria including a slope of $< 15^\circ$, the same soil type and a similar landform.

387 ***Sampling, charcoal separation and physicochemical analysis***

388 During sampling, the litter layer (a laminated mixing of small twigs, roots and fungal

389 hyphae) was first removed and the organic layer (O horizon) comprising dark colored

390 materials with fewer small roots and charcoal pieces were collected. Then, the top 15-cm

391 mineral soil (A horizon) was collected using a soil auger (4 cm diameter). Soils were placed

392 in the aseptic plastic bags for subsequent processing. Samples were transported to laboratory

393 on ice, immediately sieved through a 2-mm mesh and stored at -80°C (for DNA extraction)

394 and 4°C (for physicochemical analyses). Soil physicochemical properties included total C

395 (%), N (%), P (mg kg^{-1}), soil organic matter (SOM), pH, soil gravimetric water content,

396 available P, dissolved organic C, dissolved inorganic C, dissolved organic N and

397 exchangeable Na, K, Mg, and Ca. Detailed information on soil analyses is given in the

398 Supplementary Information.

399 To minimize disturbance of microbial community, we adopted a direct hand-picked

400 method (61, 62), rather than the water-floating method, to separate charcoal particles from

401 the O-horizon soil. These “charcoal” particles in this study were defined as solid residual

402 pieces derived from the pyrolysis of plant biomass by fire. Many researches showed that 1-3

403 mm char-particles (including biochar, charcoal) are easily extracted from soil (63–65), and in

404 view of this phenomenon, we separated macroscopic charcoal particles (≥ 1 mm in diameter)

405 from the sieved soil. We used a specific tweezer (N5, Dumont, Switzerland), which could
406 isolate visible charcoal particles larger than 50 μm under microscope. Followed by the same
407 criteria of microscopy and mineralogy, we collected charcoal approximately 15 g per plot for
408 later analyses and characterization. The details of charcoal isolation from soil are presented
409 in Supplementary Information. Collected charcoal pieces were crushed gently, homogenized
410 and stored at $-80\text{ }^{\circ}\text{C}$ for DNA extraction and physicochemical analyses.

411 *Illumina sequencing and bioinformatic analysis*

412 DNA was extracted from all soil (0.50 g) and charcoal (0.2-0.4 g) samples using the MP
413 FastDNA SPIN Kit for soil (MP Biomedicals, Solon, OH, USA) according to the
414 manufacturer's protocol. Extracted DNA was stored at -20°C for a maximum of one week
415 until amplicon library preparation began. PCR amplification was performed on the V4-V5
416 region of bacterial 16S rRNA gene using primers 515F/907R (66, 67) and the V4 region of
417 18S rRNA using primers 528F/706R (68, 69). The concentration and purity of DNA were
418 monitored on 1% agarose gels. According to actual concentration, DNA was diluted to about
419 $1\text{ ng }\mu\text{L}^{-1}$ using sterile water. 16S /18S rRNA genes were amplified used specific primers
420 with the barcode. All PCR reactions were carried out in a volume of 30 μL using 15 μL of
421 Phusion[®] High-Fidelity PCR Master Mix (New England Biolabs), 0.2 μM of forward
422 primers, 0.2 μM of reverse primers and 10 μL sample DNA. The PCR consisted of $98\text{ }^{\circ}\text{C}$ for
423 1 min, then 30 cycles at 98°C for 10 s, $50\text{ }^{\circ}\text{C}$ for 30 s, $72\text{ }^{\circ}\text{C}$ for 30 s, and finally $72\text{ }^{\circ}\text{C}$ for 5
424 min. PCR products were mixed in equidensity ratios and purified before sequencing libraries
425 were generated. The library quality was assessed on the Qubit[®] 2.0 Fluorometer (Thermo

426 Scientific) and Agilent Bioanalyzer 2100 system. At last, the library was sequenced on an
427 Illumina HiSeq 2500 and 250 bp paired-end reads were generated.

428 Briefly, the raw sequences from HiSeq platform were merged by using FLASH (70) and
429 were assigned to each sample according to the specific barcodes. The sequences were
430 pre-filtered and removed chimeras (71) by QIIME's (72) quality filters. All effective
431 sequences were then clustered into operational taxonomic units (OTUs) by Uparse (73)
432 based on 97% similarity level. We picked the highest frequent sequence for each OTU as
433 their representative sequences. We used the Ribosomal Database Project's classifier (74)
434 against the SILVA 132 database (75) to annotate taxonomic information for bacteria (16S)
435 and eukaryotes (18S). RDP classifier bootstrap confidence values were 0.8-1. We
436 normalized the read counts in OTU table by rarefying to the minimum sequence number
437 within all samples. Moreover, to parse fungal OTUs by ecological guilds, we used
438 FUNGuild to annotate fungal taxa into different functional groups (i.e. saprotroph,
439 pathotroph and symbiotroph) (76). Raw sequence data were deposited into the Genome
440 Sequence Archive (GSA) database under accession numbers CRA002389 that are publicly
441 accessible.

442 ***Statistical analysis***

443 All statistical analyses were performed using R v. 3.5.3 (77) using the rarefied data on
444 the following packages: phyloseq (78), ggplot2 (79), reshape2 (80), plyr (81), except for the
445 analysis using non-rarefied data (82). Here, α -diversity metrics included Shannon index and
446 OTU richness (namely species richness), which were calculated using the "plot_richness"

447 function (phyloseq) and Picante package (83). The vegan package (84) was used to assess
448 β -diversity via non-metric multidimensional scaling (NMDS) ordination using “metaMDS”
449 function, based on Bray-Curtis dissimilarities. Bacterial and fungal communities in charcoal
450 and soil samples were determined separately by the differential abundance analysis
451 compared with the unburnt soils in each soil horizon. The package ‘DESeq2’ (85) was used
452 to calculate the differential abundance (Log₂ fold change in relative abundance of OTUs)
453 using non-rarefied data. DESeq2 was run using the Wald test and an alpha of 0.01.

454 To reduce network complexity, we removed OTUs with relative abundances less than
455 0.001% of the total number of bacterial and fungal sequences, respectively. We then
456 calculated all possible Spearman’s rank correlations between OTUs more than three
457 sequences using the WGCNA package (86). The nodes in the networks represent microbial
458 taxa (OTUs) and the edges represent significant correlations between the nodes. The *p* values
459 were adjusted using the the Benjamini-Hochberg false discovery rate (FDR) controlling
460 procedure (87), as implemented in the multtest package (88). A valid co-occurrence network
461 was considered a statistically robust correlation between taxa when the correlation
462 coefficients threshold was above 0.8 and FDR-adjusted *p* value was below 0.01. Network
463 visualization was conducted using Cytoscape version 3.8.0 (89), and topographical
464 properties (including average degree, average path length, network diameter, and clustering
465 coefficient) of networks were calculated using Gephi software (90).

466 To investigate the importance of fire effect and soil properties, a permutational
467 multivariate analysis of variance (PERMANOVA) was performed on Bray-Curtis

468 dissimilarities using the “adonis” function of vegan ($n_{perm}=9999$). Partial Mantel tests were
469 conducted to test the effects of soil properties on the compositional similarity of OTUs using
470 vegan package. Moreover, the effects of spatial variation (i.e. longitude and latitude) were
471 controlled while testing the correlation of soil factors and the microbial community of burnt
472 and unburnt soils. Soil properties, the relative abundances of taxa and α -diversity indices
473 were performed by ANOVA, and pair-wise differences were assessed by Tukey’s HSD
474 post-hoc tests as implemented in agricolae package. The Kruskal-Wallis test instead was
475 used for the data with abnormal distribution or homoscedastic. To further test the effect of
476 fire (including TSF and charcoal) on soil factors and soil community structure of bacteria
477 and fungi, we built Structural Equation Modeling (SEM) by “sem” function from the lavaan
478 package. A priori models based on the results of linear mixed-effect model (Fig. S7). We
479 used fire effect (TSF) and fundamental edaphic factors (pH, dissolved organic C) as
480 exogenous variables, and soil total N and P, and microbial community in charcoal as
481 endogenous variables (Fig. 6). Exploratory analyses showed dissolved organic C had a better
482 fit and explanation than soil organic matter or total C. Importantly, the microbial community
483 colonized on the separated charcoal particles were hypothesized as key links between fire
484 effect and soil microbial community. The goodness of fit was considered on the basis of
485 chi-square test and root-mean-square error of approximation (RMSEA). A model could be
486 accepted when p -value > 0.05 , comparative fit index (CFI) > 0.95 and RMSEA < 0.05 .

487

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491 **References**

- 492 1. Bowman DMJS, Balch JK, Artaxo P, Bond WJ, Carlson JM, Cochrane MA, D'Antonio
493 CM, DeFries RS, Doyle JC, Harrison SP, Johnston FH, Keeley JE, Krawchuk MA,
494 Kull CA, Marston JB, Moritz MA, Prentice IC, Roos CI, Scott AC, Swetnam TW, van
495 der Werf GR, Pyne SJ. 2009. Fire in the Earth System. *Science* 324:481–484.
- 496 2. Bond W, Keeley J. 2005. Fire as a global ‘herbivore’: the ecology and evolution of
497 flammable ecosystems. *Trends in Ecology & Evolution* 20:387–394.
- 498 3. Certini G. 2005. Effects of fire on properties of forest soils: a review. *Oecologia*
499 143:1–10.
- 500 4. Dooley SR, Treseder KK. 2012. The effect of fire on microbial biomass: a
501 meta-analysis of field studies. *Biogeochemistry* 109:49–61.
- 502 5. Luo Y, Yu Z, Zhang K, Xu J, Brookes PC. 2016. The properties and functions of
503 biochars in forest ecosystems. *J Soils Sediments* 16:2005–2020.
- 504 6. Smith NR, Kishchuk BE, Mohn WW. 2008. Effects of Wildfire and Harvest
505 Disturbances on Forest Soil Bacterial Communities. *Applied and Environmental*
506 *Microbiology* 74:216–224.
- 507 7. Sun H, Santalahti M, Pumpanen J, Köster K, Berninger F, Raffaello T, Asiegbu FO,
508 Heinonsalo J. 2016. Bacterial community structure and function shift across a northern
509 boreal forest fire chronosequence. *Scientific Reports* 6.
- 510 8. Daniel RM, Cowan DA. 2000. Biomolecular stability and life at high temperatures.
511 *Cellular and Molecular Life Sciences* 57:250–264.
- 512 9. Holden SR, Treseder KK. 2013. A meta-analysis of soil microbial biomass responses to
513 forest disturbances. *Frontiers in Microbiology* 4.
- 514 10. Pérez-Valera E, Goberna M, Verdú M. 2019. Fire modulates ecosystem functioning
515 through the phylogenetic structure of soil bacterial communities. *Soil Biology and*
516 *Biochemistry* 129:80–89.
- 517 11. Taş N, Prestat E, McFarland JW, Wickland KP, Knight R, Berhe AA, Jorgenson T,
518 Waldrop MP, Jansson JK. 2014. Impact of fire on active layer and permafrost microbial
519 communities and metagenomes in an upland Alaskan boreal forest. *The ISME Journal*
520 8:1904–1919.

- 521 12. Randerson JT, Liu H, Flanner MG, Chambers SD, Jin Y, Hess PG, Pfister G, Mack MC,
522 Treseder KK, Welp LR, Chapin FS, Harden JW, Goulden ML, Lyons E, Neff JC,
523 Schuur EAG, Zender CS. 2006. The Impact of Boreal Forest Fire on Climate Warming.
524 *Science* 314:1130–1132.
- 525 13. Abatzoglou JT, Williams AP, Boschetti L, Zubkova M, Kolden CA. 2018. Global
526 patterns of interannual climate-fire relationships. *Global Change Biology*
527 24:5164–5175.
- 528 14. Liu Z, Yang J, Chang Y, Weisberg PJ, He HS. 2012. Spatial patterns and drivers of fire
529 occurrence and its future trend under climate change in a boreal forest of Northeast
530 China. *Global Change Biology* 18:2041–2056.
- 531 15. van Mantgem PJ, Nesmith JCB, Keifer M, Knapp EE, Flint A, Flint L. 2013. Climatic
532 stress increases forest fire severity across the western United States. *Ecology Letters*
533 16:1151–1156.
- 534 16. Shungart HH, Leemans Rik, Bonan GB. 1992. *A Systems Analysis of the Global
535 Boreal Forest*. Cambridge University Press, Cambridge.
- 536 17. Xu H. 1998. *Forests in Daxinganling Moutains China*. Science Press, Beijing, China.
- 537 18. Hu H, Wei S, Sun L. 2012. Estimation of carbon emissions due to forest fire in
538 Daxing’an Mountains from 1965 to 2010. *Chinese Journal of Plant Ecology*
539 36:629–644.
- 540 19. Chang Y, He HS, Hu Y, Bu R, Li X. 2008. Historic and current fire regimes in the
541 Great Xing’an Mountains, northeastern China: Implications for long-term forest
542 management. *Forest Ecology and Management* 254:445–453.
- 543 20. Wang X, He HS, Li X. 2007. The long-term effects of fire suppression and reforestation
544 on a forest landscape in Northeastern China after a catastrophic wildfire. *Landscape and
545 Urban Planning* 79:84–95.
- 546 21. Flannigan MD, Krawchuk MA, de Groot WJ, Wotton BM, Gowman LM. 2009.
547 Implications of changing climate for global wildland fire. *International Journal of
548 Wildland Fire* 18:483.

- 549 22. Santín C, Doerr SH, Kane ES, Masiello CA, Ohlson M, de la Rosa JM, Preston CM,
550 Dittmar T. 2016. Towards a global assessment of pyrogenic carbon from vegetation
551 fires. *Global Change Biology* 22:76–91.
- 552 23. Bird MI, Wynn JG, Saiz G, Wurster CM, McBeath A. 2015. The Pyrogenic Carbon
553 Cycle. *Annual Review of Earth and Planetary Sciences* 43:273–298.
- 554 24. Preston CM, Schmidt MWI. 2006. Black (pyrogenic) carbon: a synthesis of current
555 knowledge and uncertainties with special consideration of boreal regions.
556 *Biogeosciences* 3:397–420.
- 557 25. Gavin DG, Hallett DJ, Hu FS, Lertzman KP, Prichard SJ, Brown KJ, Lynch JA,
558 Bartlein P, Peterson DL. 2007. Forest fire and climate change in western North
559 America: insights from sediment charcoal records. *Frontiers in Ecology and the*
560 *Environment* 5:499–506.
- 561 26. Ohlson M, Dahlberg B, Økland T, Brown KJ, Halvorsen R. 2009. The charcoal carbon
562 pool in boreal forest soils. *Nature Geoscience* 2:692–695.
- 563 27. Hart S, Luckai N. 2013. Charcoal function and management in boreal ecosystems.
564 *Journal of Applied Ecology* 50:1197–1206.
- 565 28. Pietikainen J, Kiikkila O, Fritze H. 2000. Charcoal as a habitat for microbes and its
566 effect on the microbial community of the underlying humus. *Oikos* 89:231–242.
- 567 29. Wardle DA, Zackrisson O, Nilsson M-C. 1998. The charcoal effect in Boreal forests:
568 mechanisms and ecological consequences. *Oecologia* 115:419–426.
- 569 30. Bryanin S, Abramova E, Makoto K. 2018. Fire-derived charcoal might promote fine
570 root decomposition in boreal forests. *Soil Biology and Biochemistry* 116:1–3.
- 571 31. Hestrin R, Torres-Rojas D, Dynes JJ, Hook JM, Regier TZ, Gillespie AW, Smernik RJ,
572 Lehmann J. 2019. Fire-derived organic matter retains ammonia through covalent bond
573 formation. *Nature Communications* 10.
- 574 32. Wardle DA, Nilsson M-C, Zackrisson O. 2008. Fire-derived charcoal causes loss of
575 forest humus. *Science* 320:629–629.
- 576 33. Preston CM. 2009. Fire’s black legacy. *Nature Geoscience* 2:674–675.

- 577 34. Zimmermann M, Bird MI, Wurster C, Saiz G, Goodrick I, Barta J, Capek P,
578 Santruckova H, Smernik R. 2012. Rapid degradation of pyrogenic carbon. *Global*
579 *Change Biology* 18:3306–3316.
- 580 35. Butler OM, Elser JJ, Lewis T, Mackey B, Chen C. 2018. The phosphorus-rich signature
581 of fire in the soil-plant system: a global meta-analysis. *Ecology Letters* 21:335–344.
- 582 36. Bowd EJ, Banks SC, Strong CL, Lindenmayer DB. 2019. Long-term impacts of
583 wildfire and logging on forest soils. *Nature Geoscience* 12:113–118.
- 584 37. Cui X, Gao F, Song J, Sang Y, Sun J, Di X. 2014. Changes in soil total organic carbon
585 after an experimental fire in a cold temperate coniferous forest: A sequenced
586 monitoring approach. *Geoderma* 226–227:260–269.
- 587 38. Hu T, Sun L, Hu H, Weise DR, Guo F. 2017. Soil respiration of the Dahurian Larch
588 (*Larix gmelinii*) forest and the response to fire disturbance in Da Xing’an Mountains,
589 China. *Scientific Reports* 7:2967.
- 590 39. Blois JL, Williams JW, Fitzpatrick MC, Jackson ST, Ferrier S. 2013. Space can
591 substitute for time in predicting climate-change effects on biodiversity. *Proceedings of*
592 *the National Academy of Sciences* 110:9374–9379.
- 593 40. Carter Z, Sullivan B, Qualls R, Blank R, Schmidt C, Verburg P. 2018. Charcoal
594 Increases Microbial Activity in Eastern Sierra Nevada Forest Soils. *Forests* 9:93.
- 595 41. Sun H, Santalahti M, Pumpanen J, Köster K, Berninger F, Raffaello T, Jumpponen A,
596 Asiegbu FO, Heinonsalo J. 2015. Fungal community shifts in structure and function
597 across a boreal forest fire chronosequence. *Applied and Environmental Microbiology*
598 81:7869–7880.
- 599 42. Cairney JWG, Bastias BA. 2007. Influences of fire on forest soil fungal communities.
600 *Can J For Res* 37:207–215.
- 601 43. Holden SR, Gutierrez A, Treseder KK. 2013. Changes in Soil Fungal Communities,
602 Extracellular Enzyme Activities, and Litter Decomposition Across a Fire
603 Chronosequence in Alaskan Boreal Forests. *Ecosystems* 16:34–46.
- 604 44. Ferrenberg S, O’Neill SP, Knelman JE, Todd B, Duggan S, Bradley D, Robinson T,
605 Schmidt SK, Townsend AR, Williams MW, Cleveland CC, Melbourne BA, Jiang L,

- 606 Nemergut DR. 2013. Changes in assembly processes in soil bacterial communities
607 following a wildfire disturbance. *The ISME Journal* 7:1102–1111.
- 608 45. Xiang X, Shi Y, Yang J, Kong J, Lin X, Zhang H, Zeng J, Chu H. 2015. Rapid recovery
609 of soil bacterial communities after wildfire in a Chinese boreal forest. *Scientific*
610 *Reports* 4.
- 611 46. Pérez-Valera E, Goberna M, Faust K, Raes J, García C, Verdú M. 2017. Fire modifies
612 the phylogenetic structure of soil bacterial co-occurrence networks. *Environmental*
613 *Microbiology* 19:317–327.
- 614 47. Paredes-Sabja D, Setlow P, Mahfuzur S. 2011. Germination of spores of Bacillales and
615 Clostridiales species: mechanisms and proteins involved. *Trends in Microbiology*
616 19:85–94.
- 617 48. Sharples GP, Williams ST, Bradshaw RM. 1974. Spore formation in the
618 Actinoplanaceae (Actinomycetales). *Arch Microbiol* 101:9–20.
- 619 49. Mikita-Barbato RA, Kelly JJ, Tate RL. 2015. Wildfire effects on the properties and
620 microbial community structure of organic horizon soils in the New Jersey Pinelands.
621 *Soil Biology and Biochemistry* 86:67–76.
- 622 50. Duhamel M, Wan J, Bogar LM, Segnitz RM, Duncritts NC, Peay KG. 2019. Plant
623 selection initiates alternative successional trajectories in the soil microbial community
624 after disturbance. *Ecol Monogr* 89.
- 625 51. Knelman JE, Graham EB, Trahan NA, Schmidt SK, Nemergut DR. 2015. Fire severity
626 shapes plant colonization effects on bacterial community structure, microbial biomass,
627 and soil enzyme activity in secondary succession of a burned forest. *Soil Biology and*
628 *Biochemistry* 90:161–168.
- 629 52. Wang C, Wang G, Wang Y, Rafique R, Ma L, Hu L, Luo Y. 2016. Fire alters
630 vegetation and soil microbial community in alpine meadow. *Land Degradation &*
631 *Development* 27:1379–1390.
- 632 53. Dahlberg A. 2002. Effects of fire on ectomycorrhizal fungi in Fennoscandian boreal
633 forests. *Silva Fenn* 36.

- 634 54. Glassman SI, Levine CR, DiRocco AM, Battles JJ, Bruns TD. 2016. Ectomycorrhizal
635 fungal spore bank recovery after a severe forest fire: some like it hot. *ISME J*
636 10:1228–1239.
- 637 55. Oguntunde PG, Fosu M, Ajayi AE, van de Giesen N. 2004. Effects of charcoal
638 production on maize yield, chemical properties and texture of soil. *Biol Fertil Soils*
639 39:295–299.
- 640 56. Yu M, Meng J, Yu L, Su W, Afzal M, Li Y, Brookes PC, Redmile-Gordon M, Luo Y,
641 Xu J. 2019. Changes in nitrogen related functional genes along soil pH, C and nutrient
642 gradients in the charosphere. *Science of The Total Environment* 650:626–632.
- 643 57. DeLuca TH, MacKenzie MD, Gundale MJ, Holben WE. 2006. Wildfire-Produced
644 Charcoal Directly Influences Nitrogen Cycling in Ponderosa Pine Forests. *Soil Science*
645 *Society of America Journal* 70:448.
- 646 58. Wang C, Gower ST, Wang Y, Zhao H, Yan P, Bond-Lamberty BP. 2001. The influence
647 of fire on carbon distribution and net primary production of boreal *Larix gmelinii*
648 forests in north-eastern China. *Global Change Biology* 7:719–730.
- 649 59. Oliveras I, Román-Cuesta RM, Urquiaga-Flores E, Quintano Loayza JA, Kala J,
650 Huamán V, Lizárraga N, Sans G, Quispe K, Lopez E, Lopez D, Cuba Torres I, Enquist
651 BJ, Malhi Y. 2018. Fire effects and ecological recovery pathways of tropical montane
652 cloud forests along a time chronosequence. *Global Change Biology* 24:758–772.
- 653 60. Ludwig SM, Alexander HD, Kielland K, Mann PJ, Natali SM, Ruess RW. 2018. Fire
654 severity effects on soil carbon and nutrients and microbial processes in a Siberian larch
655 forest. *Global Change Biology* 24:5841–5852.
- 656 61. Carvalho LC d. S, Fearnside PM, Nascimento MT, Barbosa RI. 2018. Amazon soil
657 charcoal: Pyrogenic carbon stock depends of ignition source distance and forest type in
658 Roraima, Brazil. *Global Change Biology* 24:4122–4130.
- 659 62. Nguyen BT, Lehmann J, Kinyangi J, Smernik R, Riha SJ, Engelhard MH. 2008.
660 Long-term black carbon dynamics in cultivated soil. *Biogeochemistry* 89:295–308.
- 661 63. Dai Z, Webster TM, Enders A, Hanley KL, Xu J, Thies JE, Lehmann J. 2017. DNA
662 extraction efficiency from soil as affected by pyrolysis temperature and extractable
663 organic carbon of high-ash biochar. *Soil Biology and Biochemistry* 115:129–136.

- 664 64. Lin Y, Munroe P, Joseph S, Kimber S, Van Zwieten L. 2012. Nanoscale
665 organo-mineral reactions of biochars in ferrosol: an investigation using microscopy.
666 *Plant and Soil* 357:369–380.
- 667 65. Ye J, Joseph SD, Ji M, Nielsen S, Mitchell DRG, Donne S, Horvat J, Wang J, Munroe
668 P, Thomas T. 2017. Chemolithotrophic processes in the bacterial communities on the
669 surface of mineral-enriched biochars. *The ISME Journal* 11:1087–1101.
- 670 66. Biddle JF, Fitz-Gibbon S, Schuster SC, Brenchley JE, House CH. 2008. Metagenomic
671 signatures of the Peru Margin subseafloor biosphere show a genetically distinct
672 environment. *Proceedings of the National Academy of Sciences* 105:10583–10588.
- 673 67. Stubner S. 2002. Enumeration of 16S rDNA of *Desulfotomaculum* lineage 1 in rice
674 field soil by real-time PCR with SybrGreenk detection. *Journal of Microbiological*
675 *Methods* 50:155–164.
- 676 68. Cheung MK, Au CH, Chu KH, Kwan HS, Wong CK. 2010. Composition and genetic
677 diversity of picoeukaryotes in subtropical coastal waters as revealed by 454
678 pyrosequencing. *The ISME Journal* 4:1053–1059.
- 679 69. Elwood HJ. 1985. The small-subunit ribosomal RNA gene sequences from the
680 Hypotrichous Ciliates *Oxytricha nova* and *Stylonychia pustulata*. *Molecular Biology*
681 *and Evolution* 2:399–410.
- 682 70. Magoc T, Salzberg SL. 2011. FLASH: fast length adjustment of short reads to improve
683 genome assemblies. *Bioinformatics* 27:2957–2963.
- 684 71. Haas BJ, Gevers D, Earl AM, Feldgarden M, Ward DV, Giannoukos G, Ciulla D,
685 Tabbaa D, Highlander SK, Sodergren E, Methe B, DeSantis TZ, The Human
686 Microbiome Consortium, Petrosino JF, Knight R, Birren BW. 2011. Chimeric 16S
687 rRNA sequence formation and detection in Sanger and 454-pyrosequenced PCR
688 amplicons. *Genome Research* 21:494–504.
- 689 72. Caporaso JG, Kuczynski J, Stombaugh J, Bittinger K, Bushman FD, Costello EK,
690 Fierer N, Peña AG, Goodrich JK, Gordon JI, Huttley GA, Kelley ST, Knights D,
691 Koenig JE, Ley RE, Lozupone CA, McDonald D, Muegge BD, Pirrung M, Reeder J,
692 Sevinsky JR, Turnbaugh PJ, Walters WA, Widmann J, Yatsunencko T, Zaneveld J,

- 693 Knight R. 2010. QIIME allows analysis of high-throughput community sequencing data.
694 Nature Methods 7:335–336.
- 695 73. Edgar RC. 2013. UPARSE: highly accurate OTU sequences from microbial amplicon
696 reads. Nature Methods 10:996–998.
- 697 74. Wang Q, Garrity GM, Tiedje JM, Cole JR. 2007. Naive bayesian classifier for rapid
698 assignment of rRNA sequences into the new bacterial taxonomy. Applied and
699 Environmental Microbiology 73:5261–5267.
- 700 75. Quast C, Pruesse E, Yilmaz P, Gerken J, Schweer T, Yarza P, Peplies J, Glöckner FO.
701 2012. The SILVA ribosomal RNA gene database project: improved data processing and
702 web-based tools. Nucleic Acids Research 41:D590–D596.
- 703 76. Nguyen NH, Song Z, Bates ST, Branco S, Tedersoo L, Menke J, Schilling JS, Kennedy
704 PG. 2016. FUNGuild: An open annotation tool for parsing fungal community datasets
705 by ecological guild. Fungal Ecology 20:241–248.
- 706 77. R Development Core Team. 2011. R: A language and environment for statistical
707 computing. R foundation for statistical computing Vienna, Austria.
- 708 78. McMurdie PJ, Holmes S. 2013. phyloseq: An R Package for reproducible interactive
709 analysis and graphics of microbiome census data. Plos one 8:e61217.
- 710 79. Wickham H. 2016. ggplot2. Springer International Publishing, Cham.
- 711 80. Wickham H. 2007. Reshaping data with the reshape Package. Journal of Statistical
712 Software 21.
- 713 81. Wickham H. 2011. The split-apply-combine strategy for data analysis. Journal of
714 Statistical Software 40.
- 715 82. McMurdie PJ, Holmes S. 2014. Waste not, want not: why rarefying microbiome data Is
716 inadmissible. PLoS Comput Biol 10:e1003531.
- 717 83. Kembel SW, Cowan PD, Helmus MR, Cornwell WK, Morlon H, Ackerly DD,
718 Blomberg SP, Webb CO. 2010. Picante: R tools for integrating phylogenies and
719 ecology. Bioinformatics 26:1463–1464.
- 720 84. Oksanen J. 2007. vegan : Community Ecology Package. R package version 1.8-5.
- 721 85. Love MI, Huber W, Anders S. 2014. Moderated estimation of fold change and
722 dispersion for RNA-seq data with DESeq2. Genome Biol 15:550.

- 723 86. Langfelder P, Horvath S. 2008. WGCNA: an R package for weighted correlation
724 network analysis. *BMC Bioinformatics* 9:559.
- 725 87. Benjamini Y, Krieger AM, Yekutieli D. 2006. Adaptive linear step-up procedures that
726 control the false discovery rate. *Biometrika* 93:491–507.
- 727 88. Pollard KS, Dudoit S, van der Laan MJ. 2005. Multiple Testing Procedures: the
728 multtest Package and Applications to Genomics, p. 249–271. *In* Gentleman, R, Carey,
729 VJ, Huber, W, Irizarry, RA, Dudoit, S (eds.), *Bioinformatics and Computational*
730 *Biology Solutions Using R and Bioconductor*. Springer New York, New York, NY.
- 731 89. Shannon P, Markiel A, Ozier O, Baliga NS, Wang JT, Ramage D, Nada A,
732 Schwikowski B, Ideker T. 2003. Cytoscape: a software environment for integrated
733 models of biomolecular interaction network. *Genome Research* 13:2498–2504.
- 734 90. Bastian M, Heymann S, Jacomy M. 2009. Gephi : an open source software for
735 exploring and manipulating networks.

736

737 **Tables**

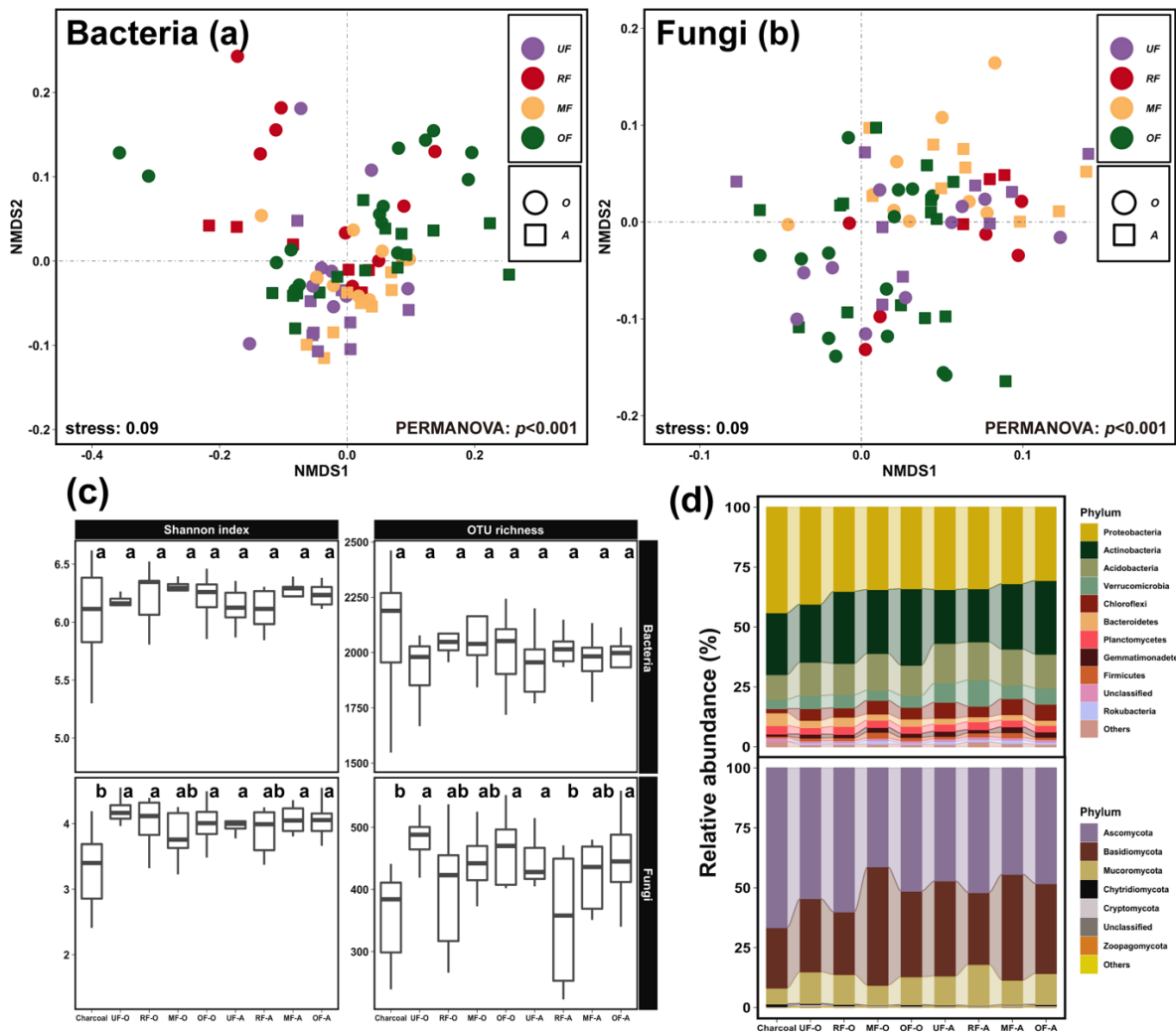
738 **Table 1.** Topological properties of networks obtained at the different fire histories in O
 739 horizon. Fire histories are recent fires (RF), middle-term fires (MF), old fires (OF) and
 740 unfired (UF).

741

Topological properties	UF	RF	MF	OF
<i>Bacteria</i>				
Number of nodes	64	60	72	222
Number of edges	57	30	62	380
Average degree	1.781	1	1.722	3.423
Modularity	0.845	0.967	0.907	0.737
Average clustering coefficient	0.723	NA	1.000	0.490
Average path length	1.68	1	1	4.052
Average number of neighbors	1.78	1	1.72	3.42
<i>Fungi</i>				
Number of nodes	67	48	16	152
Number of edges	83	38	8	253
Average degree	2.478	1.583	1	3.329
Network diameter	4	4	1	7
Modularity	0.807	0.828	0.875	0.917
Average clustering coefficient	0.702	0.661	NA ^a	0.689
Average path length	1.599	1.421	1	1.827
Average number of neighbors	2.48	1.58	1	3.33

742 ^aNA means not applicable.

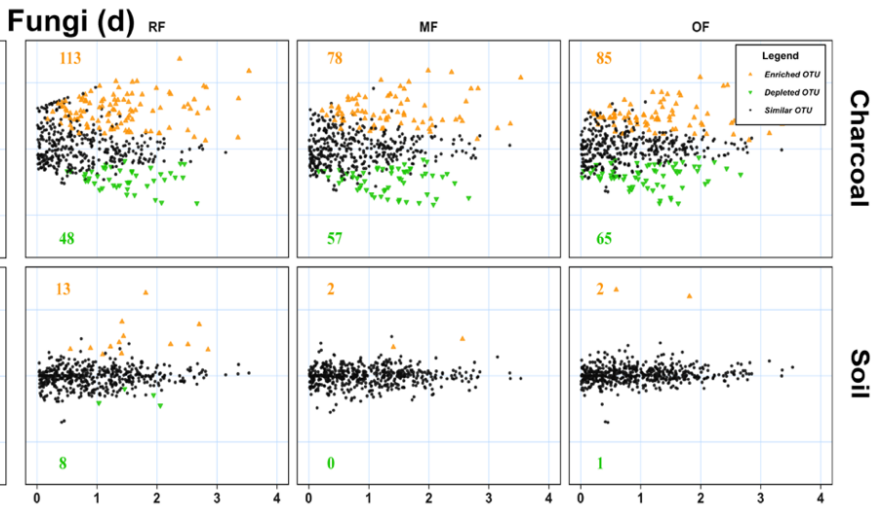
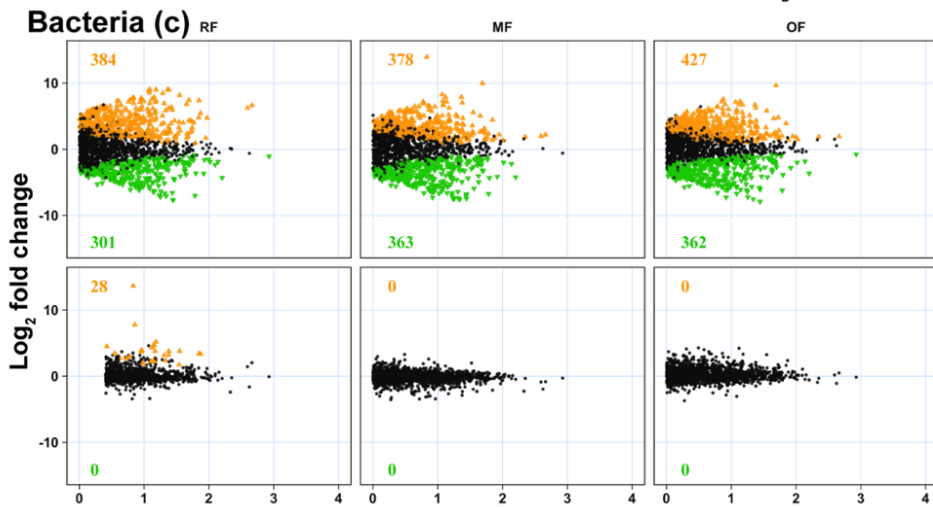
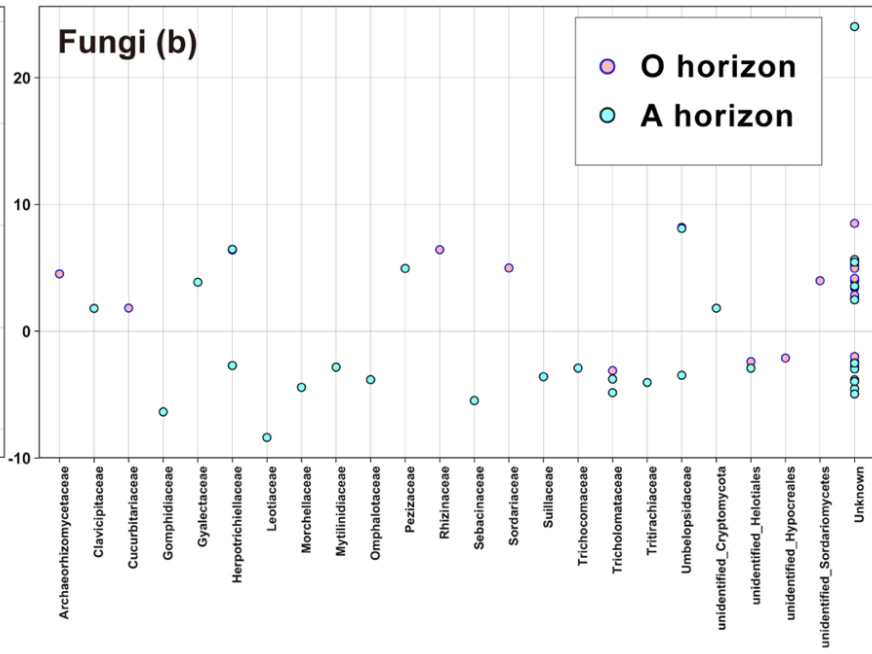
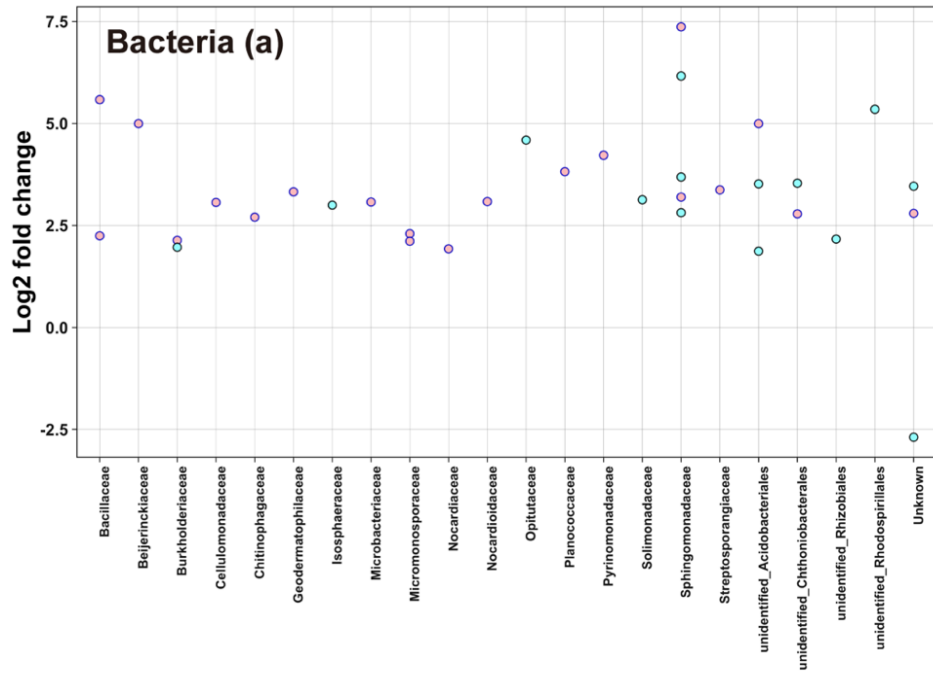
743 **Figures**



744

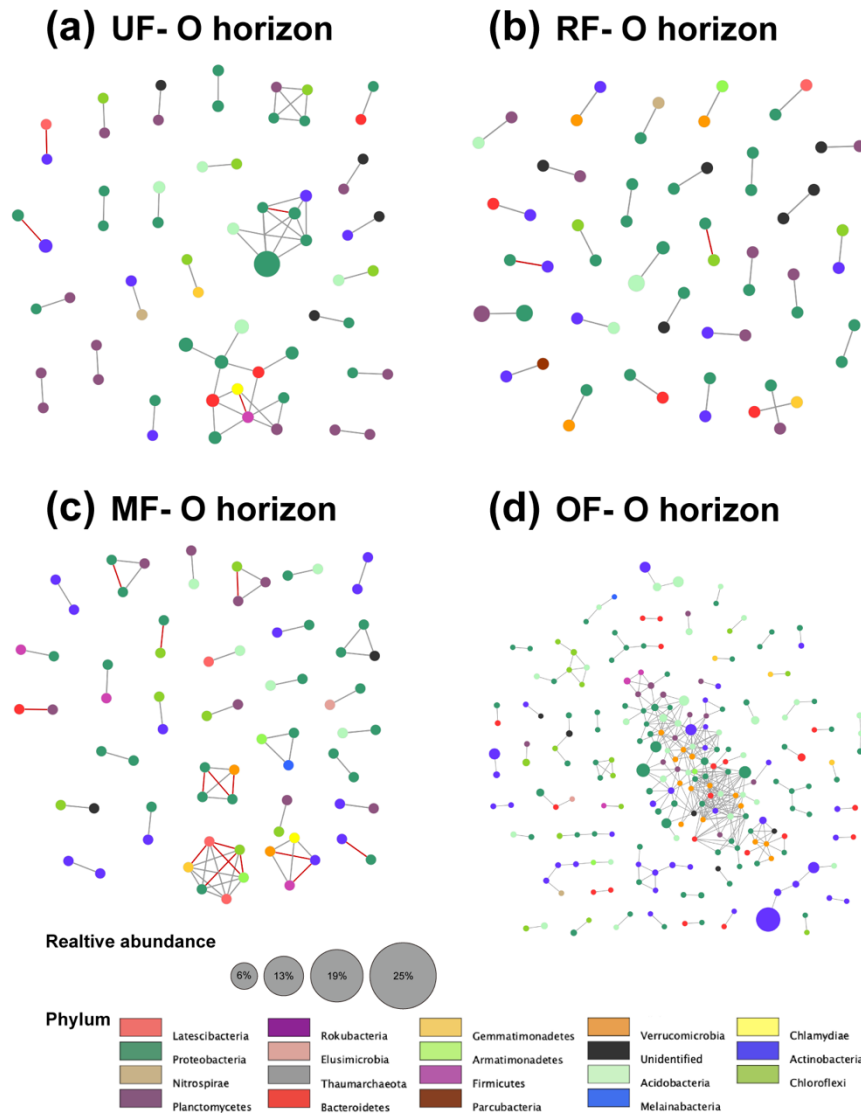
745 **Figure 1.** Microbial composition and structure in soils and charcoal samples of different fire
 746 histories and soil horizons. Non-metric multidimensional scaling ordination plot based on
 747 Bray-Curtis dissimilarity showing the change of soil bacterial (a) and fungal (b) community
 748 structure. Purple, red, orange and green symbols represent unburnt controls (UF), and recent
 749 (RF, 0-4 years), middle-term (MF, 7-15 years) and old fire (OF, 17-29 years), respectively.
 750 As listed in the legend, solid circles and squares represent samples from the O and A
 751 horizons, respectively. Boxplots of α -diversity (c) for bacterial and fungal communities in
 752 the soils UF-O and charcoal samples of different fire histories (UF, RF, MF and OF) in O and A
 753 horizons. Boxes are bounded on the first and third quartiles, divided by median lines. Boxes
 754 with different lower-case letters are significantly different ($p < 0.05$) by Bonferroni's post
 755 hoc tests. Relative abundances of bacteria and fungi among different fire histories in O and

756 A horizons at the phylum level (**d**). “Unclassified” OTUs indicate no taxonomic information
757 matched in the database. Charcoal samples only were separated from burnt soils in O
758 horizons.



Log₁₀ average OTU richness

760 **Figure 2.** Enrichment (above 0) and depletion (below 0) of bacterial **(a)** and fungal **(b)** OTUs in soils after recent fires (RF, 0-4 years) compared
761 with unburnt soils by differential abundance analysis. The OTUs (singleton is exclusive) are arranged by family on the x-axis and each point
762 represents an individual OTU. The y-axis indicates fold change in log base 2 units. Skyblue points represent differential OTUs found in O
763 horizon while pink points represent differential OTUs found in A horizon. The significant threshold is 0.01. Enrichment (yellow triangles) and
764 depletion (green triangles) of OTUs (singleton is exclusive) about bacterial community **(c)** and fungal community **(d)** by differential abundance
765 analysis. Enrichment and depletion of certain OTUs in the charcoal and soil samples of recent (RF, 0-4 years), middle-term (MF, 7-15 years) and
766 old fire (OF, 17-29 years) compared to unburnt soil samples (black dots) in O horizon. The significant threshold is 0.01. Colored numbers
767 represent the number of differential OTUs.



768

769 **Figure 3.** The co-occurrence networks of the bacterial community in O horizon under recent

770 fires (RF, **b**), middle-term fires (MF, **c**), old fires (OF, **d**) and unburnt groups (UF, **a**). A

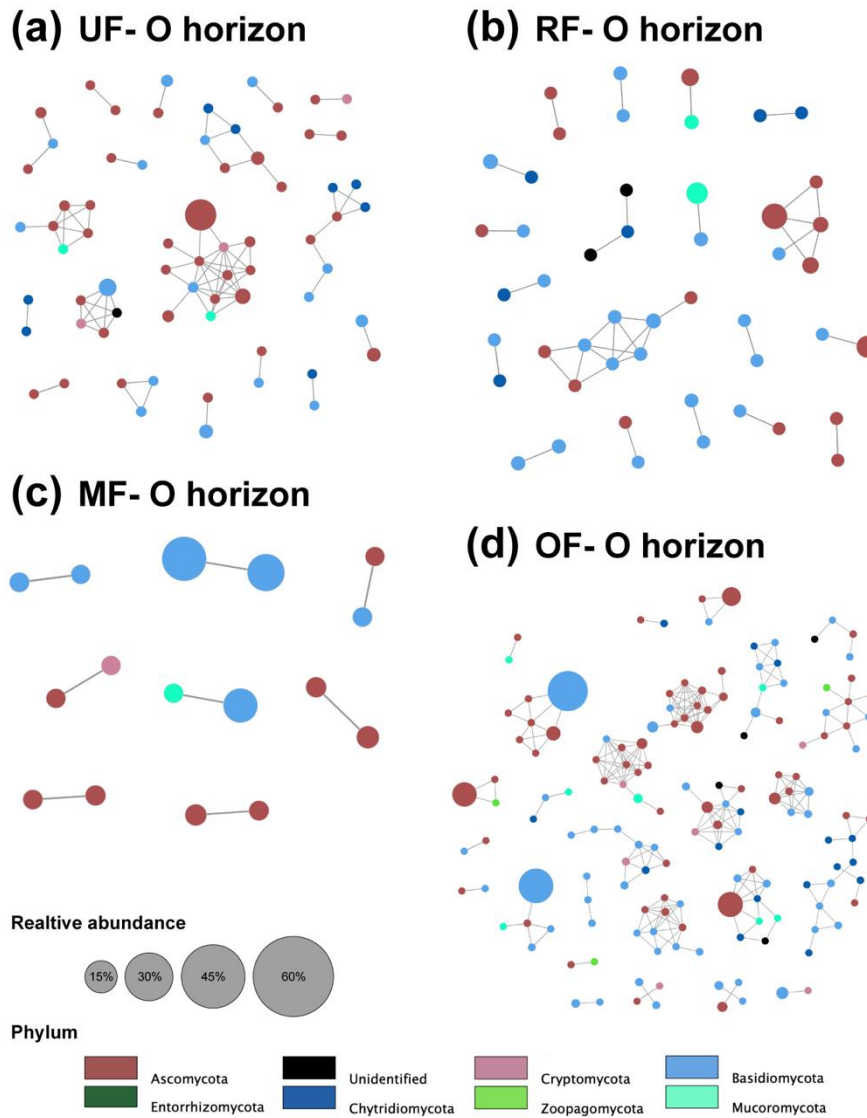
771 connection stands for a strong (Spearman's $p > 0.8$) and significant ($p < 0.01$) correlation.

772 The size of each node is proportional to the relative abundance, as is shown in the legend.

773 The thickness of each edge is proportional to the value of Spearman's correlation

774 coefficients. The grey edges represent positive interactions between two bacterial nodes,

775 while red edges represent negative interactions.



776

777 **Figure 4.** The co-occurrence networks of the fungal community in O horizon under recent

778 fires (RF, **b**), middle-term fires (MF, **c**), old fires (OF, **d**) and unburnt groups (UF, **a**). A

779 connection stands for a strong (Spearman's $p > 0.8$) and significant ($p < 0.01$) correlation.

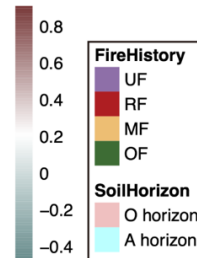
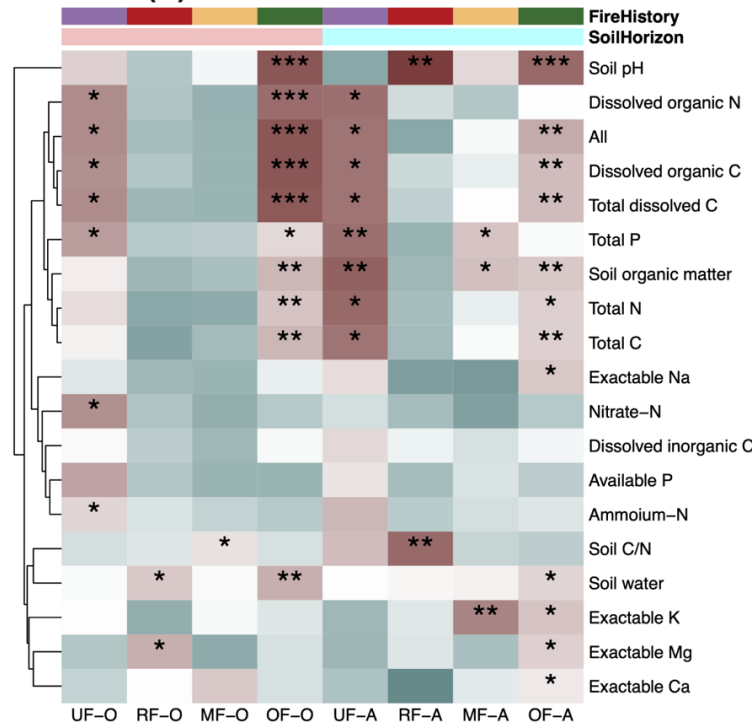
780 The size of each node is proportional to the relative abundance, as is shown in the legend.

781 And the thickness of each edge is proportional to the value of Spearman's correlation

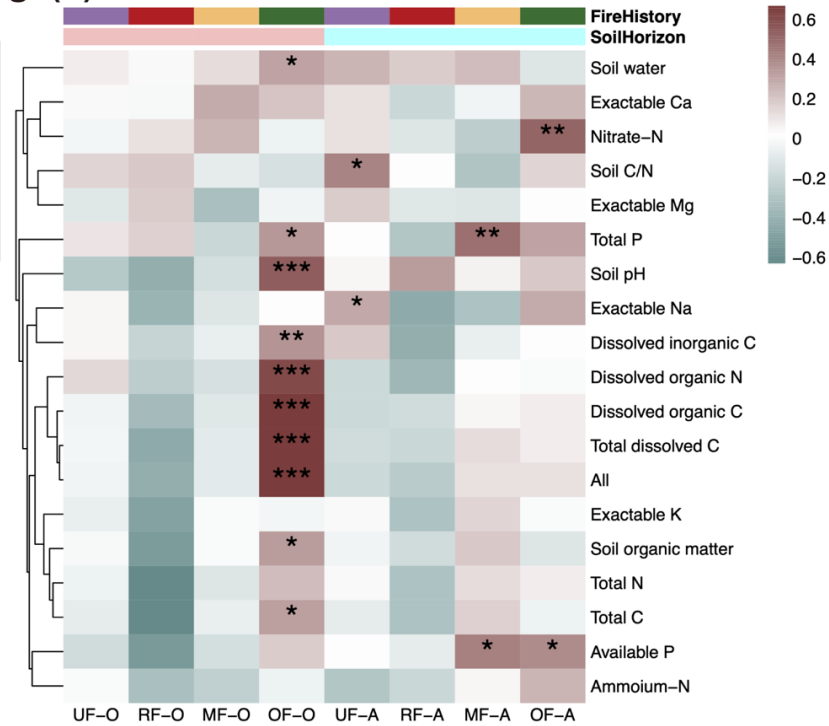
782 coefficients. All edges are grey, which means positive interactions between two fungal

783 nodes.

Bacteria (a)

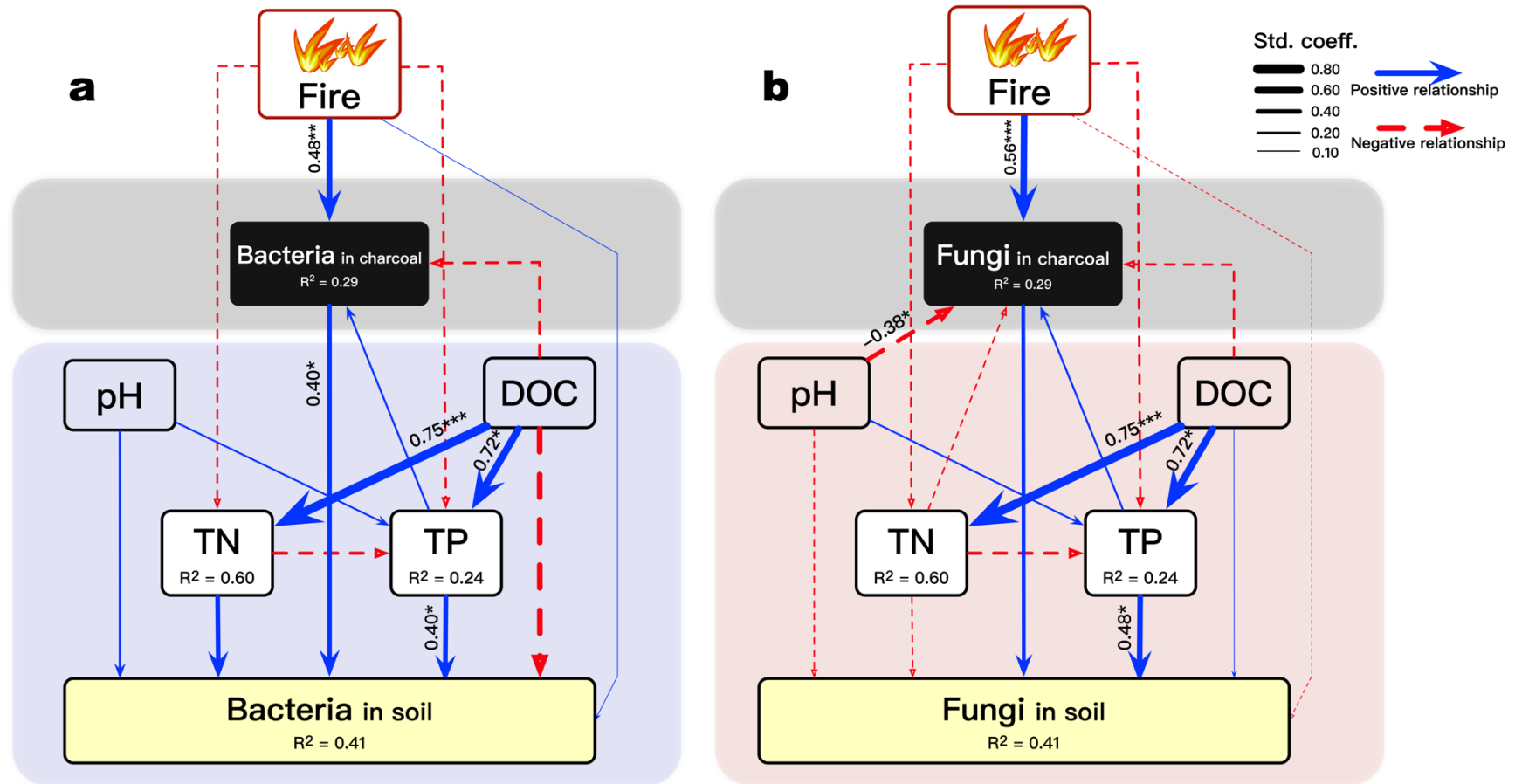


Fungi (b)



784

785 **Figure 5.** Associations of bacterial (a) and fungal (b) community structures with soil factors, by Partial Mantel test, in different fire histories and
 786 different horizons (O and A). Fire histories are recent fires (RF), middle-term fires (MF), old fires (OF) and unfired (UF). All, all environmental
 787 factors. The correlations (r) and significance (p) are determined with a Bray-Curtis distance matrix and a Euclidean distance matrix for
 788 environmental factors, controlled with spatial variation. Different colors in the cells indicate different mantel statistic correlations (r), shown in
 789 the legend. The p values in the cells with *, ** and *** indicate $p < 0.05$, $p < 0.01$, $p < 0.001$, respectively. Significance for each test is
 790 calculated from 9999 randomized Monte Carlo runs.



791

792 **Figure 6.** Fire legacy effect on bacterial (a) and fungal (b) community structures of the burnt soil in O horizon. Best-fitted structural equation
 793 models depicting the major pathways of fire, charcoal and representative environmental variables on microbial community structures. The
 794 bacterial and fungal community structures of charcoal and soil represented by the first axis of NMDS. DOC, dissolved organic carbon; TN, total
 795 nitrogen content; TP, total phosphorus content; Fire, time since fire. R^2 values denote the amount of variance explained by the model. Solid (blue)

796 and dashed arrows (red) indicate positive and negative effects, respectively. Arrow thickness is proportional to the standardized coefficients (std.
797 coeff.), interpreted as the relative importance of effects. Only significant ($*p < 0.05$, $**p < 0.01$, $***p < 0.001$) estimate numbers are noted at
798 each arrow. The goodness of fit was satisfactory ($n = 31$, p value = 0.483, CFI = 1.000, RMSEA = 0.000, SRMR < 0.015).

799 **Supplementary Information**

800 **Supplementary Methods**

801 *Additional site information*

802 *Soil analyses*

803 *Charcoal hand-picked method*

804 *References for Supplementary Methods*

805 **Supplementary Tables**

806 **Table S1.** Field site information of fire-affected and unaffected soils.

807 **Table S2.** Effects of fire history on soil properties in soil O and A horizons.

808 **Table S3.** Effects of fire history on the relative abundances (specific OTUs \times 100 / total
809 number) of bacterial and fungal phyla in soil O and A horizons.

810 **Table S4.** Results from permutational multivariate analysis of variance with Bray-Curtis
811 dissimilarity from 82 soil samples.

812 **Table S5.** Differential bacterial OTUs in the recent fires (RF) compared with unfired sites in
813 O and A horizons.

814 **Table S6.** Differential fungal OTUs in the recent fires (RF), middle-term fires (MF), old
815 fires (OF) compared with unfired sites in O and A horizons.

816 **Table S7.** Topological properties of networks obtained at the different fire histories and
817 different horizons (O and A).

818 **Table S8.** Bacterial and fungal community structure association with soil factors by Partial
819 Mental test in different fire histories and different horizons (O and A).

820 **Supplementary Figures**

821 **Figure S1.** The study region of 44 sites where samples were collected in the Chinese boreal
822 forest.

823 **Figure S2.** Soil properties for unburnt and burnt groups.

824 **Figure S3.** Non-metric multidimensional scaling ordination plot based on Bray-Curtis
825 dissimilarity showing changes in soil bacterial (a) and fungal (b) community structures.

826 **Figure S4.** All samples using NMDS ordination based on Bray-Curtis dissimilarity to depict
827 fire legacy effect on microbial communities.

828 **Figure S5.** The bacterial co-occurrence networks under recent fires (RF, b), middle-term
829 fires (MF, c), old fires (OF, d) and fire-unaffected groups (UF, a) in A horizon.

830 **Figure S6.** The fungal co-occurrence networks under recent fires (RF, b), middle-term fires
831 (MF, c), old fires (OF, d) and fire-unaffected groups (UF, a) in A horizon.

832 **Figure S7.** The general structure of *a priori* structural equation model including all possible
833 pathways (1-17).